Our schedule



Gas Dynamics (Chapter 14 in Leveque)

1 17 2 24	7 Aug 2009		торіс	Chapter in LeVeque
2 24		Monday 13.15-15.00	introduction to conservation laws, Clawpack	1&2&5
0.00	4 Aug 2009	Monday 13.15-15.00	the Riemann problem, characteristics	3
3 28	3 Aug 2009	Friday 13.15-15.00	finite volume methods for linear systems	4
4 8	3 Sep 2009	Tuesday 13.15-15.00	high resolution methods	6
5 2	I Sep 2009	Monday 13.15-15.00	boundary conditions, accuracy, variable coeff.	7,8, part of 9
6 29	9 Sep 2009	Tuesday 13.15-15.00	nonlinear conservation laws, finite volume methods	11 & 12
7	5 Oct 2009	Monday 13.15-15.00	nonlinear equations & systems	13, part of 14
8 1	2 Oct 2009	Monday 13.15-15.00	finite volume methods for nonlinear systems	part of 14, 15
9 1	9 Oct 2009	Monday 13.15-15.00	varying flux functions, source terms	part of 16, 17
10 2	6 Oct 2009	Monday 13.15-15.00	multidimensional hyperbolic problems & methods	18 & 19
11 2	2 Nov 2009	Monday 13.15-15.00	systems & applications; project planning	20 & 21
12 16	3 Nov 2009	Monday 13.15-15.00	applications: tsunamis, pockmarks, venting, impacts	
13 23	3 Nov 2009	Monday 13.15-15.00	applications: volcanic jets, pyroclastic flows, lahars	
14 30) Nov 2009	Monday 13.15-15.00	review; progress, problems & projects	
16 7	7 Dec 2009	Monday 13.15-15.00	FINAL PROJECT REPORTS DUE	

The barotropic equations and the shallow-water equations

$$\rho_t + (\rho u)_x = 0$$
$$\rho u_t^2 + (\rho u^2 + p)_x = 0$$

The conservation laws for the barotropic system (i.e. with $p = P(\rho)$)

are exactly like the shallow water equations if we identify ρ with h and use

$$p = P(\rho) = \frac{1}{2}g\rho^2$$

Other barotropic forms include the isothermal equation of state

 $p = P(\rho) = a^2 \rho,$

and the polytropic (or gamma-law) equation of state

$$p = P(\rho) = K \rho^{\gamma}$$

But next we add the energy equation...

the equation of state

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The Euler equations of gas dynamics

This is the full system of three conservation laws, for mass, momentum, and energy, for fully compressible gas dynamics:

$$q_t + f(q)_x = 0,$$

where

 $q = \begin{bmatrix} \rho \\ \rho u \\ E \end{bmatrix}, f(q) = \begin{bmatrix} \rho u \\ \rho u^2 + p \\ u(E+p) \end{bmatrix}.$

The total energy is composed of internal energy plus kinetic energy,

$$E = \rho e + \frac{1}{2}\rho u^2$$

and the system is completed by an equation of state $e = e(p, \rho)$.

The Jacobian f'(q) has eigenvalues u-c, c, u+c where the speed of

sound is $C = \sqrt{\frac{dp}{d\rho}}$ at constant entropy. FYS-GEO 4500 Galen Gisler, Physics of Geological Processes, University of Oslo

Entropy

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In the system of Euler equations for gas dynamics, we have the advantage of having an explicit formula for entropy that we can use as an entropy condition.

The specific entropy s (i.e. entropy per unit mass) is given by the formula:

$$s = c_v \log\left(\frac{p}{\rho^{\gamma}}\right) + \text{constant}$$

The additive constant is unimportant and may be omitted, since the important thing to keep track of is changes in entropy. In smooth flow, entropy is constant; at shocks it jumps to a higher value.

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The equation of state and associated relations for a polytropic gas

The ideal gas law:	$\frac{p}{\rho} = nkT$						
The internal energy:	$e = c_v T = \frac{\alpha}{2} n$	$kT = \frac{p}{(\gamma - 1)\rho}$					
The <i>enthalpy</i> :	$h = e + \frac{p}{\rho} = c_{j}$	$_{p}T = \left(1 + \frac{\alpha}{2}\right)nkT$					
Relations between the specific heats: $\begin{cases} c_p - c_v = nk \\ \gamma = \frac{c_p}{c_v} = \left(\frac{\alpha + 2}{\alpha}\right) \end{cases}$							
n: number of molecules per unit mass k : Boltzmann's constant							
c_v : specific heat at co	nstant volume	lpha : number of degrees of free	edom				
c_p : specific heat at co	nstant pressure	γ : adiabatic exponent					
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Primitive variables

It is often useful to examine the equivalent equations in directly observable "primitive" variables, rather than the conserved variables.

The Euler equations in primitive form for a polytropic gas:

$$\rho_t + u\rho_x + \rho u_x = 0$$
$$u_t + uu_x + \frac{1}{\rho}p_x = 0$$

 $p_t + \gamma p u_x + u p_x = 0$

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where

are:

 $\lambda^1 = u - c$

is the speed of sound in the polytropic gas.

Then the eigenvalues and eigenvectors

 $\lambda^2 = u$

 $r^{1} = \begin{bmatrix} -\rho/c \\ 1 \\ -\rho c \end{bmatrix} r^{2} = \begin{bmatrix} 1 \\ 0 \\ 0 \end{bmatrix} r^{3} = \begin{bmatrix} \rho/c \\ 1 \\ \rho c \end{bmatrix}$

 $\lambda^3 = u + c$

in matrix notation:



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The Jacobian for the conservation laws This (for the polytropic gas) is slightly more complex, though equivalent: $f'(q) = \begin{bmatrix} 0 & 1 & 0\\ \frac{1}{2}(\gamma-3)u^2 & (3-\gamma)u & \gamma-1\\ \frac{1}{2}(\gamma-1)u^3 - uH & H - (\gamma-1)u^2 & \gamma u \end{bmatrix}$ where $H = \frac{E+p}{\rho} = h + \frac{1}{2}u^2$ is the total specific enthalpy. And the eigenvalues and eigenvectors are: $\lambda^1 = u - c \qquad \lambda^2 = u \qquad \lambda^3 = u + c$ $r^1 = \begin{bmatrix} 1\\ u - c\\ H - uc \end{bmatrix} r^2 = \begin{bmatrix} 1\\ u\\ \frac{1}{2}u^2 \end{bmatrix} r^3 = \begin{bmatrix} 1\\ u+c\\ H + uc \end{bmatrix}$

A wave (or field, we may say, referring to the collection of waves of the same

Genuine nonlinearity, linear degeneracy

family in all accessible space) is genuinely nonlinear if

$$\nabla \lambda^p \cdot r^p(q) \neq 0$$
 for all q

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Physically this means that the characteristics are either compressing or expanding.

The opposite case is linear degeneracy,

$$\nabla \lambda^p \cdot r^p(q) = 0$$
 for all q

in this case the characteristics are parallel to one another.

Genuine nonlinearity, linear degeneracy





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The Riemann invariants for the polytropic gas

Thus, of the three eigenvectors, 1 and 3 represent waves that can become either rarefactions or shocks, while 2 is linearly degenerate and can only be a contact discontinuity.

For any *simple wave* (not a rarefaction or a shock), the Riemann invariants are constant along particle paths through the wave. These are, for the 3 waves:



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a membrane into sections with different

The membrane is suddenly removed, and

region into the lower density region.

Three waves develop: a shock wave, a

contact discontinuity, and a rarefaction

wave (or fan). The first two travel to the

continuous, but their derivatives are not.

At the shock, velocity, pressure and density are all discontinuous. At the contact, only density is discontinuous. In the rarefaction fan, all variables are

right, the third to the left.

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densities.



x

The Riemann problem for 3 waves: rarefaction contact shock t_{n+1} ρ_l ρ_r u^* u^* p^* $p^{\tilde{}}$ ρ_{l} ρ_r u_i u_r p_l p_r Q_{i-1}^n Q_i^n Ē $x_{i-1/2}$

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The 2-field is linearly degenerate. Across the contact *u* and *p* will be constant and only o will jump.

The strategy for solving the problem is to use the Hugoniot loci and integral curves for the 1-field and 3-field, in the phase plane of *u* and *p*, in the same way as for the shallow-water equations to obtain (p^*, u^*) . Then we calculate the densities on either side of the contact. Finally we solve for the solution within the rarefaction fan.

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The general (exact) Riemann solver for the Euler equations for a polytropic gas

As before, we define functions

$$u = \varphi_{l}(p) = \begin{cases} u_{l} + \frac{2c_{l}}{\gamma - 1} \left[1 - \left(p / p_{l} \right)^{\frac{\gamma - 1}{2\gamma}} \right] & \text{if } p \leq p_{l} \\ u_{l} + \frac{2c_{l}}{\sqrt{2\gamma(\gamma - 1)}} \left[\frac{1 - p / p_{l}}{\sqrt{1 + \beta p / p_{l}}} \right] & \text{if } p \geq p_{l} \end{cases}$$
$$u = \varphi_{r}(p) = \begin{cases} u_{r} - \frac{2c_{r}}{\gamma - 1} \left[1 - \left(p / p_{r} \right)^{\frac{\gamma - 1}{2\gamma}} \right] & \text{if } p \leq p_{r} \\ u_{r} - \frac{2c_{r}}{\sqrt{2\gamma(\gamma - 1)}} \left[\frac{1 - p / p_{r}}{\sqrt{1 + \beta p / p_{r}}} \right] & \text{if } p \geq p_{r} \end{cases}$$
where $\beta = \frac{\gamma + 1}{\gamma - 1}$.

We then require that $\varphi_l(p_m) = \varphi_r(p_m)$, using an iterative procedure to find the intersection (p^*, u^*) of the curves. The densities on either side of the contact will then be given by

$$\rho_l^* = \left(\frac{1+\beta p^*/p_l}{p^*/p_l+\beta}\right)\rho_l; \quad \rho_r^* = \left(\frac{1+\beta p^*/p_r}{p^*/p_r+\beta}\right)\rho_r$$

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and solve for u as a function of ξ .

Hugoniot loci and integral curves for the Euler Equations (polytropic gas)

Some integral curves (solid) and Hugoniot loci (dotted) for the Euler equations. Just as in the shallow water equations, these curves are close together in many places. An iterative solver can start from the intersection of the integral curves, which can be obtained explicitly.

These curves are computed for γ =1.4 and

densities $Q_{l}=3$ and

 Q_r =1. For lower γ and higher density contrasts, the curves spread further apart.

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We now have everything except the rarefaction:



Then we have that $\xi = \lambda^1 = u - c$ within the rarefaction wave, so we can rewrite the Riemann invariant as

$$u+\frac{2(u-\xi)}{\gamma-1}=u_{l}+\frac{2c_{l}}{\gamma-1},$$

and solve for u as a function of ξ .

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Multifluid problems and other equations of state

The easiest multi-fluid case is when you have two ideal gases with different values of γ . Then you set up the Riemann problem at the interface between the two fluids with right and left values for γ . Things get complicated when mixing occurs.

Other analytical or tabular equations of state can also be incorporated into a finite volume conservative scheme. Sage, for example, uses the Sesame library of tabular equations of state for lots of materials. The Sesame library, developed at Los Alamos National Laboratory, contains mostly industrial materials, with a few materials of geological interest.

Leveque gives lots of references to papers in which Riemann solvers are developed for other equations of state in his section 14.15. Some of these will be worth looking into for geological applications.

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An equation of state for a dusty gas*

One version of a dusty-gas equation of state is: $p = \frac{(1-K)}{(1-T)}\rho RT$,

where K is the mass concentration and Z is the volume fraction of solid particles. These are related through $K = \frac{Z\rho_s}{\rho}$, where ρ_s is the particle solid density.

The speed of sound is then $c_{ds} = \sqrt{\frac{\Gamma p}{2(1-T)}}$

f specific heats for the mixture is
$$\Gamma = \frac{\gamma C_{\nu}(1-K) + C_{sp}K}{C_{sp}(1-K) + C_{-K}K}$$
.

The specific heat of the dust particles is C_{sp} and C_{v} is the specific heat at constant volume of the gas.

> *from Vishwakarma, Nath, & Singh (2008), Physica Scripta 78 035402. http://stacks.iop.org/PhysScr/78/035402

where the ratio of











Vents, kimberlite pipes, and geysers may be natural deLaval nozzles for a dusty gas



Shocks can be stronger in dusty gases

0.50

0.75 p / p_0

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The maximum ratio of upstream to downstream densities across a shock is (see Landau & Lifshitz, Fluid Dynamics)

$$\frac{\rho_u}{\rho_d} = \frac{\gamma + 1}{\gamma - 1}$$

For a diatomic gas (like air), $\gamma = 1.4$, so $\frac{\rho_u}{\rho_d} = 6.4$

For a dusty gas,
$$\gamma \Rightarrow 1$$
, so $\frac{\rho_u}{\rho_d}$ can be arbitrarily large. At $\gamma = 1.01$, $\frac{\rho_u}{\rho_d} = 201$.



Once again, we extend from what we've learned for linear systems of equations

We intend to solve the nonlinear conservation law $q_t + f(q)_x = 0$

using a method that is in conservative form:

$$Q_{i}^{n+1} = Q_{i}^{n} - \frac{\Delta t}{\Delta x} \left(F_{i+1/2}^{n} - F_{i-1/2}^{n} \right)$$

and yielding a *weak solution* to this conservation law. To get the *correct* weak solution we must use an appropriate entropy condition.

Finite Volume Methods for Nonlinear Systems (Chapter 15 in Leveque)

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Recall Godunov's method: Given a set of cell quantities Q_i^n at time *n*: 1. Solve the Riemann problem at $x_{i-1/2}$ to obtain $Q_{i-1/2}^{\downarrow} = q^{\downarrow}(Q_{i-1}^n, Q_i^n)$ 2. Define the flux: $F_{i-1/2}^n = f(Q_{i-1/2}^{\downarrow})$ 3. Apply the flux differencing formula: $Q_i^{n+1} = Q_i^n - \frac{\Delta t}{\Delta x} (F_{i+1/2}^n - F_{i-1/2}^n)$ This will work for any general system of conservation laws. Only the formulation of the Riemann problem itself changes with the system.

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Recall what TVD means:

Monotonicity preserving methods:

If a grid function that is initially monotone, i.e. $Q_i^n \ge Q_{i-1}^n$ for all *i* at step *n* $Q_i^{n+1} \ge Q_{i-1}^{n+1}$ for all *i* at step n+1remains monotone at the next time: then the method is monotonicity preserving. Total Variation Diminishing (TVD) methods: Define the total variation of a grid function Q as: $TV(Q) = \sum_{i=1}^{n} |Q_i - Q_{i-1}|$ $\mathrm{TV}(Q^{n+1}) \leq \mathrm{TV}(Q^n)$ A method is Total Variation Diminishing if If Q^n is monotone, then so is Q^{n+1} , and no spurious oscillations are generated. This gives a form of stability necessary for proving convergence, also for nonlinear conservation laws. FYS-GEO 4500 Galen Gisler, Physics of Geological Processes, University of Oslo Autumn 2004 Monday, 12 October 2009

Slope limiters and flux limiters

The slope limiter formula for advection is:

$$Q_i^{n+1} = Q_i^n - \frac{u\Delta t}{\Delta x} \left(Q_i^n - Q_{i-1}^n \right) - \frac{1}{2} \frac{u\Delta t}{\Delta x} \left(\Delta x - u\Delta t \right) \left(\sigma_i^n - \sigma_{i-1}^n \right)$$

The flux limiter formulation for advection is:

$$Q_{i}^{n+1} = Q_{i}^{n} - \frac{\Delta t}{\Delta x} \left(F_{i+1/2}^{n} - F_{i-1/2}^{n} \right)$$

with the flux:

$$F_{i-1/2}^{n} = uQ_{i-1}^{n} + \frac{1}{2}u(\Delta x - u\Delta t)\sigma_{i-1}^{n}$$

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High-resolution wave-propagation scheme

The fluctuation notation is more useful in the nonlinear case:

$$\begin{aligned} Q_{i}^{n+1} &= Q_{i}^{n} - \frac{\Delta t}{\Delta x} \Big(\mathcal{A}^{-} \Delta Q_{i+1/2} + \mathcal{A}^{+} \Delta Q_{i-1/2} \Big) - \frac{\Delta t}{\Delta x} \Big(F_{i+1/2} - F_{i-1/2} \Big) \\ \text{where} \qquad F_{i-1/2}^{n} &= \frac{1}{2} \sum_{p=1}^{m} \Big| s_{i-1/2}^{p} \Big| \Big(1 - \frac{\Delta t}{\Delta x} \Big| s_{i-1/2}^{p} \Big| \Big) \widetilde{\mathcal{W}}_{i-1/2}^{p} \end{aligned}$$

 $\widetilde{\mathcal{W}}_{i-1/2}^{p}$ represents the limited version of $\mathcal{W}_{i-1/2}^{p}$.

This is obtained by comparing $\mathcal{W}_{i-1/2}^p$ with $\mathcal{W}_{i-1/2}^p$ where

$$l = \begin{cases} i - 1 \text{ if } s_{i-1/2}^p > 0\\ i + 1 \text{ if } s_{i-1/2}^p < 0 \end{cases}$$

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Wave propagation for nonlinear systems

An **approximate** Riemann solver is typically used to get the wave decomposition

$$Q_i - Q_{i-1} = \sum_{p=1}^m \mathcal{W}_{i-1/2}^p,$$

where the wave $\mathcal{W}_{i-1/2}^p$ propagates at a speed $s_{i-1/2}^p$.

If we define $\hat{A}_{i-1/2} = \hat{A}(Q_i^n, Q_{i-1}^n)$ as a linearised approximation to f'(q) valid in the neighbourhood of (Q_{i}, Q_{i-1}) ,

then we can solve the simpler linear Riemann problem at that cell interface for the linearised equation:

 $q_t + \hat{A}_{i-1/2} q_x = 0,$

to obtain

$$\mathcal{W}_{i-1/2}^{p} = \alpha_{i-1/2}^{p} \hat{r}_{i-1/2}^{p}, \ s_{i-1/2}^{p} = \hat{\lambda}_{i-1/2}^{p}.$$

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Riemann solvers in CLAWPACK

In CLAWPACK, the hyperbolic problem is specified by providing a **Riemann** solver with

Input: the value of *q* in each grid cell

Output: the solution to the Riemann problem at each cell interface:

The Waves \mathcal{W}^p , p = 1, 2, ..., m for a system of *m* equations

The Speeds s^p , p = 1, 2, ..., m

The Fluctuations $\mathcal{A}^{\pm}\Delta Q$, for high-resolution corrections

Because the problem is solve entirely using Riemann solvers, you won't see anything in the code that resembles the original system of partial differential equations.

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Approximate Riemann Solvers

Approximate the true Riemann solution by a set of waves consisting of finite jumps propagating at *constant* speeds (as in the linear case).

Use a *local linearisation*: replace $q_t + f(q)_x = 0$ by $q_t + \hat{A}_{i-1/2}q_x = 0$, where $\hat{A} = \hat{A}(q_t, q_r) \approx f'(q_{ave})$.

Then decompose $q_l - q_r =$

$$q_l - q_r = \sum_{p=1}^m \alpha^p \hat{r}^p$$

to obtain the waves $\mathcal{W}^p = \alpha^p \hat{r}^p$ with speeds $s^p = \hat{\lambda}^p$.

But how do we chose \hat{A} ?

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Approximate Riemann Solvers

Properties desired for \hat{A} :

 \hat{A} must be diagonalisable with real eigenvalues

$$\hat{A}_{i-1/2} \to f'(q) \text{ as } Q_{i-1}, Q_i \to q$$

 $\hat{A}_{i-1/2}(Q_i - Q_{i-1}) = s(Q_i - Q_{i-1}) = f(Q_i) - f(Q_{i-1})$

With these properties, the method will be conservative, it will give the right answer across shocks (why?), and it is a good approximation for smooth flow.

We could take $\hat{A}_{i-1/2} = f'(\hat{Q}_{i-1/2})$ where $\hat{Q}_{i-1/2}$ is a suitable average of (Q_{i-1}, Q_i) . We'll use basically this approach.

Or we could take
$$\hat{A}_{i-1/2} = \frac{1}{2} \Big[f'(Q_{i-1}) + f'(Q_i) \Big]$$

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Roe solver for Shallow Water
Given
$$h_l, u_l, h_r, u_r$$
, define $\overline{h} = \frac{h_l + h_r}{2}$, $\hat{u} = \frac{\sqrt{h_l}u_l + \sqrt{h_r}u_r}{\sqrt{h_l} + \sqrt{h_r}}$
Then if \hat{A} is defined as the Jacobian matrix evaluated at the special state $\hat{q} = (\overline{h}, \overline{h}\hat{u})$,
we find that:
the Roe conditions are satisfied.

an isolated shock is modelled well,

and the wave propagation algorithm is conservative.

If we use limited waves, we obtain high-resolution methods as before.



Roe solver for Shallow Water Given h_l, u_l, h_r, u_r , define $\overline{h} = \frac{h_l + h_r}{2}$, $\hat{u} = \frac{\sqrt{h_l}u_l + \sqrt{h_r}u_r}{\sqrt{h_l} + \sqrt{h_r}}$ Then if \hat{A} is defined as the Jacobian matrix evaluated at the special state $\hat{q} = (\overline{h}, \overline{h}\hat{u}),$ we find that: the Roe conditions are satisfied, an isolated shock is modelled well, and the wave propagation algorithm is conservative.

If we use limited waves, we obtain high-resolution methods as before.

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The Harten-Lax-van Leer (HLL) Solver

This solver uses only 2 waves with

 s^1 = minimum characteristic speed

 s^2 = maximum characteristic speed

Write

$$\mathcal{W}^1 = Q^* - q_l, \quad \mathcal{W}^2 = q_r - Q^*$$

where the middle state $\boldsymbol{Q}^{^{*}}$ is uniquely determined by the conservation requirement:

$$s^{1}\mathcal{W}^{1} + s^{2}\mathcal{W}^{2} = f(q_{r}) - f(q_{l})$$

$$\Rightarrow Q^{*} = \frac{f(q_{r}) - f(q_{l}) - s^{2}q_{r} + s^{1}q_{l}}{s^{1} - s^{2}}$$

Modifications of this include positivity constraints and the addition of a third wave.

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$$\begin{array}{l} \textbf{The Shallow-water Riemann solver in constraints a result of the second secon$$

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f-wave approximate Riemann solver

Instead of splitting Q into waves, we might consider splitting the flux f into M_w "waves" $(M_w \le w)$:

$$f(Q_i) - f(Q_{i-1}) = \sum_{p=1}^{m_w} \mathcal{Z}_{i-1/2}^p$$

It turns out this is useful for spatially varying flux functions, i.e.

$$q_t + f(q, x)_x = 0,$$

with applications, for example, in:

wave propagation in heterogeneous nonlinear media,

flow in heterogeneous porous media,

traffic flow with varying road conditions,

conservation laws on curved manifolds,

and certain kinds of source terms.

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f-wave approximate Riemann solver Let \hat{A} be any averaged Jacobian matrix, for example: $\hat{A} = f'\left(\frac{q_i + q_r}{2}\right)$ Use eigenvectors of \hat{A} to do *f*-wave splitting. Then $\mathcal{A}^- \Delta Q_{i+1/2} + \mathcal{A}^+ \Delta Q_{i-1/2} = f(Q_i) - f(Q_{i-1})$, so the method is conservative. If \hat{A} is the Roe average, then this is equivalent to the normal Roe Riemann $\mathcal{Z}^p = s^p \mathcal{W}^p$.

Assignment for next time

CAMBRIDGE TEXTS IN APPLIED MATHEMATICS Finite Volume Methods for Hyperbolic Problems

Read Chapter 14 and Chapter 15.

Write (in Fortran) an approximate Riemann solver for the Euler equations using the Roe average. Test it on the shock tube problem, or (optionally) on the Woodward-Collella blast-wave problem. Use the shallow-water Riemann solver as a guide.

RANDALL J. LEVEQUE

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Hand in your code and results to me by Monday 19 October.

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