Water

GEF4400 "The Earth System"

Prof. Dr. Jon Egill Kristjansson,

Prof. Dr. Kirstin Krüger (UiO)

Email: kkruegergeo.uio.no

Lecture/ interactive seminar/ field excursion
 Teaching language: English

Time and location: Monday 12:15-14:00

Wednesday 10:15-12:00, CIENS Glasshallen 2.

Study program

Master of meteorology and oceanography
PhD course for meteorology and oceanography students

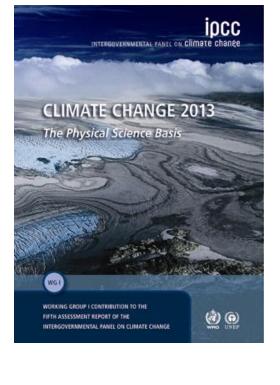
Credits and conditions:

The successful completion of the course includes an **oral presentation (weight 50%)**, a **successful completion of the Andøya field excursion** (mandatory), a **field report**, as well as a final **oral examination (50%)**. Student presentations will be part of the



IPCC Chapter 3: Observations: Ocean

- Background
- Introduction (Appendix 3A)
- Ocean temperature and heat content (Section 3.2)
- Salinity and fresh water content (Section 3.3)
- Ocean surface fluxes (Section 3.4)
- Ocean circulation (Section 3.6)
- Sea level change (Section 3.7)
- Executive Summary (Ch. 3)

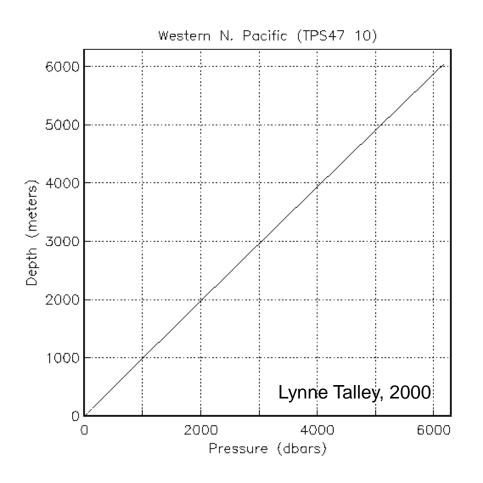


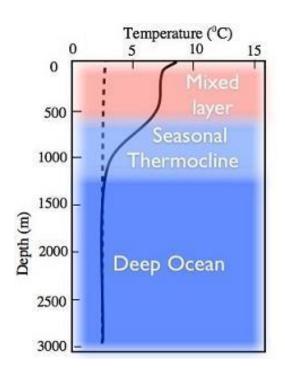
Rhein, M., et al., 2013: Observations: Ocean. In: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change. Cambridge University Press.



Background

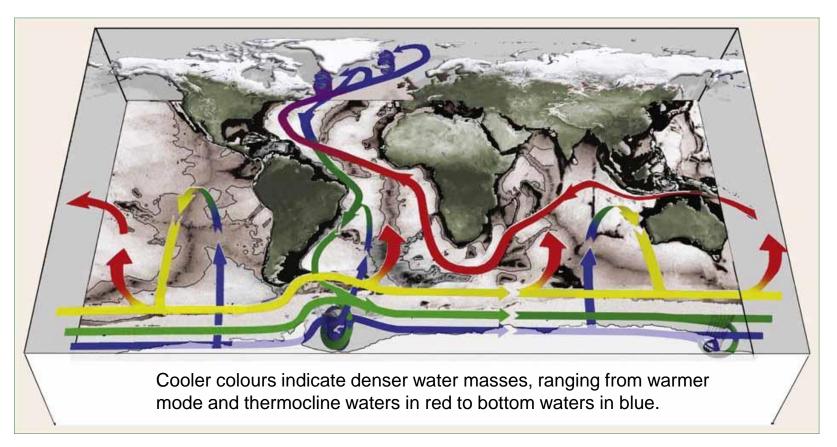
Ocean vertical structure (Tropical Oceans)





Ocean pressure is usually measured in decibars because the pressure in decibars is almost exactly equal to the depth in meters.

Ocean circulation



Marshall and Speer (2012, Nature)

Meridional Overturning Circulation (MOC) schematic driven mainly by the difference in heat, salinity, wind and eddies. In the early schematic of the conveyer belt analogy by Broecker (1991) the role for the Southern Ocean was neglected.

Surface Ocean Currents

- Surface ocean currents are driven by the circulation of wind above surface waters, interacting with evaporation, sinking of cold water at high latitudes, and the Coriolis force generated by the earth's rotation. Frictional stress at the interface between the ocean and the wind causes the water to move in the direction of the wind.
- Large surface ocean currents are a response of the atmosphere and ocean to the flow of energy from the tropics to polar regions.
- On a global scale, large ocean currents are constrained by the continental masses found bordering the three oceanic basins.
 Continental borders cause these currents to develop an almost closed circular pattern called a gyre
 - closed circular pattern called a gyre.
- Each ocean basin has a large gyre located at approximately 30° North/South. The currents in these gyres are driven by atmospheric flow produced by subtropical high pressure systems.



Ocean observations

In-situ: buoys, Argo floats, gliders, mooring, ships, ROV

Remote sensing: satellite, aircraft, radar

Observed ocean properties

- Sea Surface Temperature (SST): satellite and in-situ
- Sea Surface Salinity (SSS): in-situ
- Sea surface wind (stress): satellite and in-situ
- Sea level height: satellite and in-situ
- Ocean current: in-situ
- Ocean colour (chlorophyll): satellite and in-situ
- Air-sea fluxes (Carbon): in situ
- Sea ice: satellite and in-situ

Introduction and Motivation to Chapter 3

Why do we care about the Oceans influence on climate?

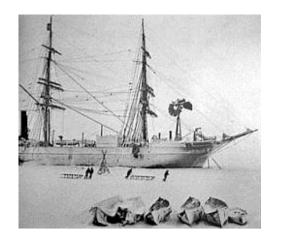
- Storing and transporting large amounts of heat, freshwater, and carbon; exchanging these properties with the atmosphere.
- ~93% of the excess heat energy stored in the ocean over last 50 yrs;
- >3/4 of total exchange of water (evaporation, precipitation) takes
 place over the oceans;
- 50 times more carbon than in the atmosphere, presently absorbing about 30% of human emissions of carbon dioxide (CO₂);
- ocean changes may result in climate feedbacks that either increase or reduce the rate of climate change;
- large inertia of the oceans means can provide a clearer signal of longer-term change than other components of the climate system.
- →Observations of ocean change to track the evolution of climate change, and a relevant benchmark for climate models.

Oceanography expeditions



- Early oceanography expeditions in the 1870s (e.g. *Challenger* voyage around the world);
- Arctic and Antarctic explorations (1893 to 1912) with *Fram*;



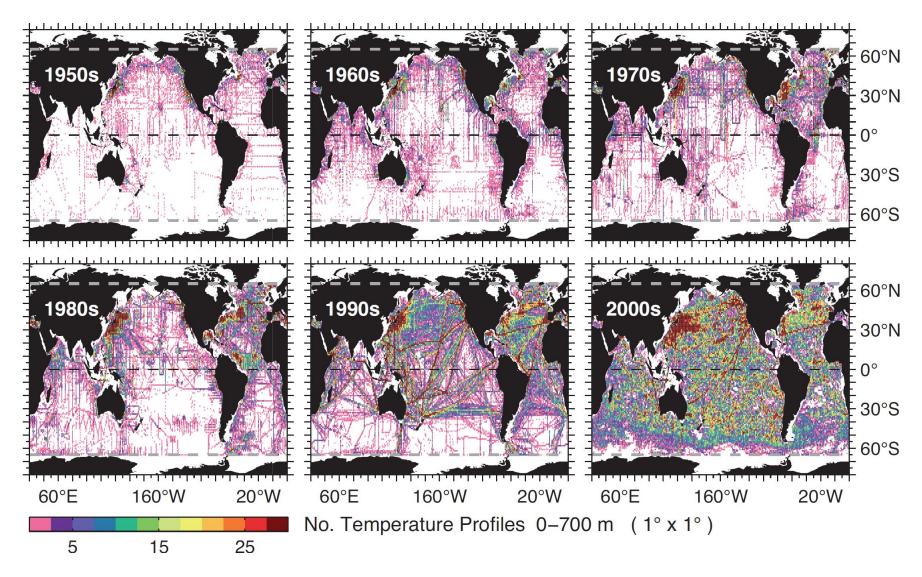


- Meteor survey to the Atlantic in the 1920s;
- Discovery investigations to the Southern Ocean in the 1920s.

With the International Geophysical Year (IGY) in 1957/58 a more frequent sampling began.



Ocean observations evolution



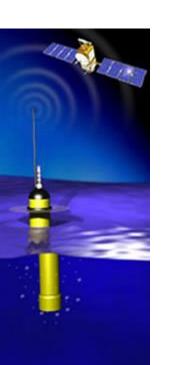
Ocean observations evolution



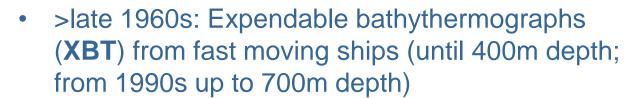
 Reversing thermometers and Nansen bottles from ships on stations



1960s: Conductivity-Temperature-Depth (CTD) casts with Niskin bottles



 1950s-1970s: subsurface measurements with mechanical bathythermographs from slow moving ship









Ocean observations coverage

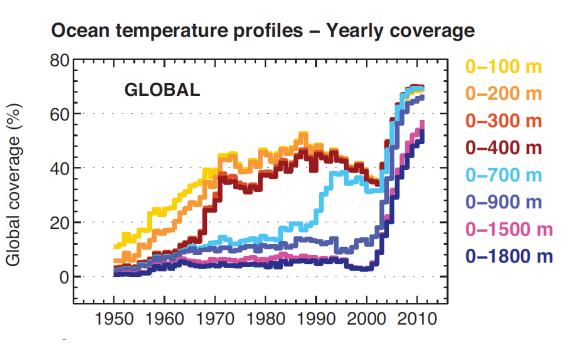
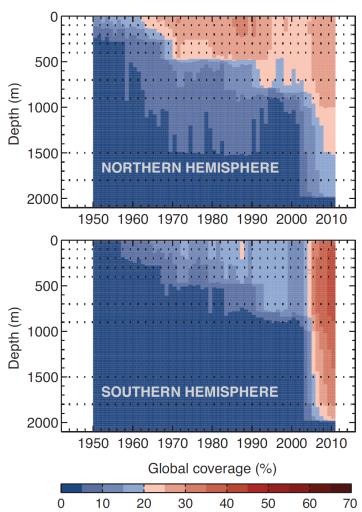


Figure 3.A.2 | (Top) Percentage of global coverage of ocean temperature profiles as a function of depth in 1° latitude by 1° longitude by 1-year bins (top panel) shown versus time. Different colours indicate profiles to different depths (middle panel). Percentage of global coverage as a function of depth and time, for the Northern Hemisphere. (Bottom panel) As above, but for the Southern Hemisphere.



Ocean observations improvements since AR4

Lack of long-term ocean measurements \rightarrow documenting and understanding oceans changes is an ongoing challenge.

Since AR4, substantial progress has been made in improving the **quality** and **coverage** of **ocean observations**:

- Biases in historical measurements have been identified and reduced, providing a clearer record of past change.
- Argo floats have provided near-global, year-round measurements of temperature and salinity in the upper 2000 m since 2005.
- Satellite altimetry record is now >20 years in length.
- Longer continuous time series of the meridional overturning circulation and tropical oceans have been obtained.
- Spatial and temporal coverage of biogeochemical measurements in the ocean has expanded.
- →Understanding ocean change has improved.

3.2 Ocean temperature and heat content

"Temperature is the most often measured subsurface ocean variable."

- How is the temperature in the shallow, medium and deep ocean changing?
- How is the ocean heat content changing?

$$H = \rho c_p \int_{h2}^{h1} T(z) dz$$

-H: Ocean heat content

-ρ: water density

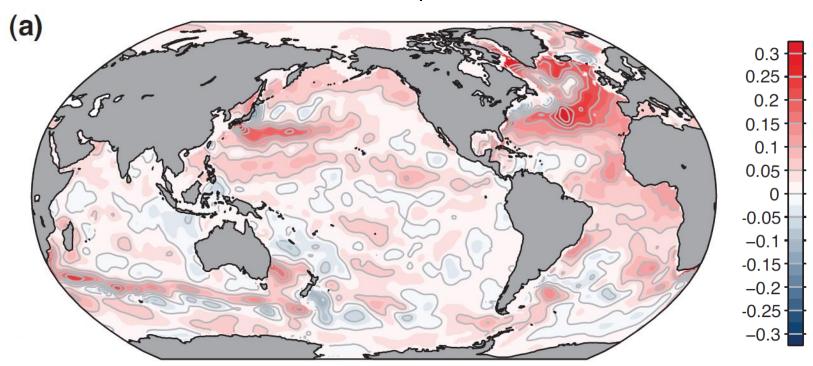
-c_p: specific heat capacity for sea water

-h_{1,2}: ocean depth

-T: Temperature

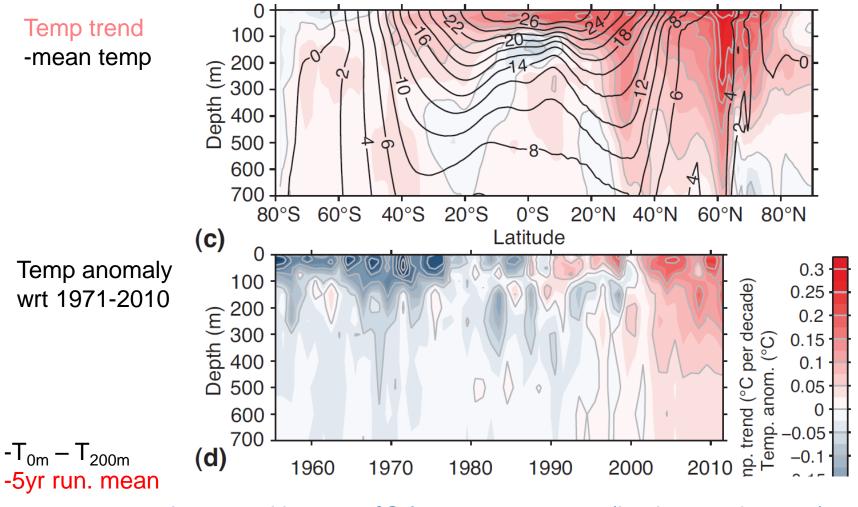
Temperature trend 1971-2010 (deg C/decade)

0-700 m depth



- Positive temperature change over most of the globe (Levitus et al., 2009).
- Warming is more prominent in the NH, especially the North Atlantic.
- This result holds in different analyses, using different time periods, bias corrections and data sources.

Temperature trend - Global



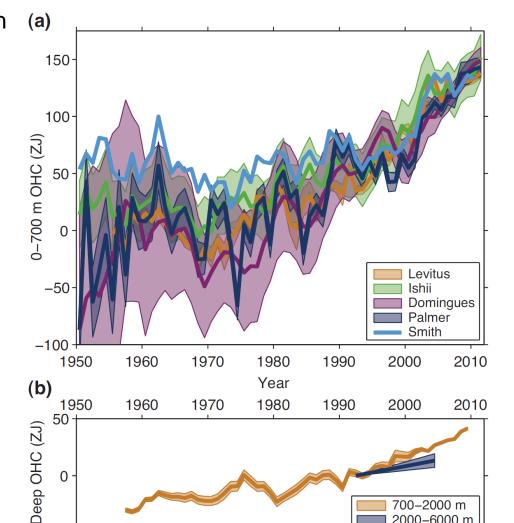
- Increased by ~0.25°C from 1971 to 2010 (Levitus et al., 2009);
- corresponds to a 4% increase in density stratification;
- is widespread in all oceans north of 40°S.

Why is the Northern Ocean warming stronger than the Southern Ocean?

Discuss together

Ocean heat content (OHC)

0-700 m depth $(1\sigma \text{ std dev.})$



700-2000 m 2000-6000 m

> Slowing of the upper OHC between 2003 and 2010(?)

700-2000 m

2000-6000 m

Global integrals of 0 to 700 m upper OHC estimated from Ocean temp. measurements show a gain from 1971-2010.

Increasingly uncertain for earlier years, especially prior to 1970.

Decreases few years following major volcanic eruptions (1963, 1982, 1991).

-50 -

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GEF4400/9400 changed time schedule

Changed GEF4400/9400 time schedule during September and November 2015:

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Mo. 14.09.15, 12:15-14:00, Wed -
Mo. 21.09.15: 10:00-12:30, Wed -
Mo. 28.09.15: 10:00-12:30, Wed -
Mo. 02.11.15: 10:00-12:30, Wed 04.11.15 10:15-12:00
Mo. 09.11.15: 10:00-12:30, Wed 11.11.15 10:15-12:00
Mo. 16.11.15: 10:00-12:30, Wed 18.11.15 10:15-12:00
Mo. 23.11.15: 10:00-12:30, Wed 25.11.15 10:15-12:00
```

GEF4400/ GEF9400 topic presentations

Christine Smith-Johnsen

- 5.1: Polar amplification in the past (KK)
- 8.1: Has solar activity contributed significantly to global warming? (JEK)

Malte Ziemek

5.2: How do volcanic eruptions affect climate? (KK)

Susanne Foldvik (JEK)

4.2: Stability of the Antarctic ice sheet (JEK)

Charalampos Sarachidis

3.2: El Niño in the past, present and future (KK)

Hans Brenna

- 6.2: Causes and relevance of oxygen minimum zones (KK)
- 8.2: CO2 doubling (JEK)

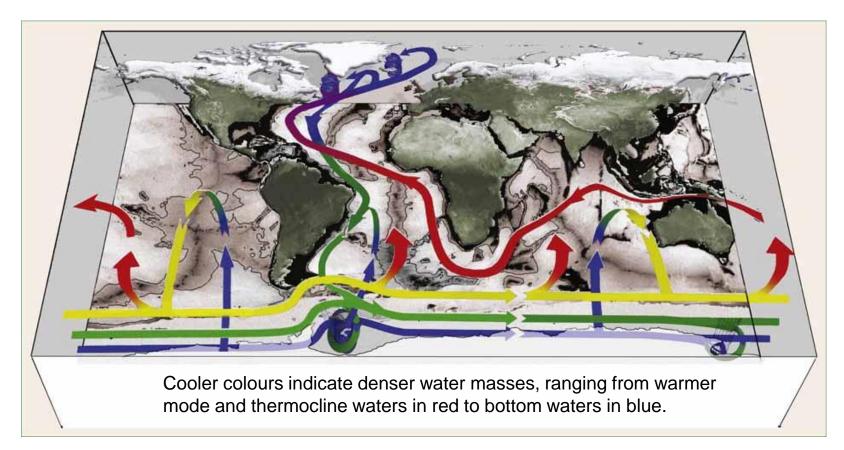
Sara Marie Blichner

- 7.1) What is the sign of the Cloud Feedback? (JEK)
 or
 - 7.2) Basic aspects of Aerosol-Cloud Interactions (JEK)

Suggested presentation time plan

	Charalampos:	
	El Nino?	
Chapter 5: Paleo climate archives (KK)	28.09.2015	-
	(JEK)	Sara: tba
Chapter 7: Clouds and Aerosols	05.10.2015	07.10.2015
		Christine: Solar
	(JEK)	activity
Chapter 8: Radiative Forcing	12.10.2015	14.10.2015
Andøya	JEK 19.10.2015	JEK 21.10.2015
	Susanne: AA ice	Hans: CO2
	sheet	doubling
Chapter 8: Radiative Forcing	26.10.2015	28.10.2015
	Charalampos:	
Chapter 5 Paleo climate archives and	El Nino	KK
Ch. 6: Biogeochemistry	02.11.2015	05.11.2015
	Hans: Oxygen	
	minimum zones	KK
Chapter 6: Biogeochemistry	09.11.2015	11.11.2015
	Christine: Polar	
	Amplification	KK
Chapter 12: Long-term Climate Change	16.11.2015	18.11.2015
	Malte: Volcano-	
	Climate	KK
Chapter 12: Long-term Climate Change	23.11.2015	25.11.2015

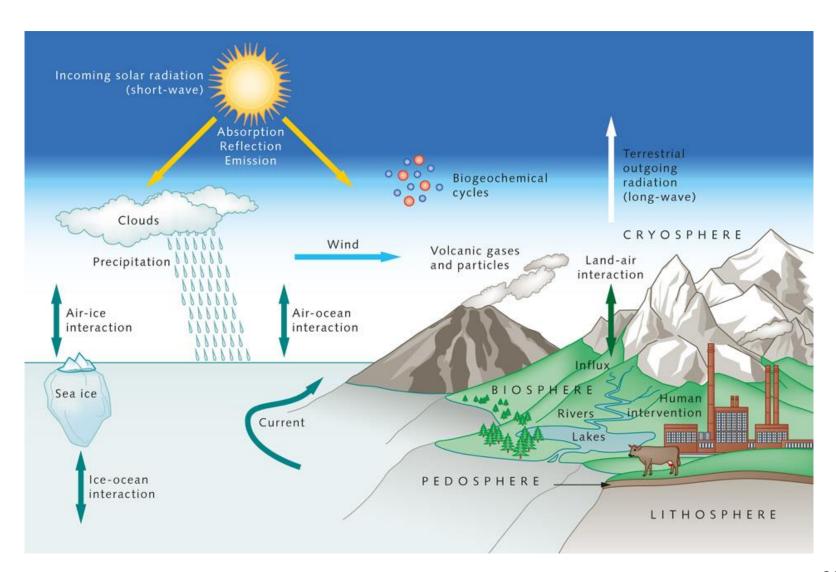
Ocean circulation



Marshall and Speer (2012, Nature)

Meridional Overturning Circulation (MOC) schematic driven mainly by the difference in **heat**, **salinity**, **wind** and **eddies**. In the early schematic of the conveyer belt analogy by Broecker (1991) the role for the Southern Ocean was neglected.

Ocean-atmosphere interactions



Ocean heat content (OHC)

0-700 m depth (a) $(1\sigma \text{ std dev.})$ 150 100 0-700 m OHC (ZJ) evitus Ishii **Domingues** Palmer Smith -100 -700-2000 m 1950 1960 1970 1980 1990 2000 2010 Year 2000-6000 m (b) 1950 1960 1970 1980 2000 2010 1990 50 Deep OHC (ZJ) 0 -

Global integrals of 0 to 700 m upper OHC estimated from Ocean temp. measurements show a gain from 1971-2010.

Increasingly uncertain for earlier years, especially prior to 1970.

Decreases for a few years following major volcanic eruptions (1963, 1982, 1991).

Slowing of the upper OHC between 2003 and 2010(?) 35

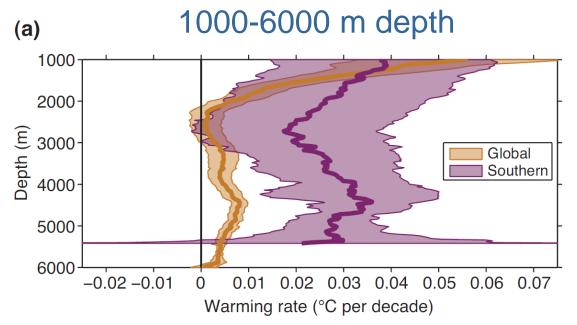
700-2000 m

2000-6000 m

1 Zepta Joule (ZJ): 10²¹ Joules (=kg·m²/s²)

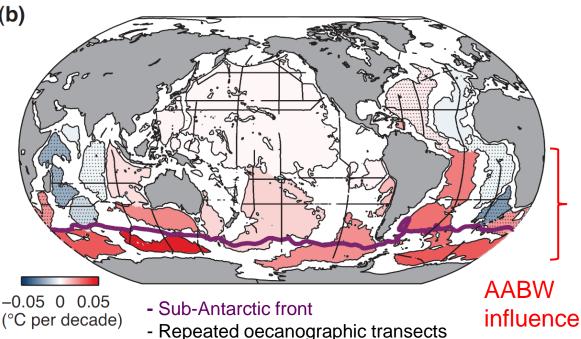
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Warming rates for 1992-2005 (deg C/decade)



<4000 m depth (b)

- No significant temperature trend between 2000-3000 m depth;
- <3000 m depth warming >0 is likely esp. in recently formed Antarctic Bottom Water (AABW);
- highest warming rates near 4500 m, usually near sea floor, where AABW influence is strongest.

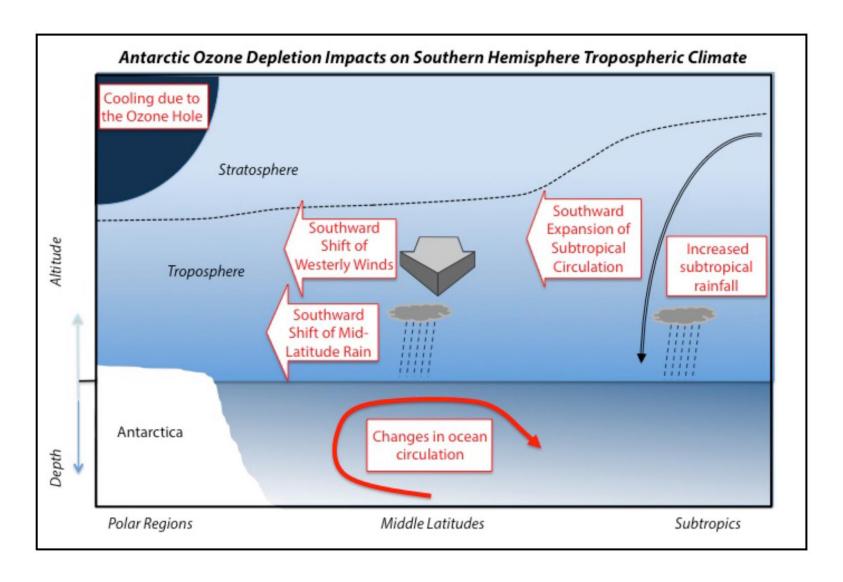


Why is the AABW warming?

Discuss together...

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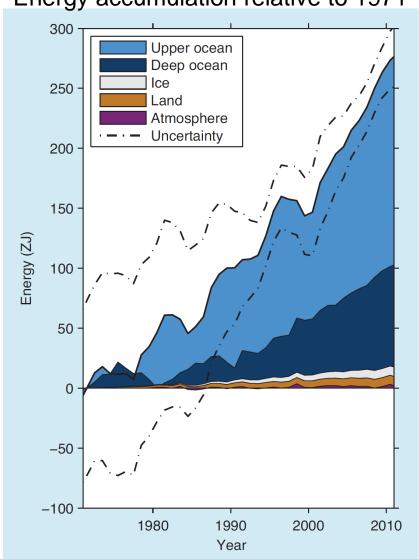


Stratospheric O₃ depletion > strengthening westerly winds (positive Southern Annular Mode) > increased surface wind stress > strengthening overturning circulation of the Southern Ocean

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Box 3.1 - Change in Global Energy Inventory

Energy accumulation relative to 1971



Estimated from satellite data since 1970:

Ocean warming dominates the total energy change inventory, accounting for ~93% from 1971 to 2010 (high confidence).

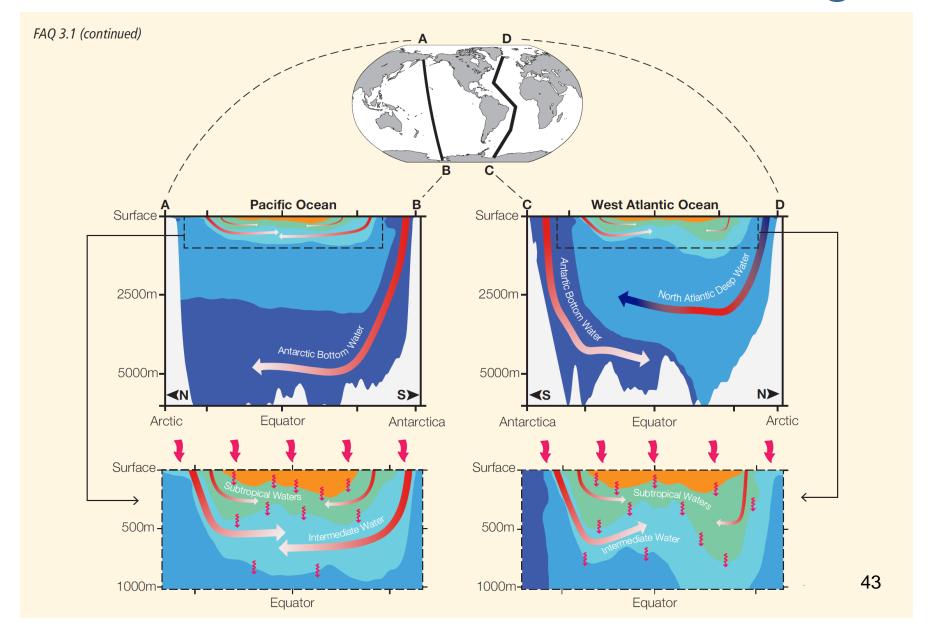
- The upper ocean (0-700 m) accounts for about 64% of the total energy change inventory.
- The deep ocean (below 700 m depth).
- Melting ice (including Arctic sea ice, ice sheets and glaciers) accounts for 3% of the total.
- Warming of the continents 3%.
- Warming of the atmosphere makes up the remaining 1%.

FAQ 3.1: Is the ocean warming?

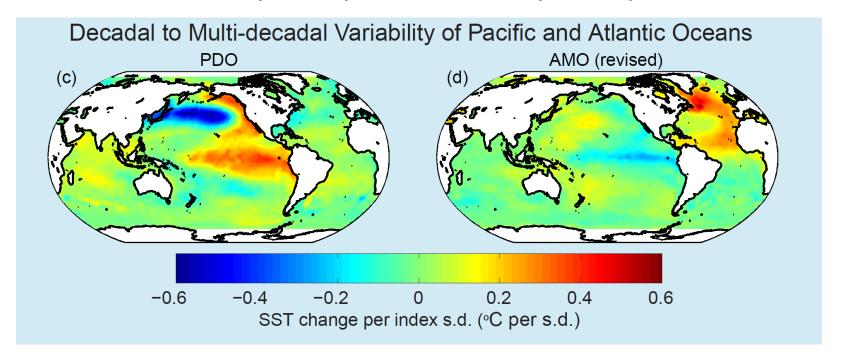
First, let's discuss in groups pro's and con's.

- "Yes, the ocean is warming over many regions, depth ranges and time periods,
- although neither everywhere nor constantly.
- The signature of warming emerges most clearly when considering global, or even ocean basin, averages over time spans of a decade or more.
- Ocean temperature at any given location can vary greatly with the seasons.
- It can also fluctuate substantially from year to year or even decade to decade - because of variations in ocean currents and the exchange of heat between ocean and atmosphere."

FAQ 3.1: Is the ocean warming?

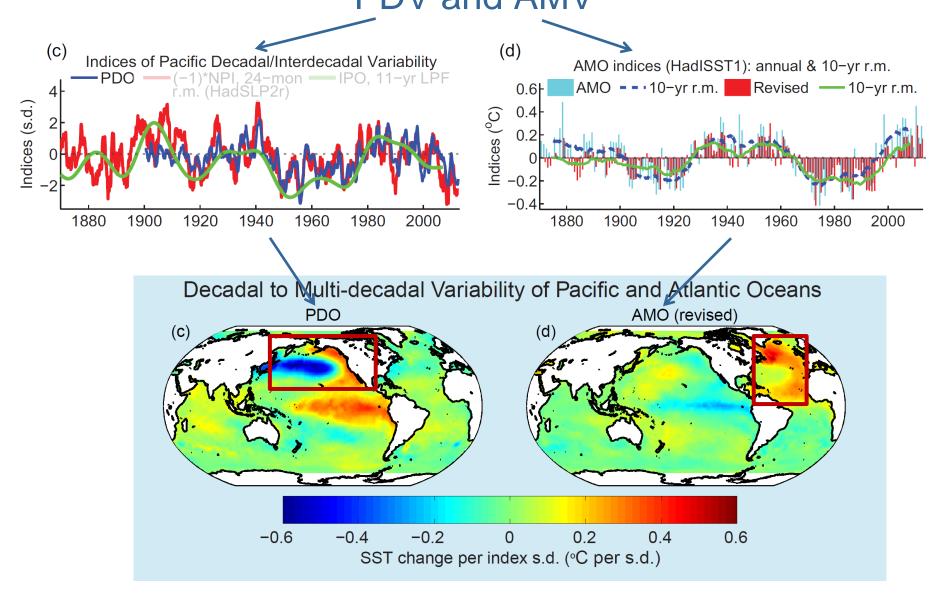


Box 2.5 - Patterns and Indices of Climate Variability – PDV (PDO) and AMV (AMO)



	Pacific Decadal Oscillation	1st PC of monthly N. Pacific SST _a field [20°N–70°N] with subtracted global
Pacific Decadal and Interdecadal Variability	(PDO)	mean
	Inter-decadal Pacific Oscillation (IPO)	Projection of a global SST _a onto the IPO pattern, which is found as one of the leading Empirical Orthogonal Functions of a low-pass filtered global SST _a field
Atlantic Ocean Multidecad	Atlantic Multi-decadal Oscillation (AMO) index	10-year running mean of linearly detrended Atlantic mean SST _a [0°–70°N]
Variability	Revised AMO index	As above, but detrended by subtracting SST _a [60°S–60°N] mean

Box 2.5 - Patterns and Indices of Climate Variability - PDV and AMV



3.2 Conclusions - Temperature and Heat Content Changes

- "It is virtually certain that the upper ocean (above 700 m) has warmed from 1971 to 2010, and likely that it has warmed from the 1870s to 1971. Confidence in the assessment for the time period since 1971 is high.
- It is **likely** that the **ocean warmed between 700 and 2000 m from 1957 to 2009**, based on 5-year averages. It is likely that the **ocean warmed from 3000 m to the bottom from 1992 to 2005**, while no significant trends in global average temperature were observed between 2000 and 3000 m depth during this period.
- It is virtually certain that upper ocean (0 to 700 m) heat content increased during the relatively well-sampled 40-year period from 1971 to 2010.
- Warming of the ocean between 700 and 2000 m likely contributed about 30% of the total increase in global ocean heat content (0 to 2000 m) between 1957 and 2009.
- Ocean warming dominates the global energy change inventory.
- Warming of the ocean accounts for about 93% of the increase in the Earth's energy inventory between 1971 and 2010 (high confidence), with warming of the upper (0 to 700 m) ocean accounting for about 64% of the total."