

# GEF4400 “The Earth System”

Prof. Dr. Jon Egill Kristjansson,

**Prof. Dr. Kirstin Krüger (UiO)**

- **Lecture/ interactive seminar/ field excursion**

Teaching language: English

Time and location: Monday 12:15-14:00

Thursday 14:15-16:00, CIENS Glasshallen 2.

- **Study program**

Master of meteorology and oceanography

PhD course for meteorology and oceanography students

- **Credits and conditions:**

The successful completion of the course includes an **oral presentation (weight 50%)**, a **successful completion of the Andøya field excursion (mandatory)**, a **field report**, as well as a final **oral examination (50%)**. Student presentations will be part of the course.



# GEF4400/9400 **changed** time schedule

Changed GEF4400/9400 time schedule during November 2015:

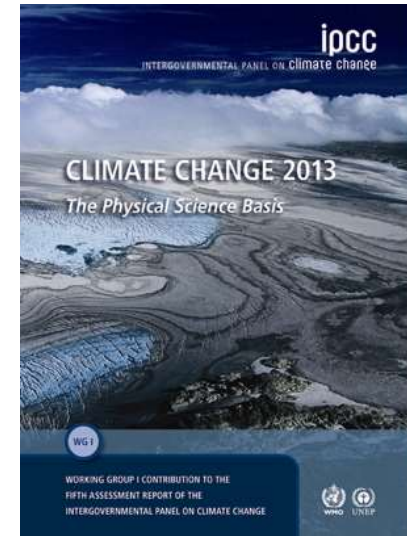
Mo. 02.11.15: **10:00-12:30**, Wed 04.11.15 10:15-12:00

Mo. 09.11.15: **10:00-12:30**, Wed 11.11.15 10:15-12:00

Mo. 16.11.15: **10:00-12:30**, Wed 18.11.15 10:15-12:00

Mo. 23.11.15: **10:00-12:30**, Wed 25.11.15 10:15-12:00

# IPCC Chapter 6: Carbon and other biogeochemical cycles



- Background
- Introduction: Global Carbon Cycle (*Section 6.1*)
- Evolution of biogeochemical cycles since industrial era (*Section 6.3*)
- Variations in Carbon cycle before the fossil fuel era (*Section 6.2*)
- Projections of future carbon cycles (*Section 6.4*)
- Global Carbon Budget in 2014
- Executive Summary (Ch. 6)

Ciais, P., et al., 2013: Carbon and Other Biogeochemical Cycles. In: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change. Cambridge University Press.

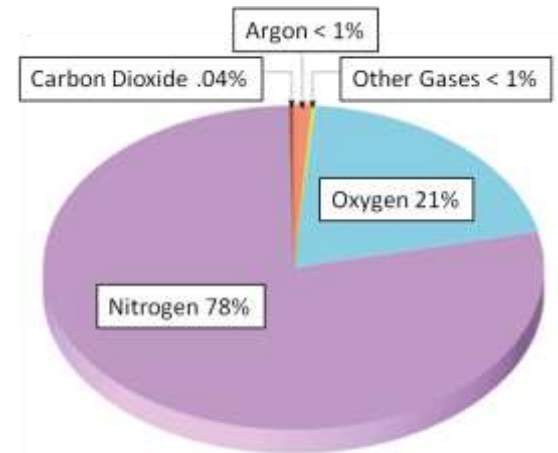


# Carbon dioxide (CO<sub>2</sub>) in the atmosphere

# Content of air

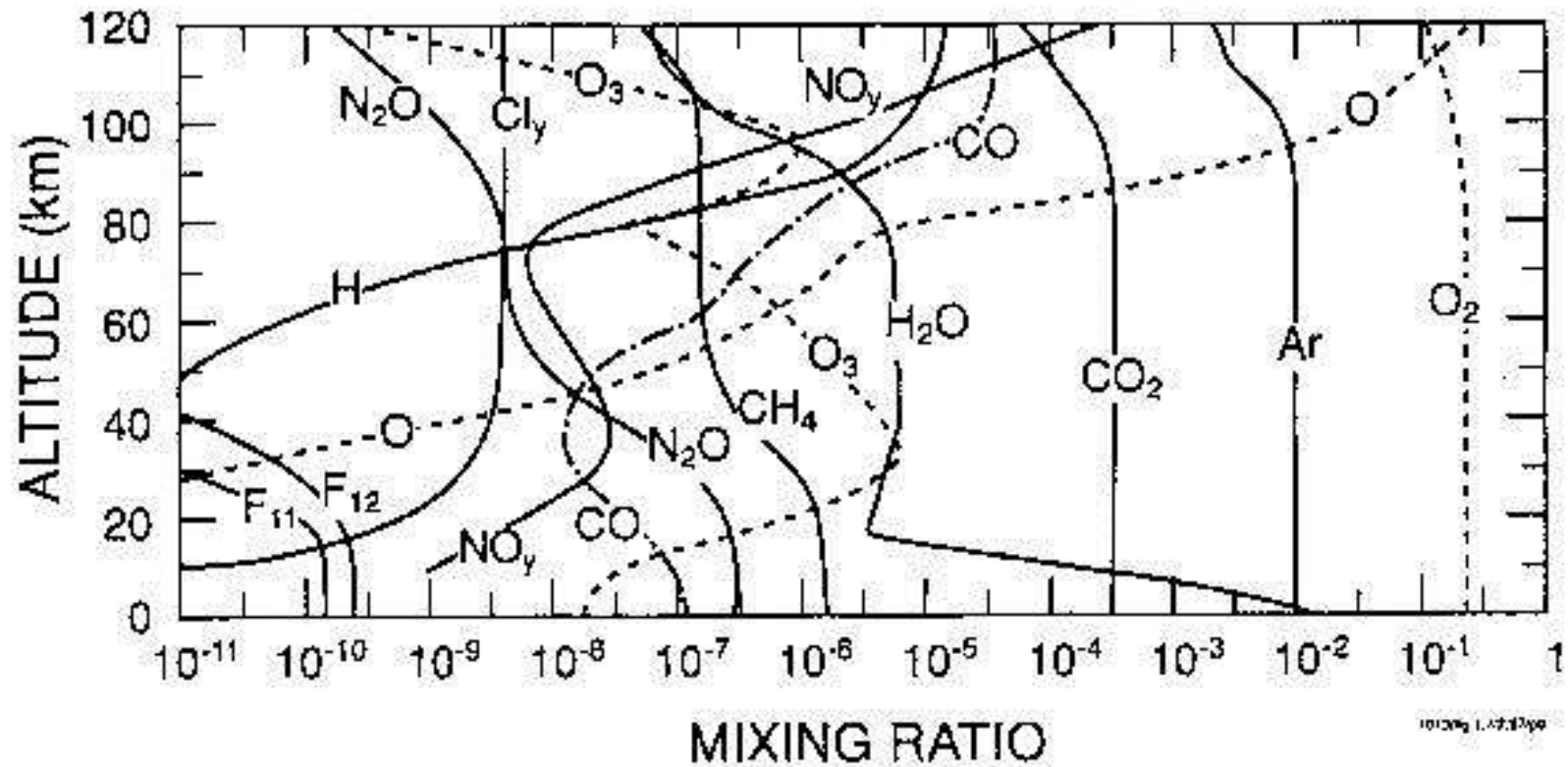
Nitrogen N <sub>2</sub>	78
Oxygen O <sub>2</sub>	21
Inert gases (Ar)	0.9
	<hr/>
	99.9

Percent of volume (%)



Carbon dioxide CO <sub>2</sub>	0.04, varying
Ozone O <sub>3</sub>	0.00005, varying
Water vapour H <sub>2</sub> O	highly varying
+ other trace gases	

## Background



Typical vertical distribution of chemical species within the air [Brasseur, 1999].

# Mixing ratios of trace gases

1 ppm (1 part per million) 1 particle CO<sub>2</sub> per 10<sup>6</sup> particles air

1 ppb ( 1 part per billion): 1 particle CO<sub>2</sub> per 10<sup>9</sup> particles air

1 ppt ( 1 part per trillion): 1 particle CO<sub>2</sub> per 10<sup>12</sup> particles air

“v”: per volume

“m”: per mass

Mixing ratio is a relative unit → taking the air density into account

Absolute unit: concentration of a trace gas (e.g. given in mPa, nbar,  
Dobson units for ozone)

# Carbon dioxide

Molecular formula

CO<sub>2</sub>

Molar mass

44.010 g/mol

Appearance

colorless, odorless gas

(gas at 1 atm and 0 °C; 1 atm =1013.25 hPa)

(solid at 1 atm and -78.5 °C)

(liquid at 56 atm and 20 °C)

Dipole moment

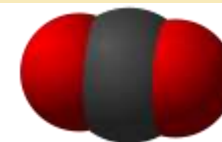
zero

Molecular shape

linear

Spectral data

UV, IR



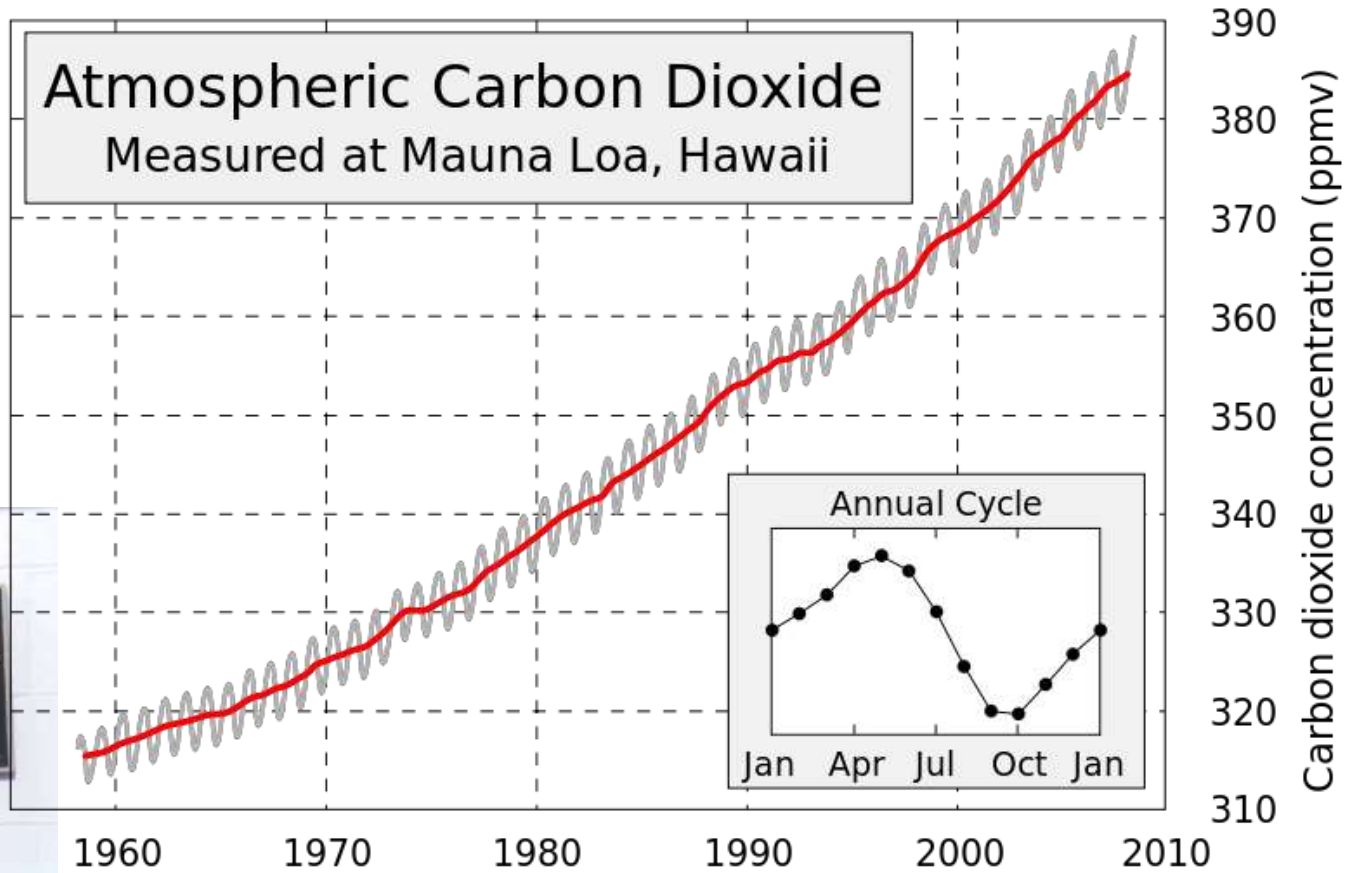
**CO<sub>2</sub>:** trace gas, **0.0398%** concentration of the atmosphere, **sources and sinks** are at the surface, uniform distribution up to 90 km.

**Sources:** combustion of fossil fuels, burning of vegetable matter, chemical processes, respiration, volcanoes , geothermal processes, dissolution of carbonates in crustal rocks.

**Sinks:** ocean, sediments, biosphere(photosynthesis)

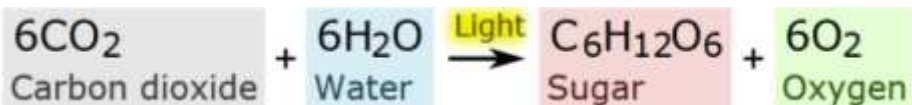
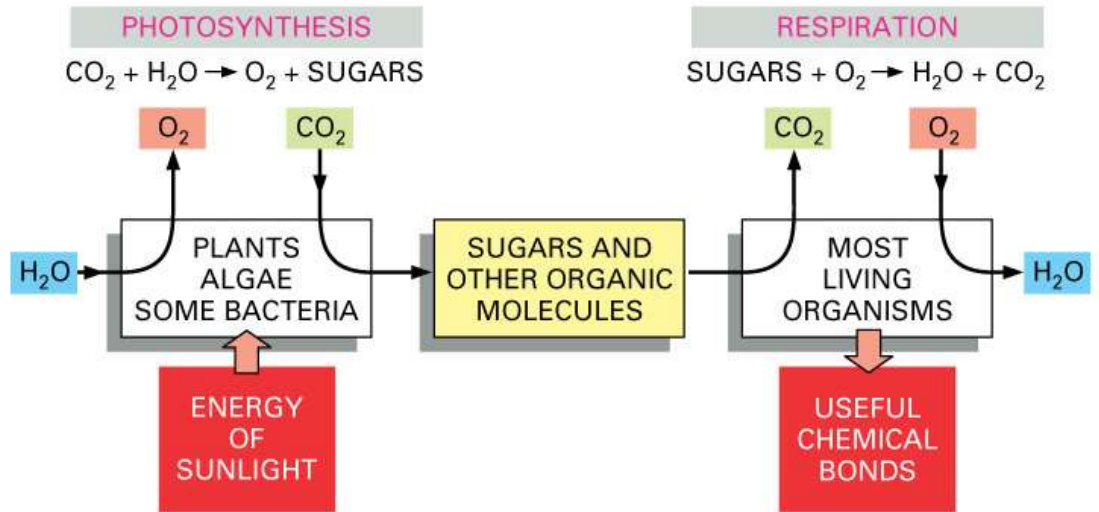
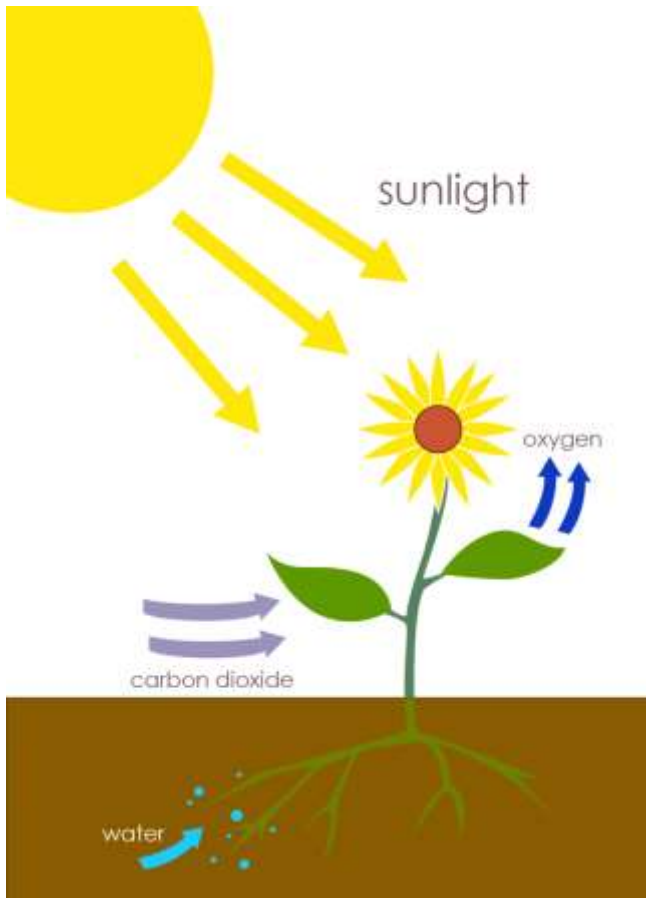


# Mauna loa curve (Keeling curve)

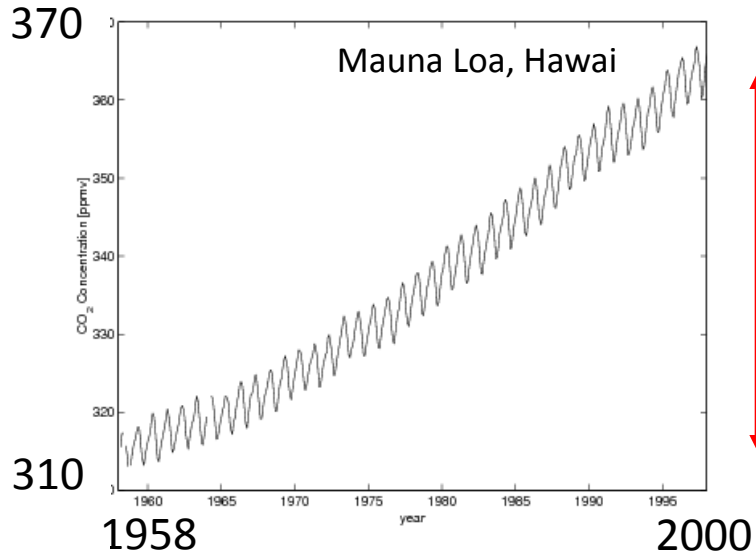


Charles D. Keeling  
1928-2005

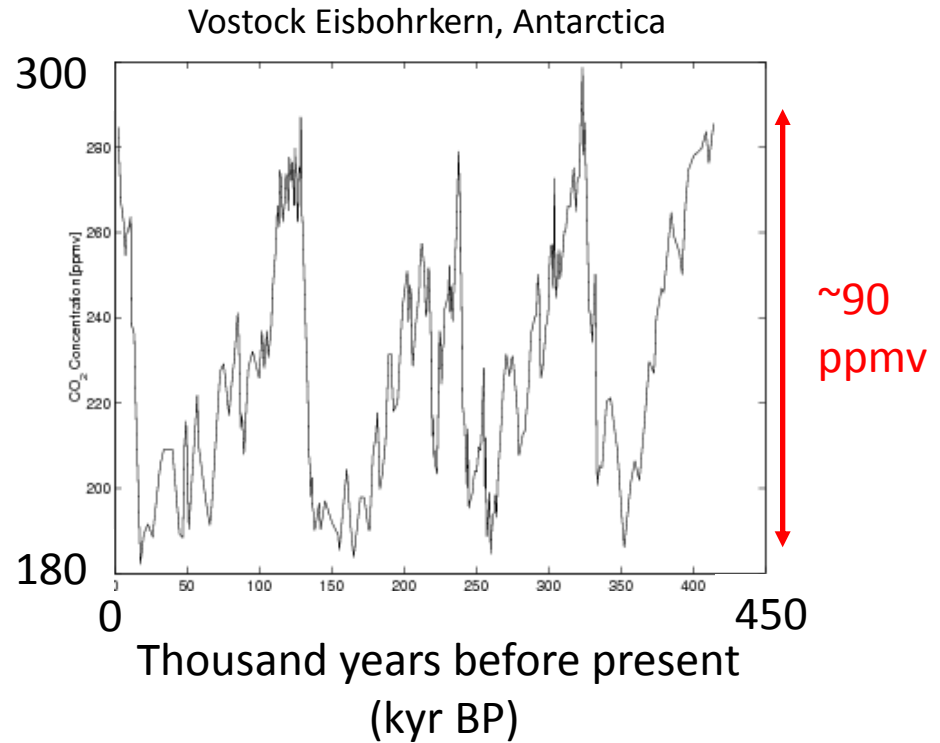
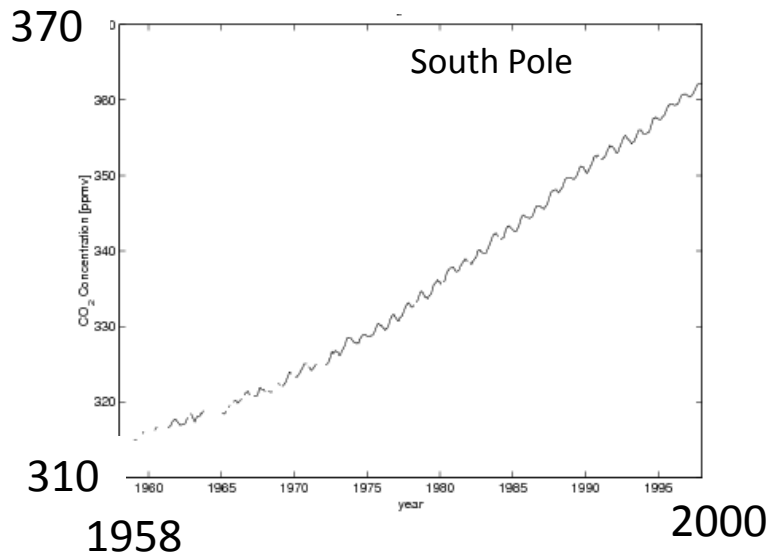
# Photosynthesis + Respiration



# Past and present CO<sub>2</sub> mixing ratio (ppmv)



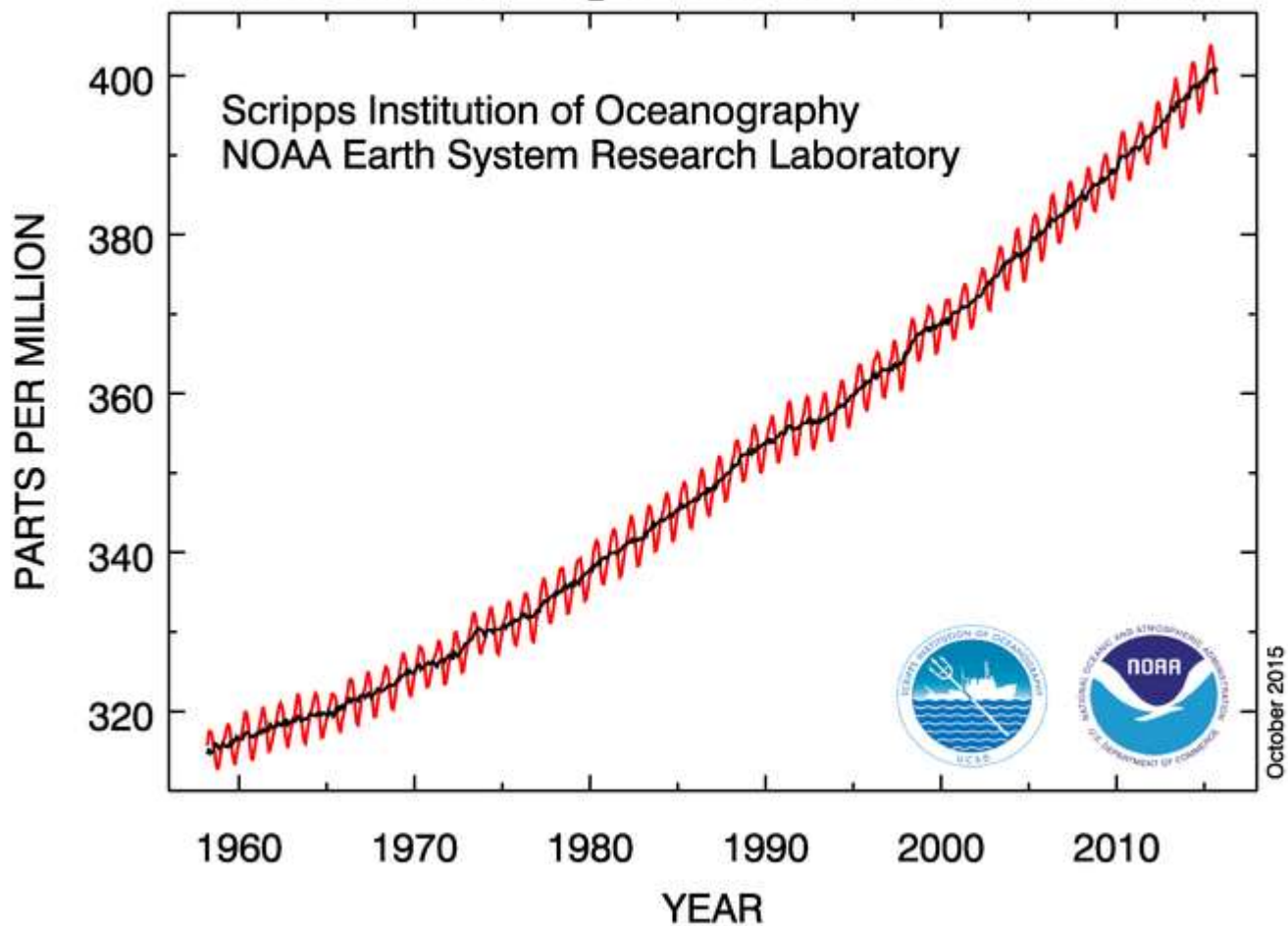
~55  
ppmv



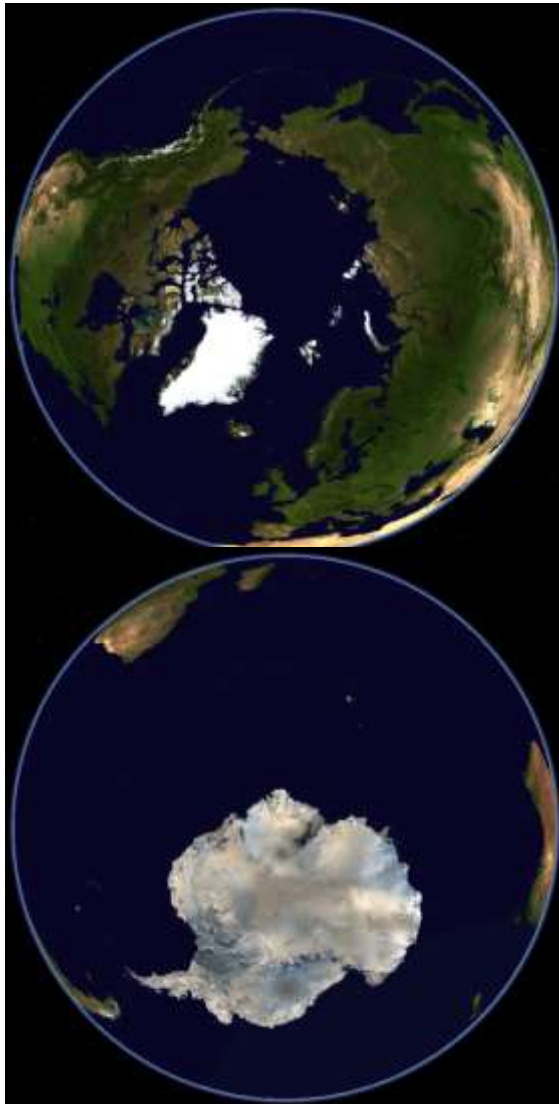
~90  
ppmv

(Macke, 2004)

## Atmospheric CO<sub>2</sub> at Mauna Loa Observatory



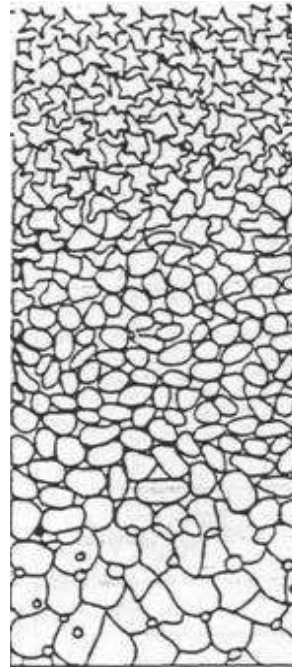
# Ice cores as climate archives



young



old



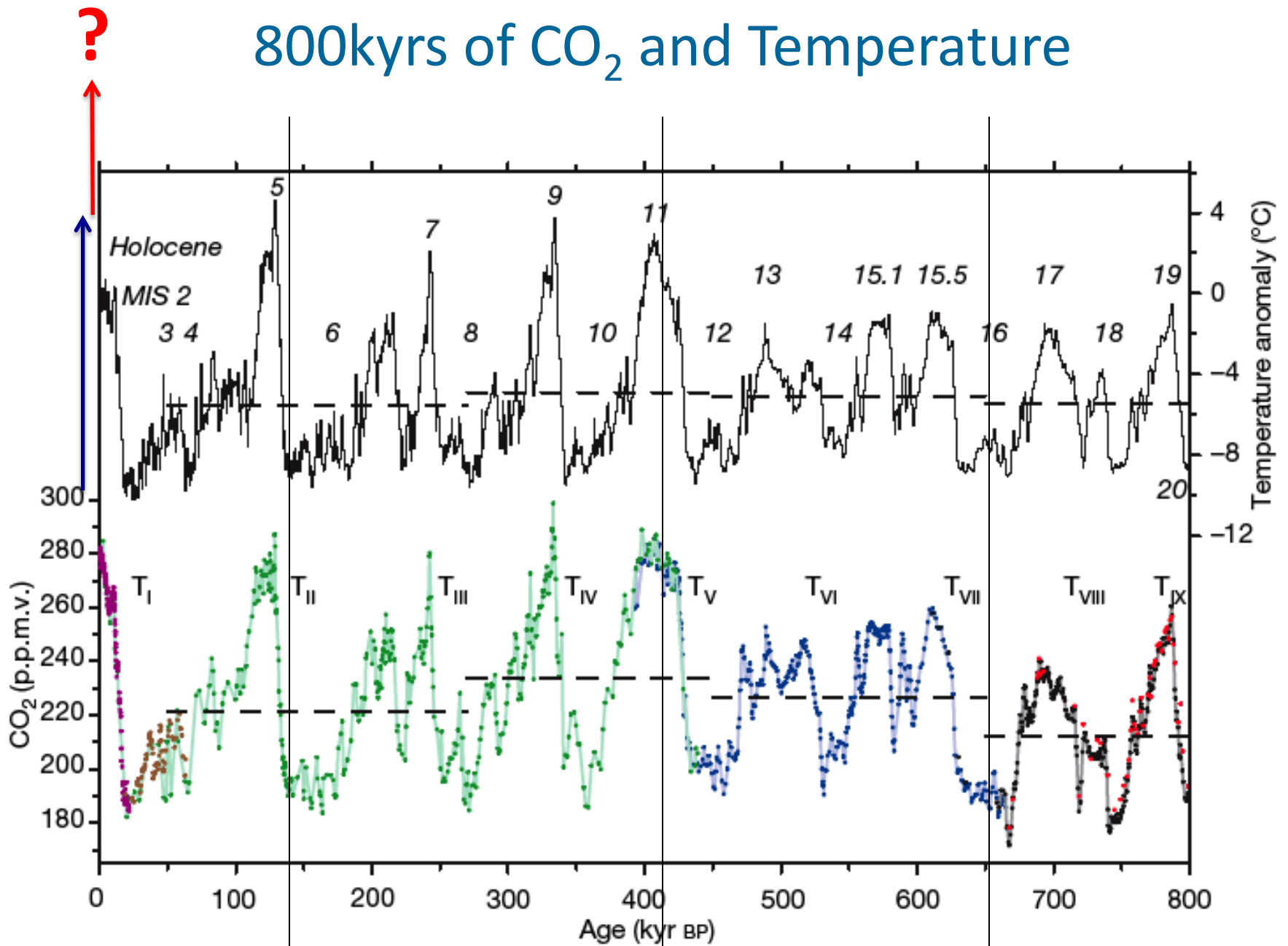
snow

firn

ice



# 800kyrs of CO<sub>2</sub> and Temperature



Barnola et al., 1987

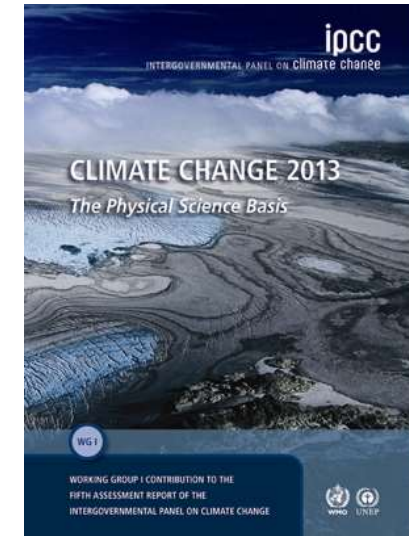
Petit et al., 1999 Siegenthaler et al., 2005 Lüthi et al., 2008





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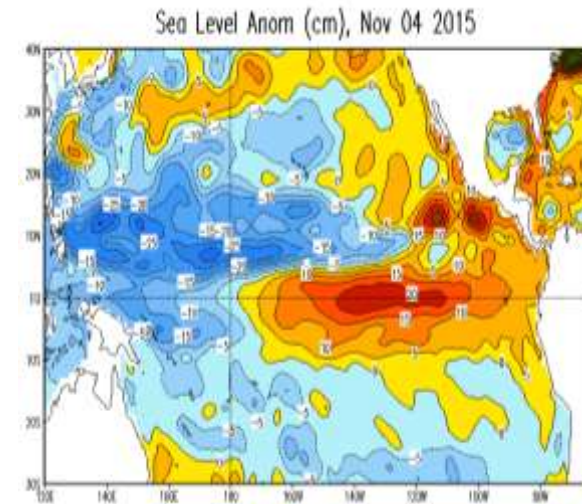
Ciais, P., et al., 2013: Carbon and Other Biogeochemical Cycles. In: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change. Cambridge University Press.



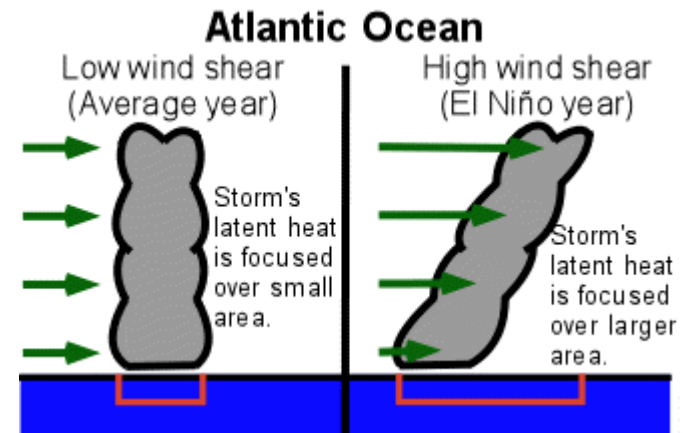


# Questions

- El Nino SLH anomaly between East and West Pacific? > 20 to 50 cm
- to respire, see-saw (in air pressure)
- Hurricane increase during El Nino?

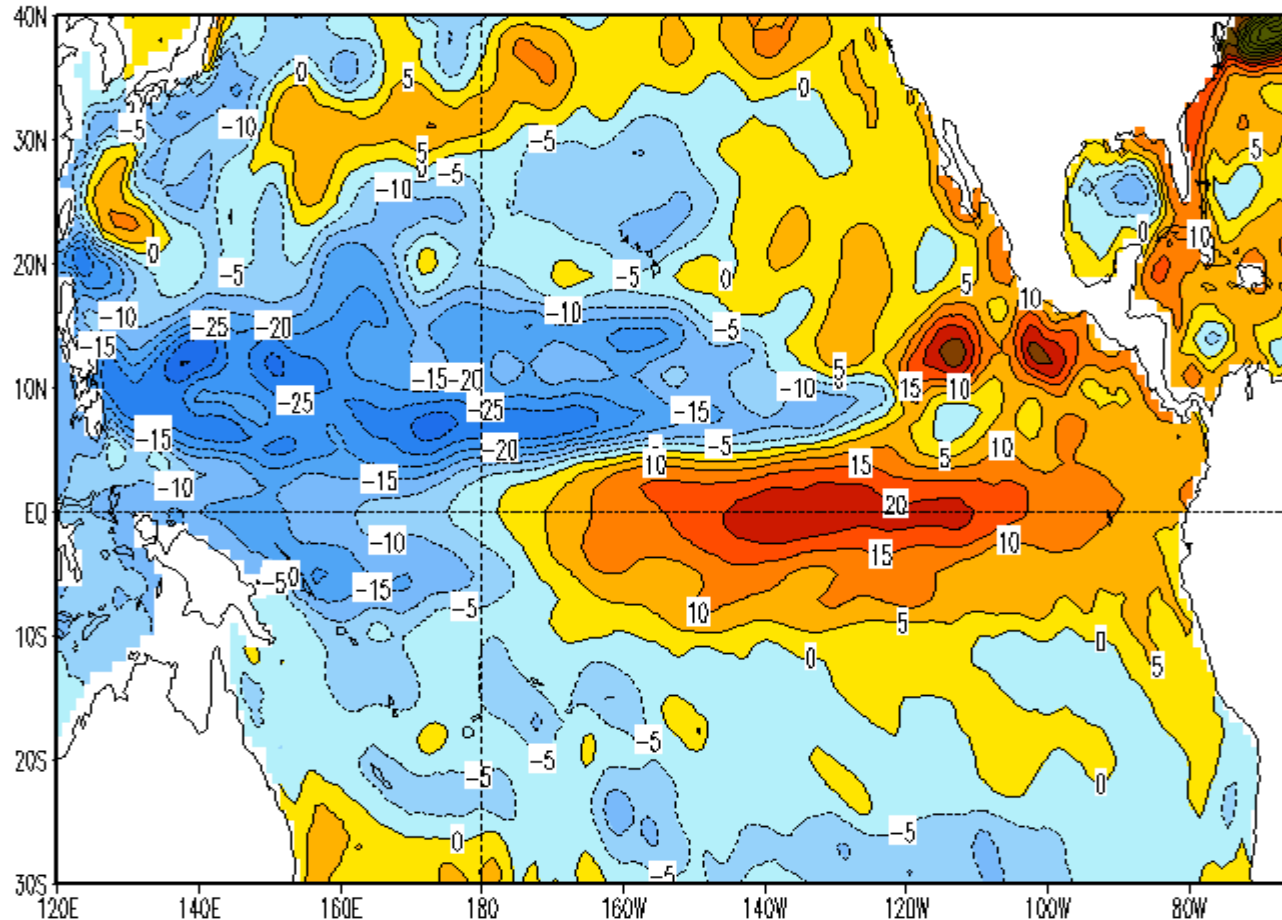


	Atlantic		Eastern Pacific	
	Average	El Niño Avg.	Average	El Niño Avg.
Named storms	9.4	7.1	16.7	17.6
Hurricanes	5.8	4.0	9.8	10.0
Intense Hurricanes	2.5	1.5	4.8	5.5



# SLH anomalies (cm) Nov, 04, 2015

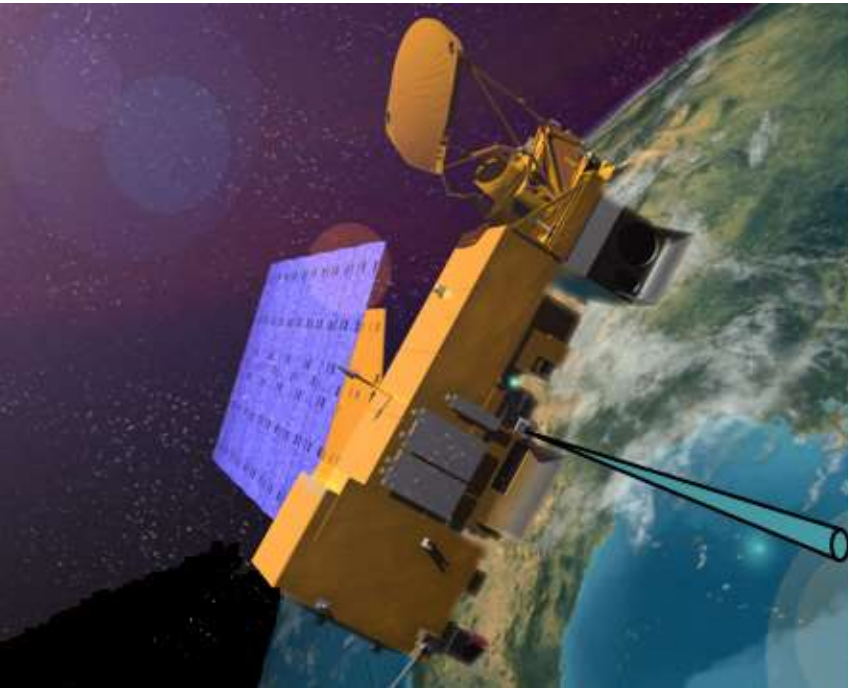
Sea Level Anom (cm), Nov 04 2015



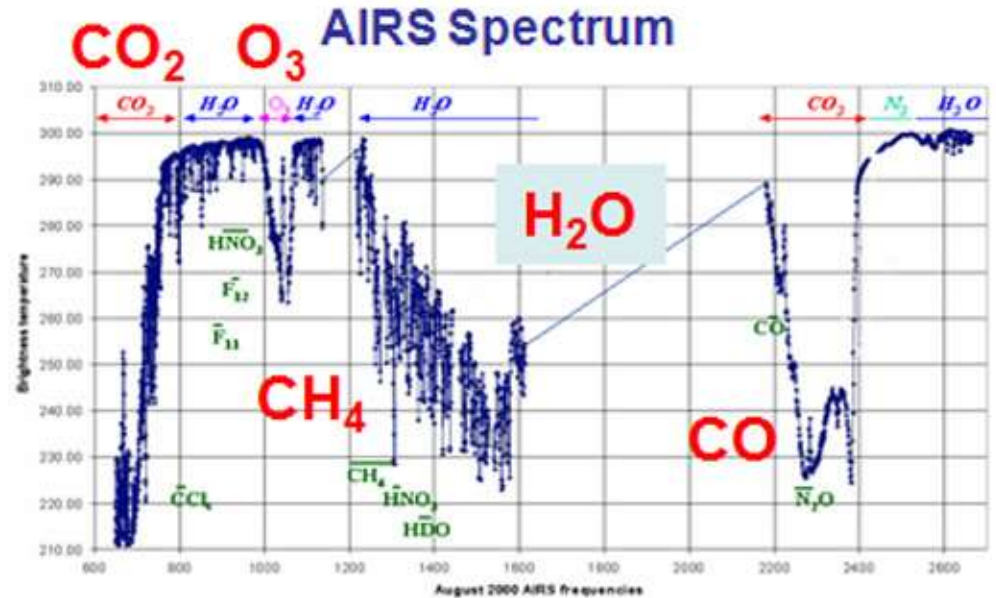
NCEP-CPC

# Carbon dioxide (CO<sub>2</sub>) in the atmosphere

# The Atmospheric Infrared Sounder (AIRS) instrument



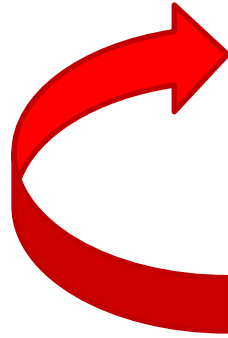
on NASA's Aqua satellite



- Launched on May 4, 2002,
- still operational (<http://airs.jpl.nasa.gov/data/near-real-time>)
- cross-track scanning instrument,
- scan mirror rotates around an axis along the line of flight and directs infrared energy from the Earth into the instrument,
- ground scan ~800 km from either side.

# CO<sub>2</sub> Increase

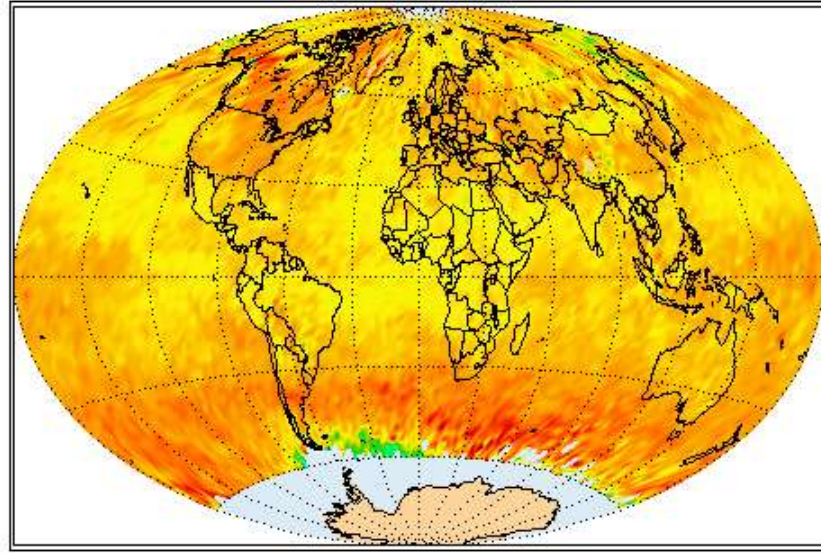
+17 ppm/7 years



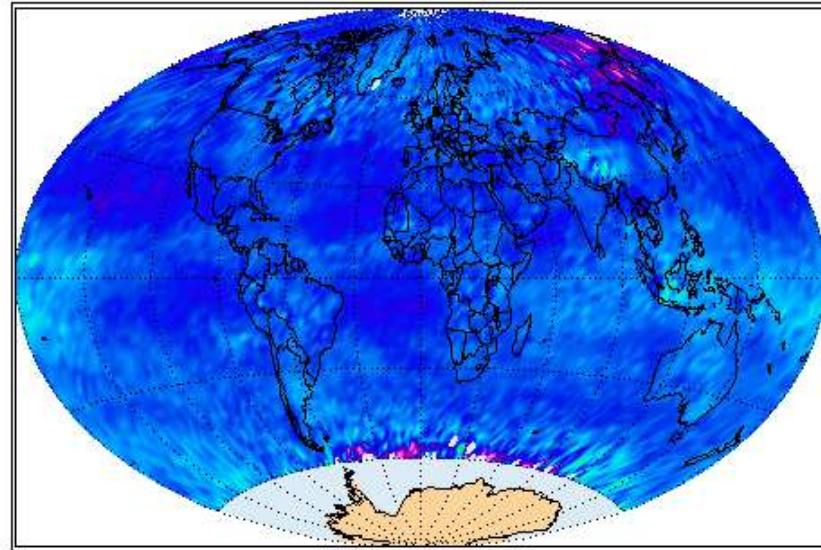
Global averaged annual mean  
Marine station data:  
-2002: 372 ppm  
-2009: 386 ppm  
-2014: 397 ppm

([www.esrl.noaa.gov/gmd/ccgg/trends/](http://www.esrl.noaa.gov/gmd/ccgg/trends/))

November, 2009



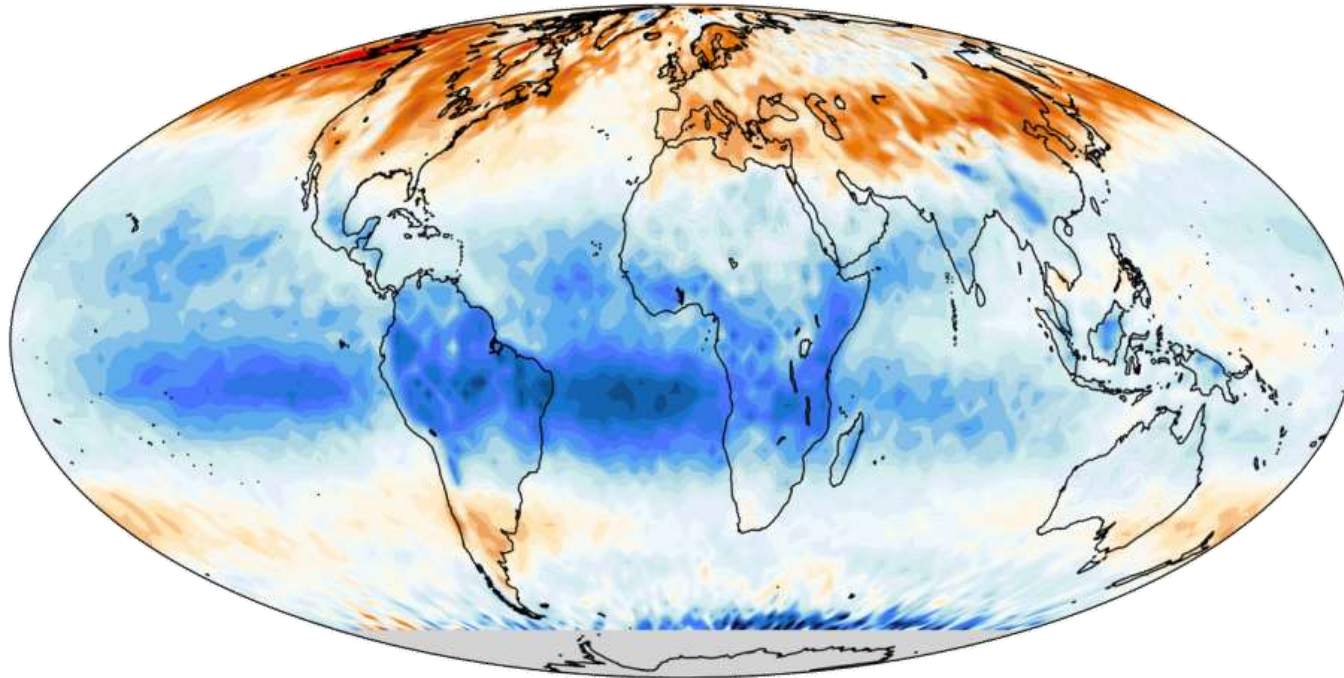
November, 2002



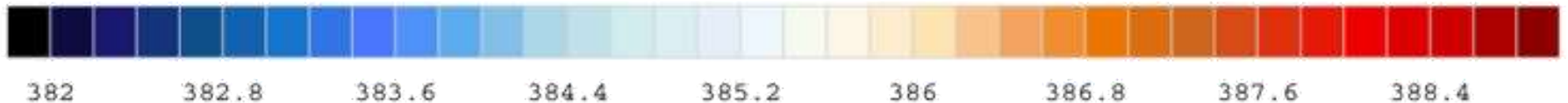


# CO<sub>2</sub> seasonal and hemispheric variations

CO<sub>2</sub> mid-troposphere (3-7 km) AIRS instrument  
Annual mean 2008



Carbon Dioxide 2008 Concentration (ppm)



<http://mirador.gsfc.nasa.gov>

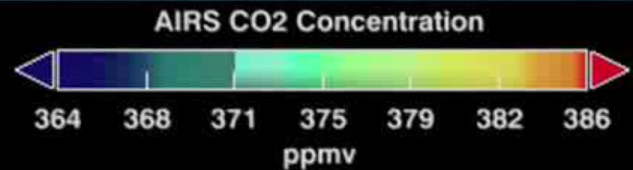
- seasonal cycle in CO<sub>2</sub> uptake by vegetation
- maximum in NH due to more emitters and higher fraction of vegetation

# CO<sub>2</sub>: Sept 2002 - July 2008

## AIRS Mid-Tropospheric Carbon Dioxide



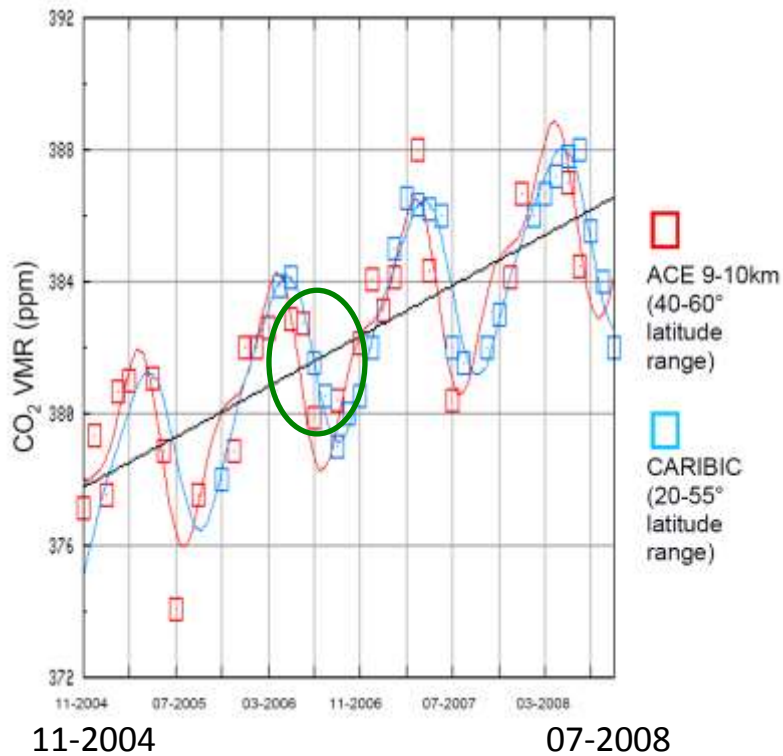
Sep 2002



<http://www.nasa.gov>



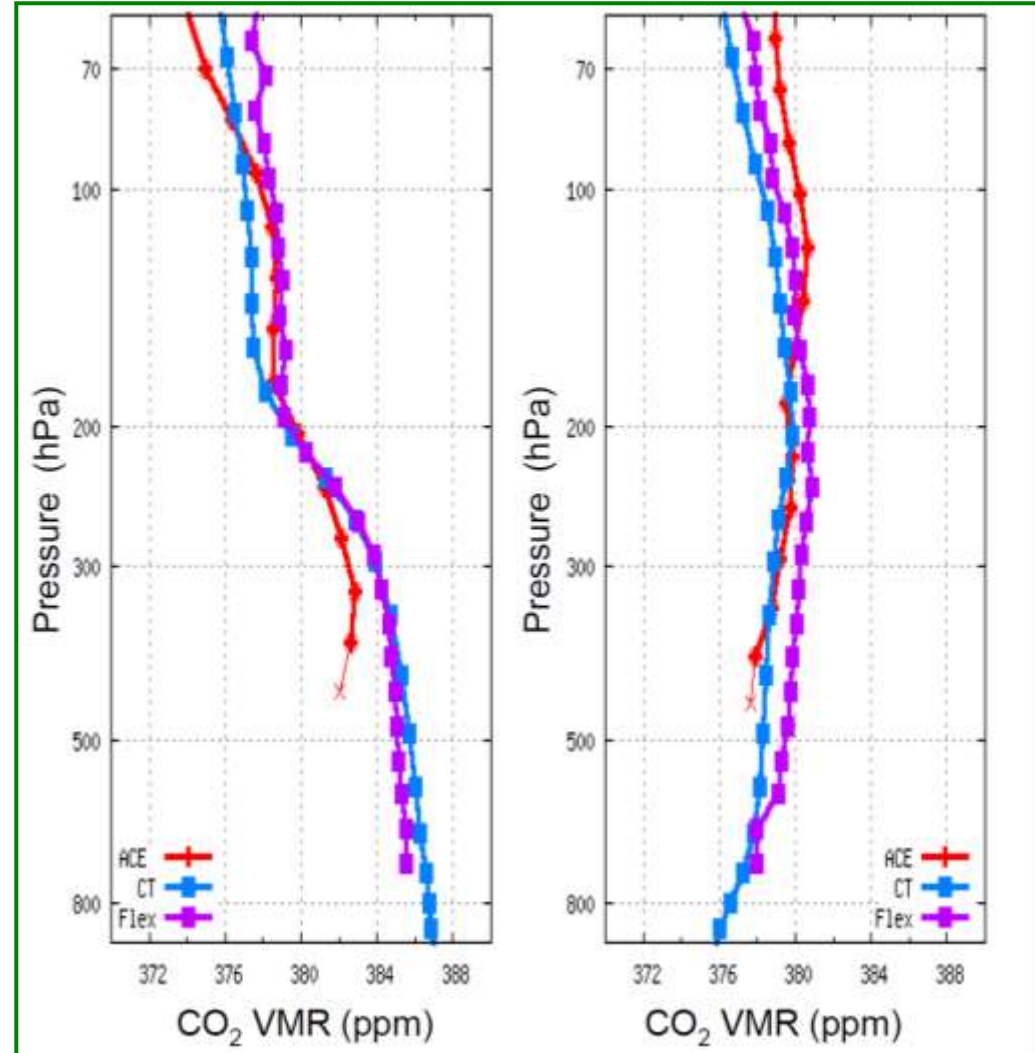
# CO<sub>2</sub> as a time distribution tracer - troposphere



- ACE: Atmospheric Chemistry Experiment (satellite)
- CARIBIC (aircraft)
- CT: Carbon Tracker (model)
- FLEXPART (model)

May 2006

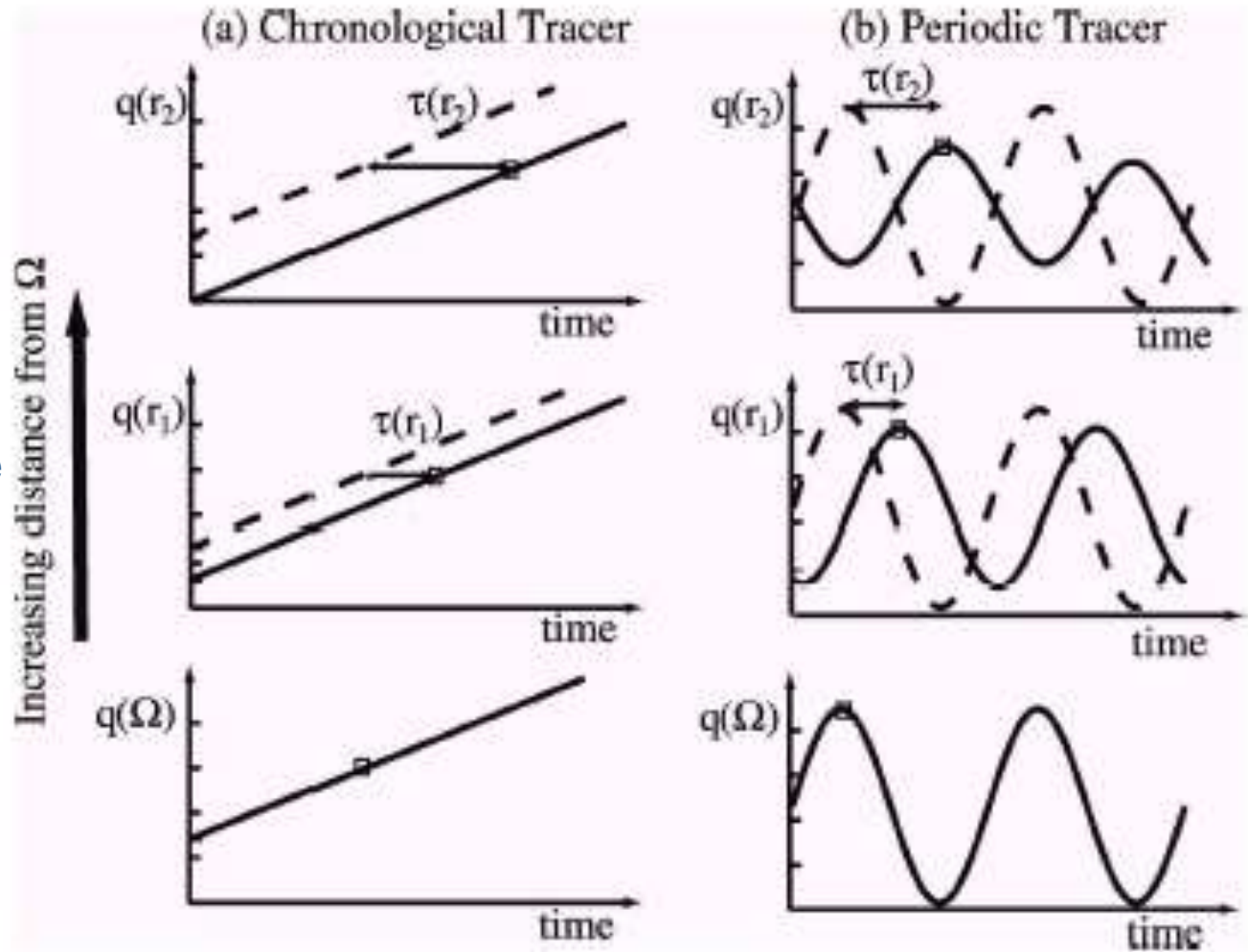
July 2006





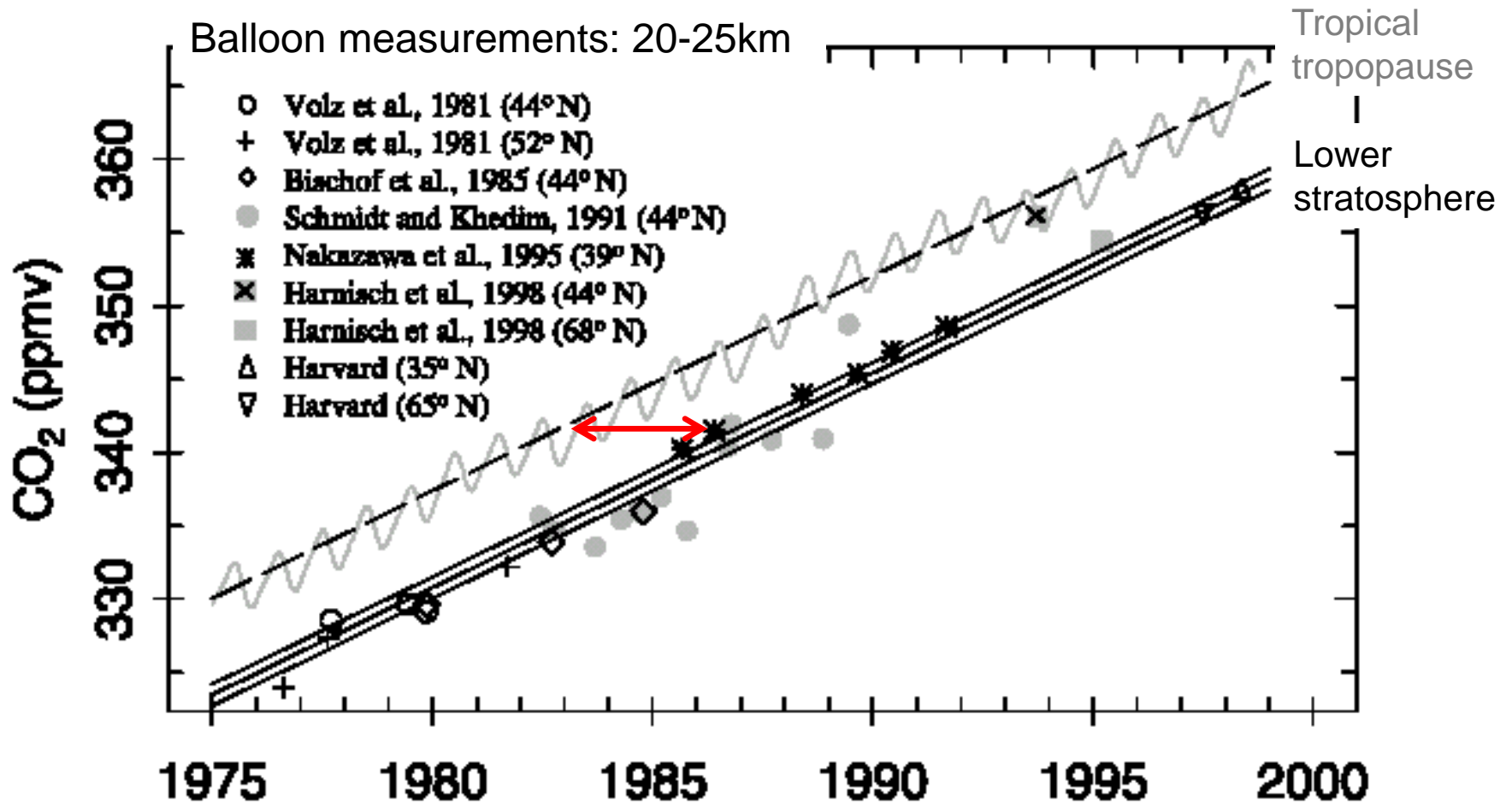
# Tracer-distribution in the stratosphere

↑  
Increasing distance  
of the tropical  
tropopause



[Waugh et al., GR 2002]

# CO<sub>2</sub> as a time distribution tracer - stratosphere



The transport from the tropical tropopause towards mid-latitudes, at 20-25 km altitude takes more than 4 years.

Andrews et al. (2001)

# Carbon Cycle

<b>Gas</b>		<b>Mixing ratio in ppm</b>	<b>Residence Time</b>	<b>Increase in % per year</b>
<b>Carbon dioxide</b>	CO <sub>2</sub>	398	5 – 200 a	0,4
<b>Methane</b>	CH <sub>4</sub>	1.8	12 a	1,5
<b>Carbon monoxide</b>	CO	0.05 – 0.2	60 – 180 d	
<b>Chlorofluorocarbon</b>	CFC	10 <sup>-3</sup>	70 – 100 a	

+ many more  
Halocarbons

# Greenhouse Gas Concentrations

GAS	Pre-1750 tropospheric concentration <sup>1</sup>	Recent tropospheric concentration <sup>2</sup>	GWP* (100-yr time horizon)	Atmospheric lifetime <sup>4</sup> (years)	Increased radiative forcing <sup>5</sup> (W/m <sup>2</sup> )
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Mixing ratios in parts per million (ppm)

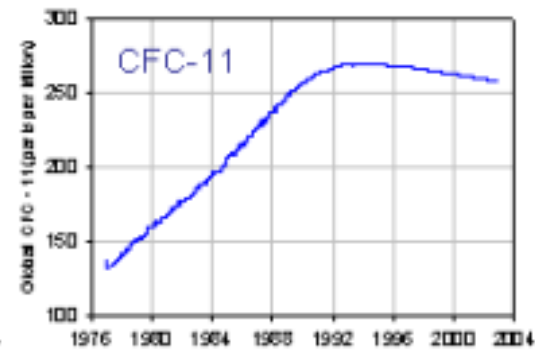
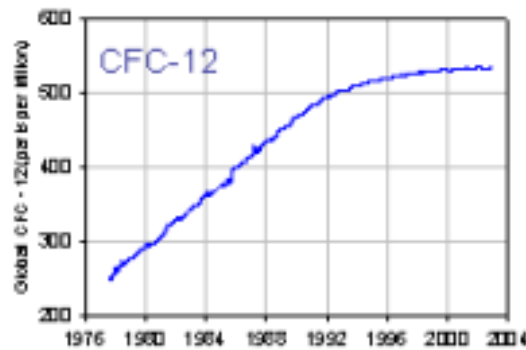
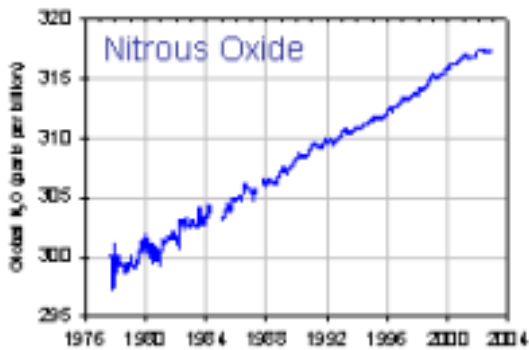
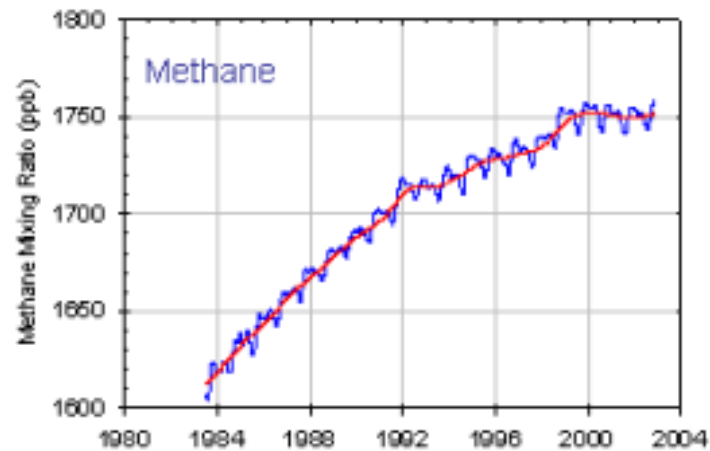
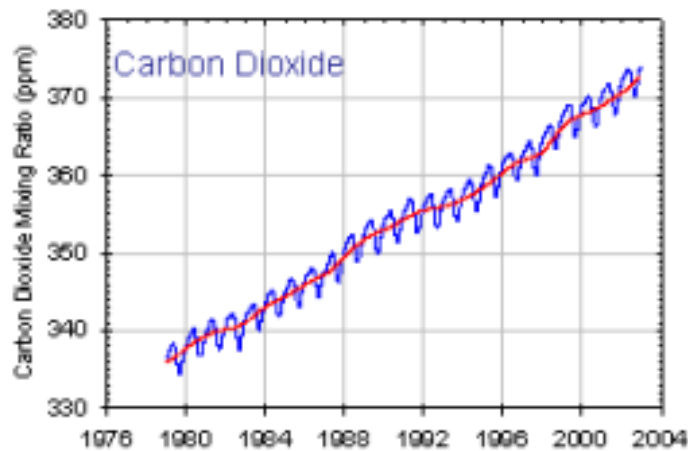
<b>Carbon dioxide (CO<sub>2</sub>)</b>	280 <sup>6</sup>	384.8 <sup>7</sup>	<b>1</b>	~ 100 <sup>4</sup>	1.66
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Mixing ratios in parts per billion (ppb)

<b>Methane (CH<sub>4</sub>)</b>	700 <sup>8</sup>	1865 <sup>9</sup> /1741 <sup>9</sup>	<b>25</b>	12 <sup>4</sup>	0.48
<b>Nitrous oxide (N<sub>2</sub>O)</b>	270 <sup>10</sup>	322 <sup>9</sup> /321 <sup>9</sup>	<b>298</b>	114 <sup>4</sup>	0.16
Tropospheric ozone (O <sub>3</sub> )	25 <sup>1</sup>	34 <sup>4,1</sup>	n.a. <sup>4</sup>	hours-days	0.35 <sup>4</sup>

\*Global Warming Potential is the ratio of the radiative forcing of a trace gas relative to that of CO<sub>2</sub>.

# Global Trends in Major Greenhouse Gases to 1/2003

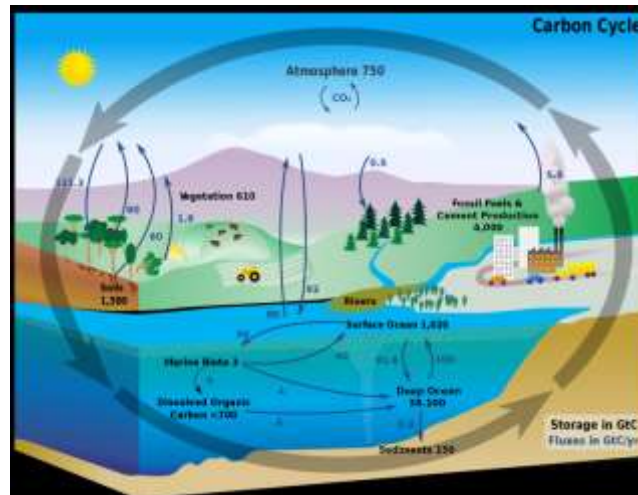


IPCC-AR  
models:  
WMGHG

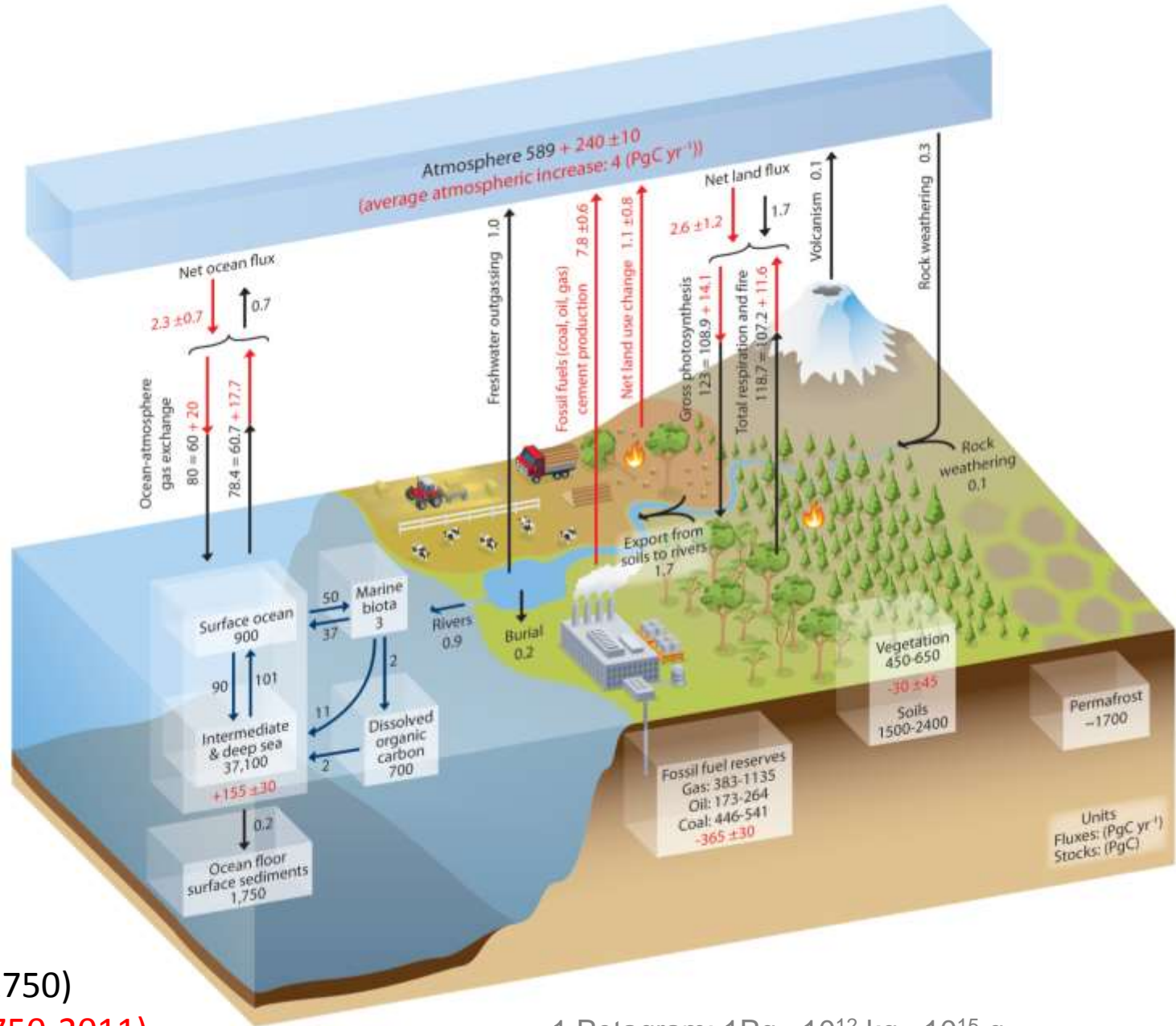
Global trends in major long-lived greenhouse gases through the year 2002. These five gases account for about 97% of the direct climate forcing by long-lived greenhouse gas increases since 1750. The remaining 3% is contributed by an assortment of 10 minor halogen gases, mainly HCFC-22, CFC-113 and  $\text{CCl}_4$ .

# The Carbon Cycle in the IPCC FAR

- Introduction: Carbon Cycle (*Section 6.1*)



# Global Carbon Cycle



→ Fluxes: PgC yr<sup>-1</sup>

█ Stocks: Pg C

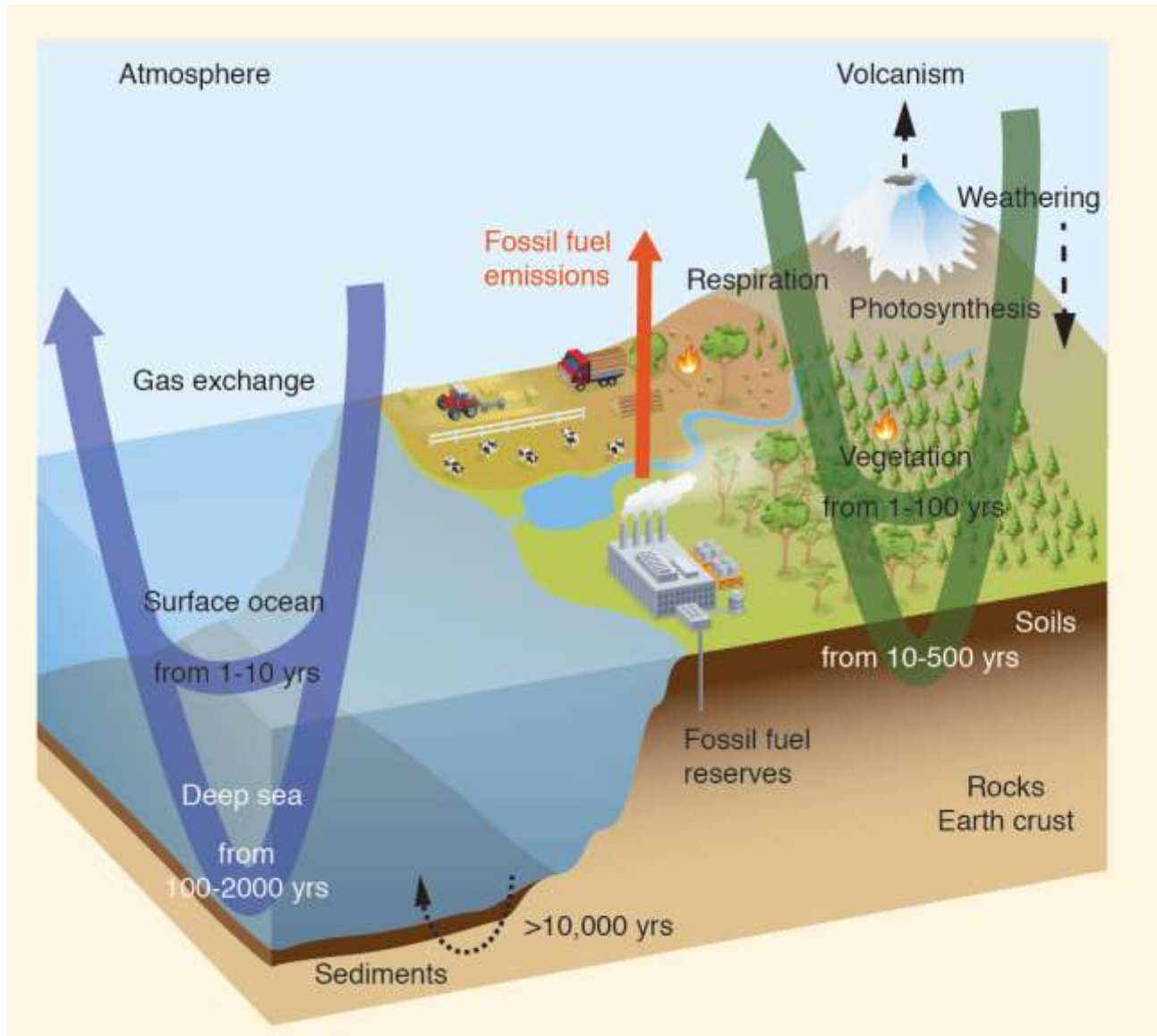
--pre-industrial (pre 1750)

--industrial period (1750-2011)

1 Petagram: 1Pg = 10<sup>12</sup> kg = 10<sup>15</sup> g



# Global Carbon Cycle - turn over times scales





## Box 6.1 | Multiple Residence Times for an Excess of Carbon Dioxide Emitted in the Atmosphere

**Box 6.1, Table 1** | The main natural processes that remove CO<sub>2</sub> consecutive to a large emission pulse to the atmosphere, their atmospheric CO<sub>2</sub> adjustment time scales, and main (bio)chemical reactions involved.

Processes	Time scale (years)	Reactions
Land uptake: Photosynthesis–respiration	1–10 <sup>2</sup>	$6\text{CO}_2 + 6\text{H}_2\text{O} + \text{photons} \rightarrow \text{C}_6\text{H}_{12}\text{O}_6 + 6\text{O}_2$ $\text{C}_6\text{H}_{12}\text{O}_6 + 6\text{O}_2 \rightarrow 6\text{CO}_2 + 6\text{H}_2\text{O} + \text{heat}$
Ocean invasion: Seawater buffer	10–10 <sup>3</sup>	$\text{CO}_2 + \text{CO}_3^{2-} + \text{H}_2\text{O} \rightleftharpoons 2\text{HCO}_3^-$
Reaction with calcium carbonate	10 <sup>3</sup> –10 <sup>4</sup>	$\text{CO}_2 + \text{CaCO}_3 + \text{H}_2\text{O} \rightarrow \text{Ca}^{2+} + 2\text{HCO}_3^-$
Silicate weathering	10 <sup>4</sup> –10 <sup>6</sup>	$\text{CO}_2 + \text{CaSiO}_3 \rightarrow \text{CaCO}_3 + \text{SiO}_2$

C<sub>6</sub>H<sub>12</sub>O<sub>6</sub>: Sugar

CaCO<sub>3</sub>: Calcium carbonate

Dissolved carbon dioxide: CO<sub>2(aq)</sub>

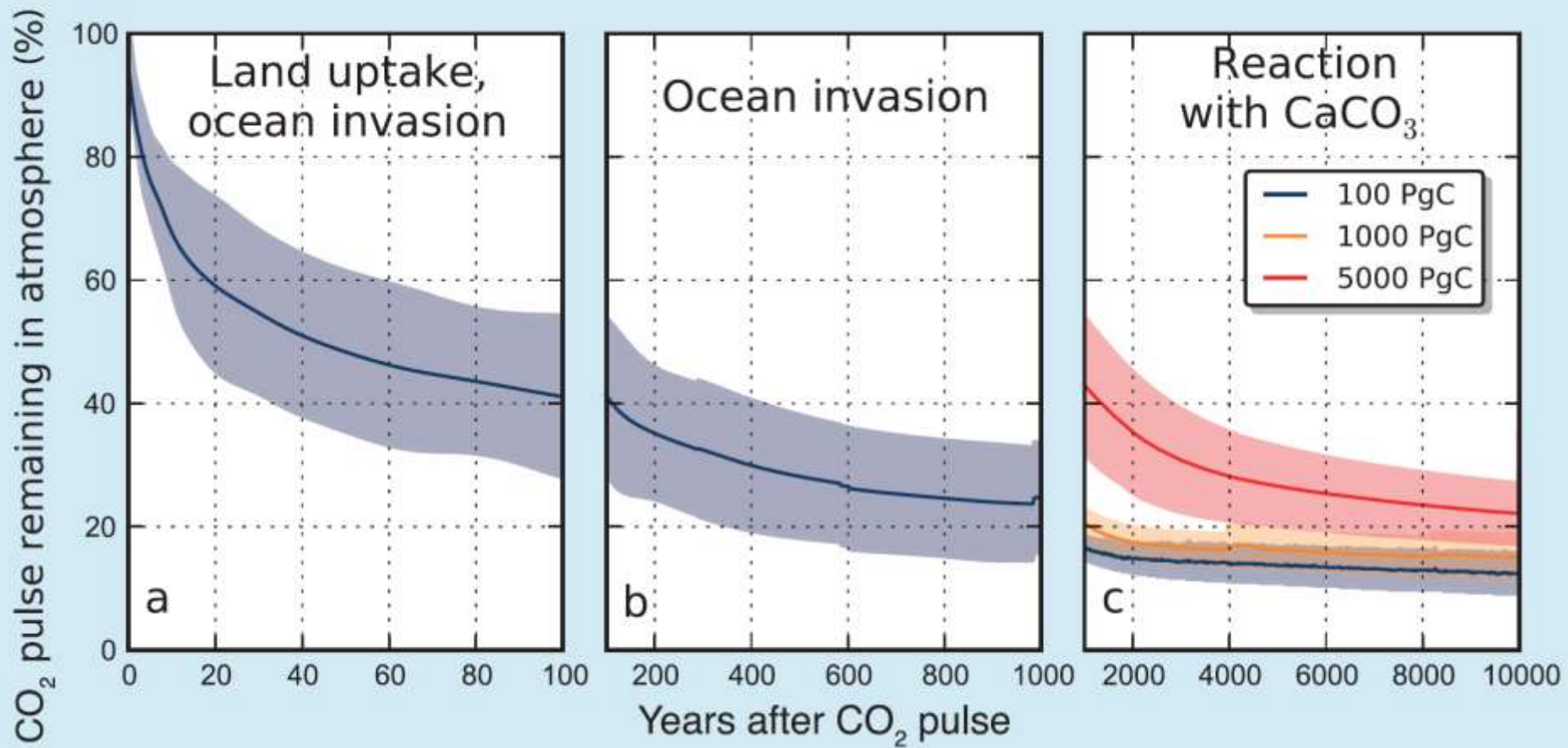
Bicarbonate HCO<sub>3</sub><sup>-</sup>

Carbonate ion: CO<sub>3</sub><sup>2-</sup>

Calcium inosilicate mineral (limestone): CaSiO<sub>3</sub>

Silicon dioxide (quartz): SiO<sub>2</sub>

## Box 6.1 | Multiple Residence Times for an Excess of Carbon Dioxide Emitted in the Atmosphere



Concept of a single, characteristic atmospheric lifetime is not applicable to CO<sub>2</sub> (→Chapter 8).

# Atmospheric CO<sub>2</sub> adjustment to anthropogenic carbon emissions

## Phase 1:

-**Within decades**, about 1/3 to 1/2 half of anthropogenic CO<sub>2</sub> goes into land and ocean, while the rest stays in the atmosphere.

-**Within centuries**, most of the anthropogenic CO<sub>2</sub> will be in the form of additional dissolved inorganic carbon in the ocean, decreasing ocean pH.

-**Within 1 kyr**, the remaining atmospheric fraction of the CO<sub>2</sub> emissions is between 15 and 40%, depending on the amount of carbon released. (>Carbonate buffer capacity of ocean decreases with higher CO<sub>2</sub> >larger cumulative emissions >higher remaining atmospheric fraction.)

## Phase 2:

-Within **several 1 kyr**, ocean pH will be restored by reaction of ocean dissolved CO<sub>2</sub> and calcium carbonate (CaCO<sub>3</sub>) of sea floor sediments, partly replenishing the buffer capacity of the ocean and further drawing down atmospheric CO<sub>2</sub> as a new balance is re-established between CaCO<sub>3</sub> sedimentation in the ocean and terrestrial weathering.

- Atmospheric CO<sub>2</sub> fraction down to 10 to 25% after about 10 kyr.

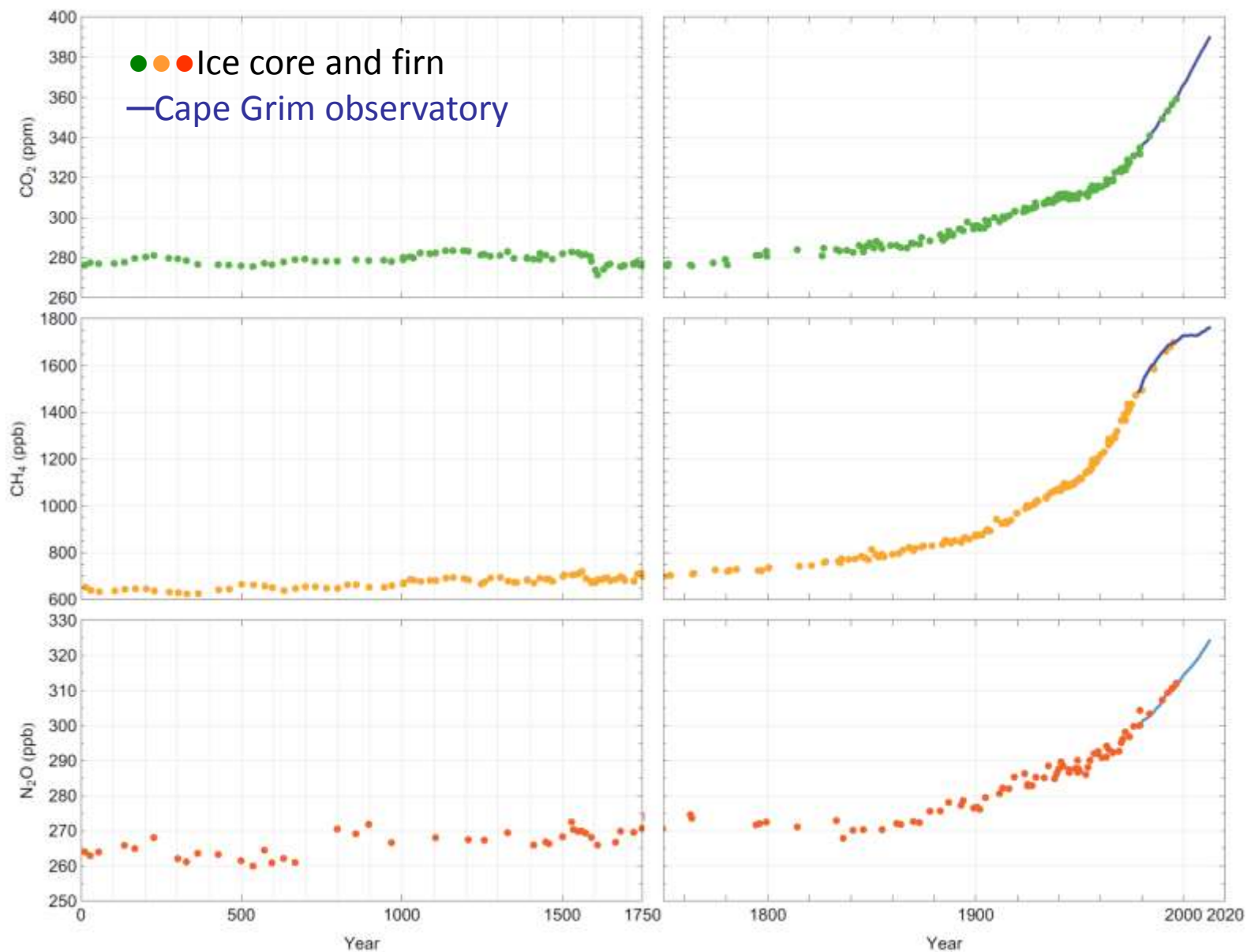
## Phase 3:

-Within **several 100 kyr**, rest of emitted CO<sub>2</sub> will be removed from atmosphere by silicate weathering (very slow process of CO<sub>2</sub> reaction with calcium silicate (CaSiO<sub>3</sub>) and other minerals of igneous rocks).

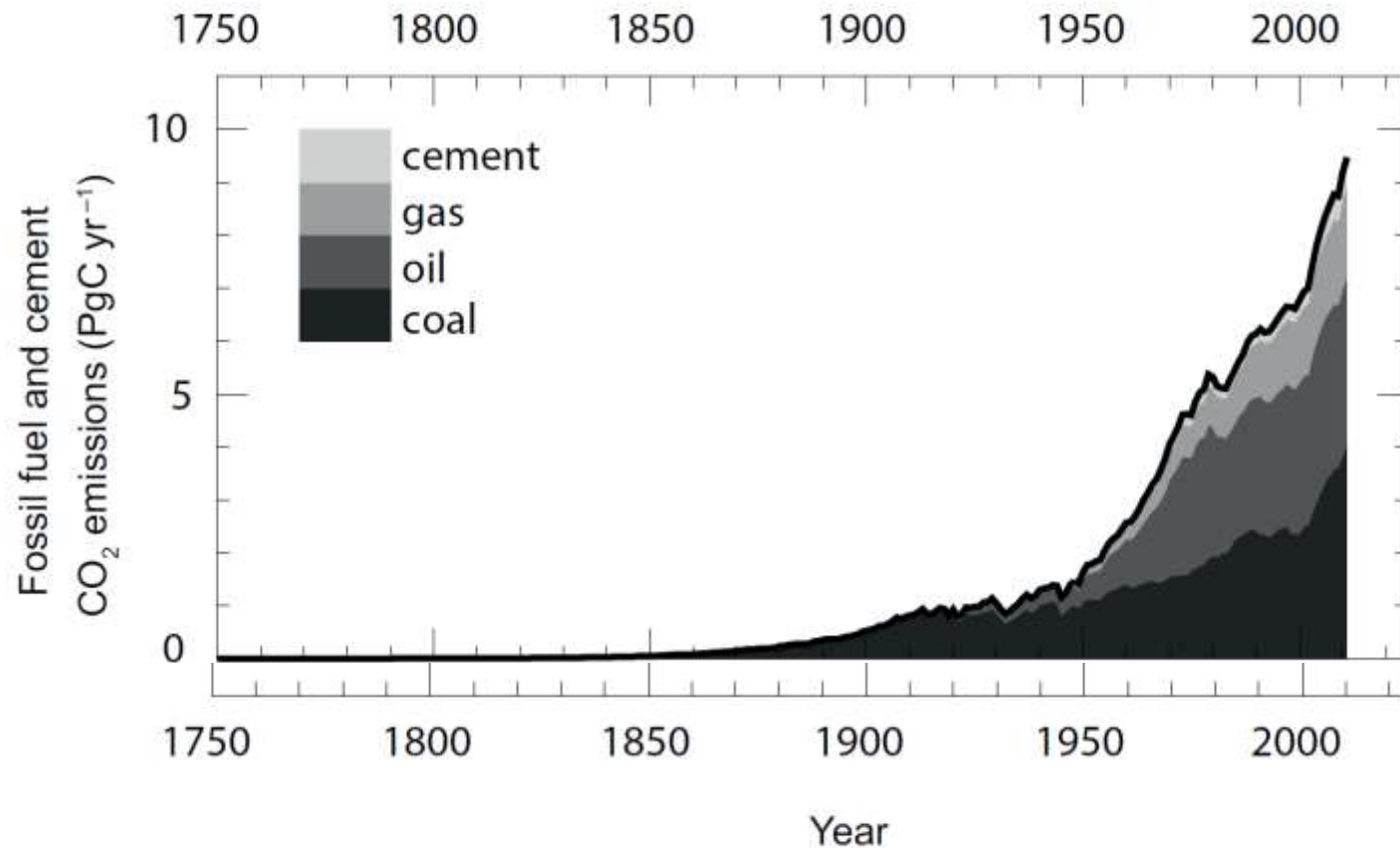
# The Carbon Cycle in the IPCC FAR

- Evolution of biogeochemical cycles since industrial era (*Section 6.3*)

# Atmospheric CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O pre-industrial era

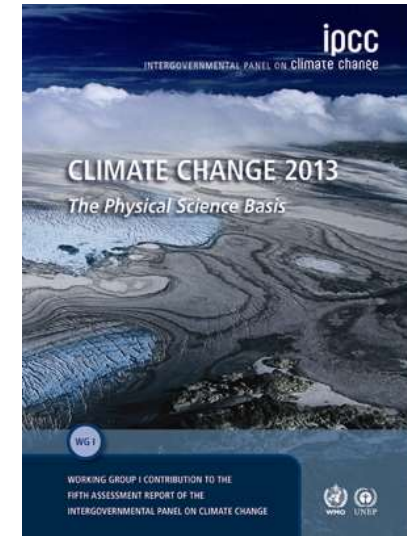


# Anthropogenic CO<sub>2</sub> emissions



# IPCC Chapter 6: Carbon and other biogeochemical cycles

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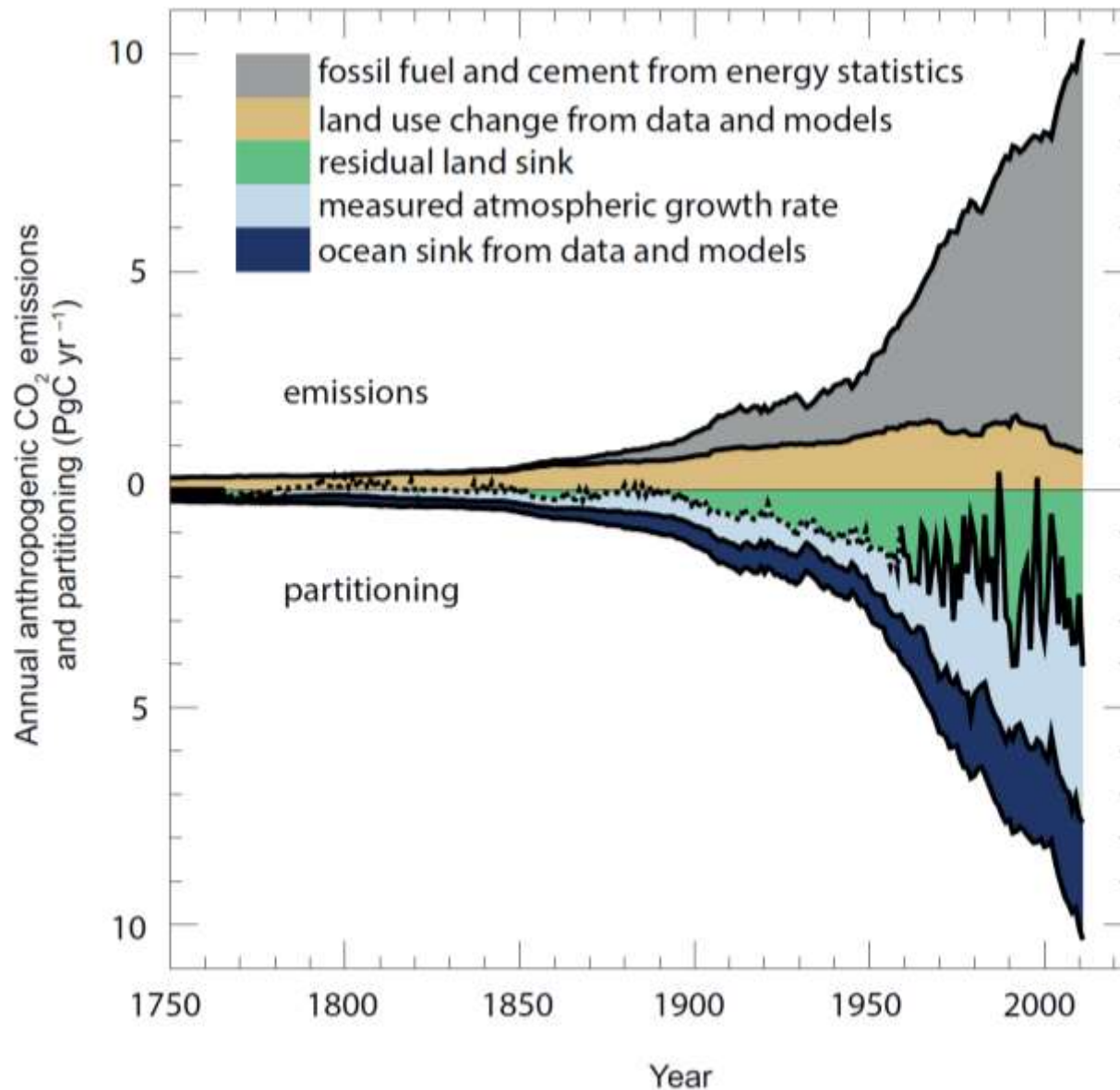


Ciais, P., et al., 2013: Carbon and Other Biogeochemical Cycles. In: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change. Cambridge University Press.





# Anthropogenic CO<sub>2</sub> emissions and partitioning





# Global anthropogenic CO<sub>2</sub> budget – time evolution

**Table 6.1** | Global anthropogenic CO<sub>2</sub> budget, accumulated since the Industrial Revolution (onset in 1750) and averaged over the 1980s, 1990s, 2000s, as well as the last 10 years until 2011. By convention, a negative ocean or land to atmosphere CO<sub>2</sub> flux is equivalent to a gain of carbon by these reservoirs. The table does not include natural exchanges (e.g., rivers, weathering) between reservoirs. The uncertainty range of 90% confidence interval presented here differs from how uncertainties were reported in AR4 (68%).

	1750–2011 Cumulative PgC	1980–1989 PgC yr <sup>-1</sup>	1990–1999 PgC yr <sup>-1</sup>	2000–2009 PgC yr <sup>-1</sup>	2002–2011 PgC yr <sup>-1</sup>
Atmospheric increase <sup>a</sup>	240 ± 10 <sup>f</sup>	3.4 ± 0.2	3.1 ± 0.2	4.0 ± 0.2	4.3 ± 0.2
Fossil fuel combustion and cement production <sup>b</sup>	375 ± 30 <sup>f</sup>	5.5 ± 0.4	6.4 ± 0.5	7.8 ± 0.6	8.3 ± 0.7
Ocean-to-atmosphere flux <sup>c</sup>	-155 ± 30 <sup>f</sup>	-2.0 ± 0.7	-2.2 ± 0.7	-2.3 ± 0.7	-2.4 ± 0.7
Land-to-atmosphere flux <i>Partitioned as follows</i>	30 ± 45 <sup>f</sup>	-0.1 ± 0.8	-1.1 ± 0.9	-1.5 ± 0.9	-1.6 ± 1.0
Net land use change <sup>d</sup>	180 ± 80 <sup>g</sup>	1.4 ± 0.8	1.5 ± 0.8	1.1 ± 0.8	0.9 ± 0.8
Residual land sink <sup>e</sup>	-160 ± 90 <sup>f</sup>	-1.5 ± 1.1	-2.6 ± 1.2	-2.6 ± 1.2	-2.5 ± 1.3

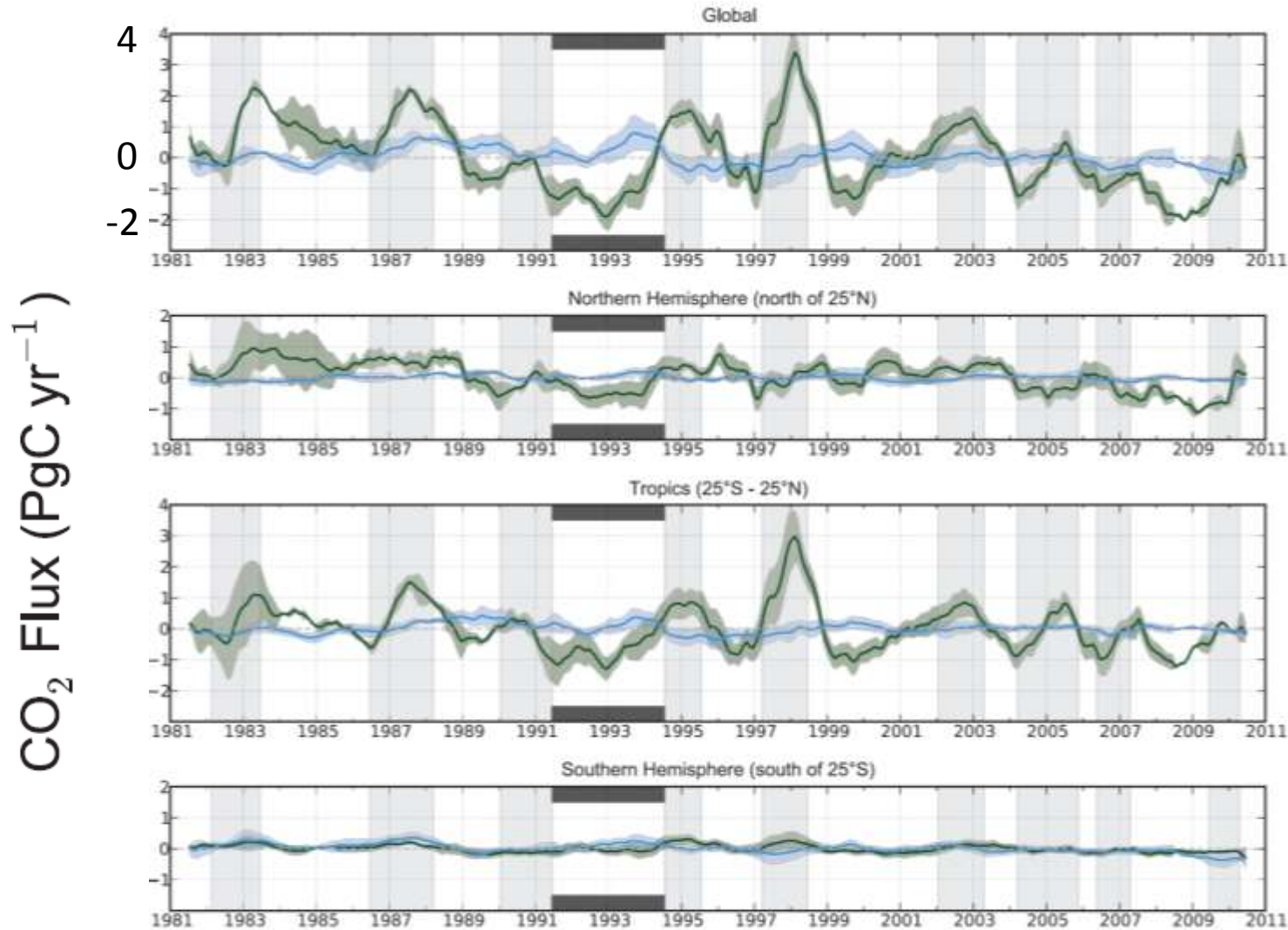
- Varying atmospheric CO<sub>2</sub> sources growth; mainly increasing since the last three decades.
- Ocean CO<sub>2</sub> sink is also slightly increasing.

# Carbon Dioxide from Fossil Fuel Combustion

This video shows the accumulation of carbon dioxide (CO<sub>2</sub>) in the atmosphere from the burning of coal, oil, and natural gas (fossil fuels) over a two-year time period (2011-2012). The video begins with a map of the world with no fossil fuel CO<sub>2</sub>. As time progresses the viewer watches the global accumulation of CO<sub>2</sub> emissions from all fossil fuel sources. Large emitters such as Eastern Asia, Western Europe, and the North East of North America stand out. By the end of 2012, the entire Northern Hemisphere is red, illustrating a total accumulation of about 9 to 10 ppm of CO<sub>2</sub>.

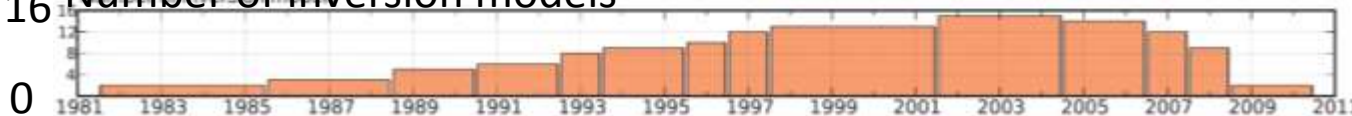
[www.esrl.noaa.gov/gmd/ccgg/trends/ff.html](http://www.esrl.noaa.gov/gmd/ccgg/trends/ff.html)

# CO<sub>2</sub> flux anomalies – inversion modeling

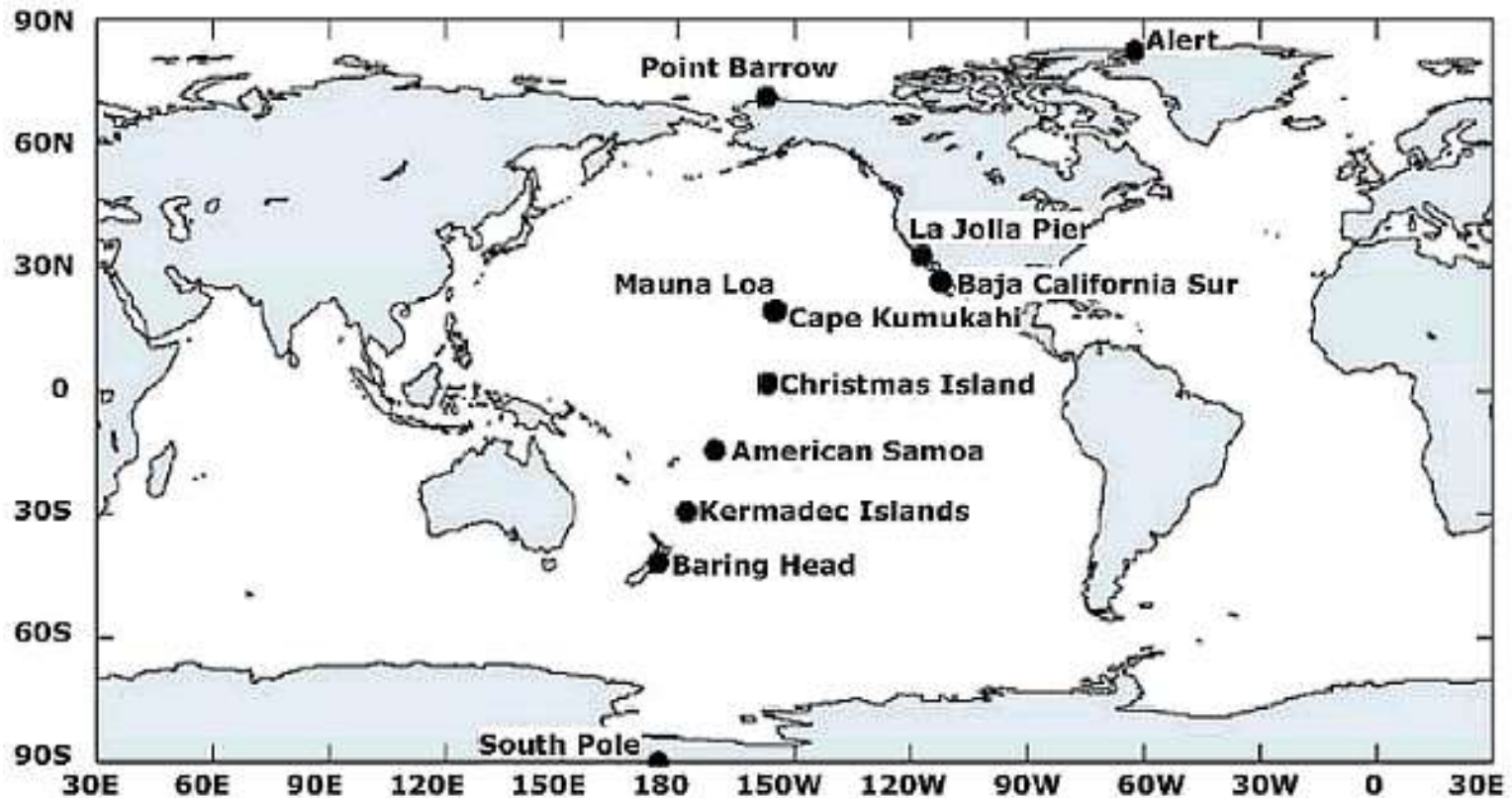


–land-atmosphere  
–ocean-atmosphere  
+/- 1  $\sigma$  model  
El Niño  
**Pinatubo**

16 Number of inversion models

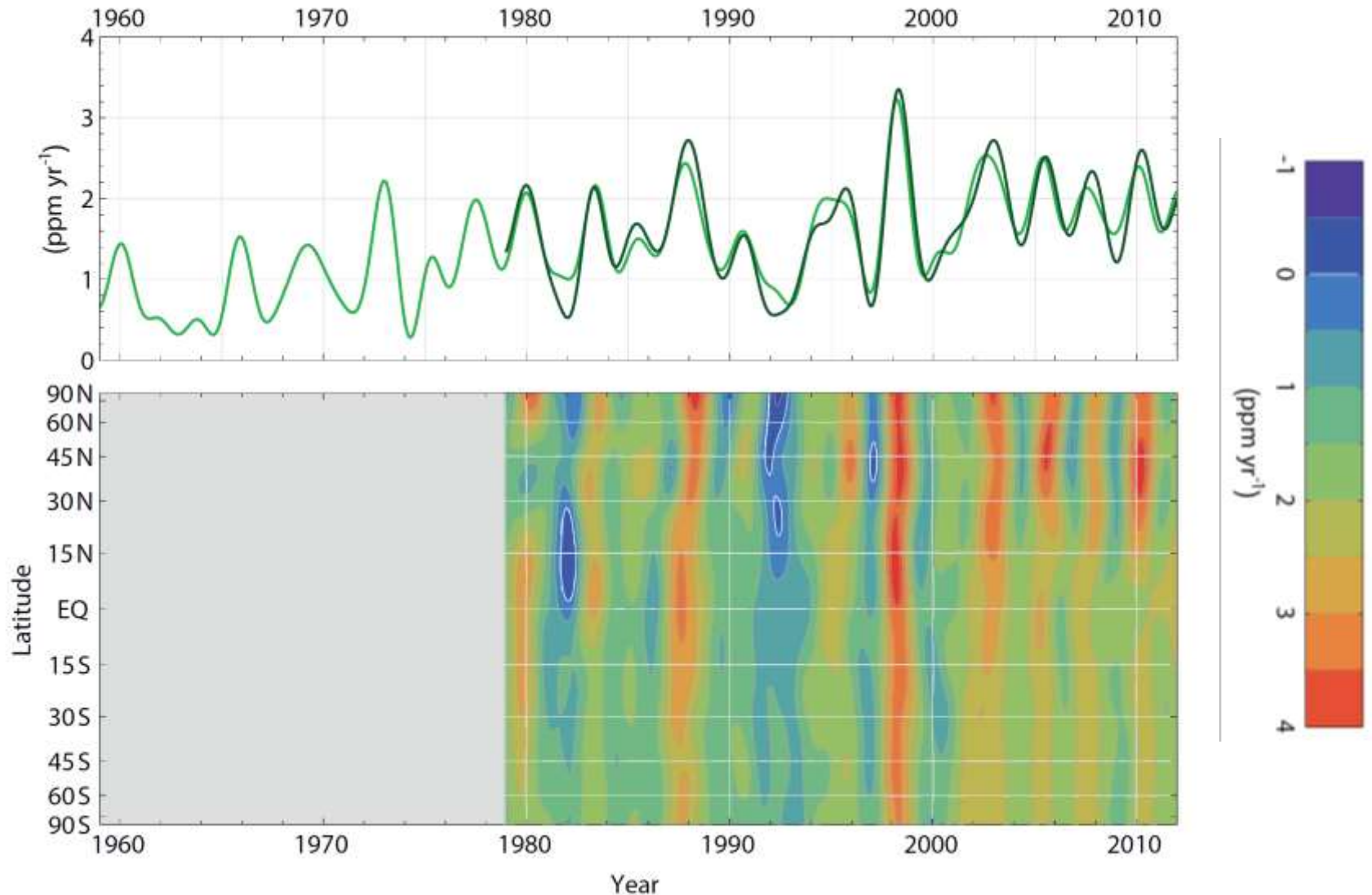


# Atmospheric CO<sub>2</sub> Records from monitoring sites of the Scripps Institution of Oceanography (SIO)



# Atmospheric CO<sub>2</sub> growth rate [ppm/yr]

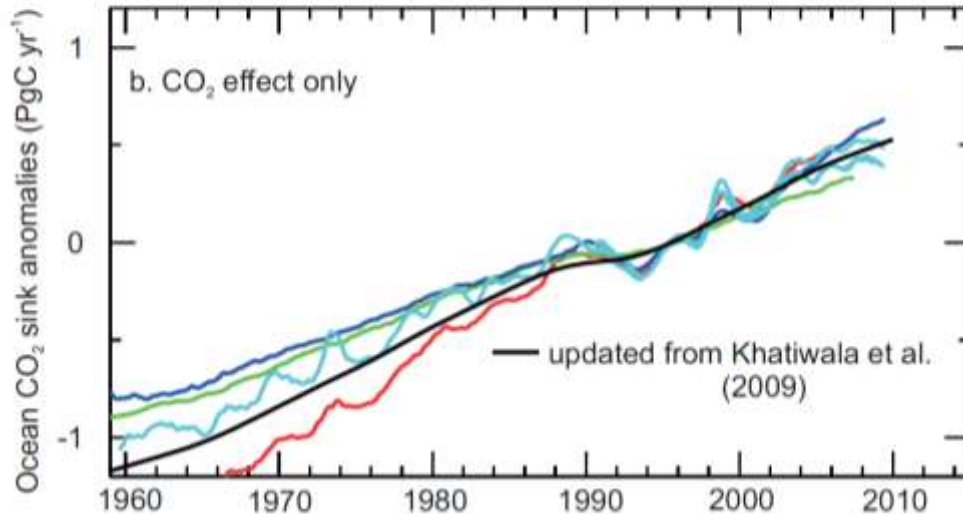
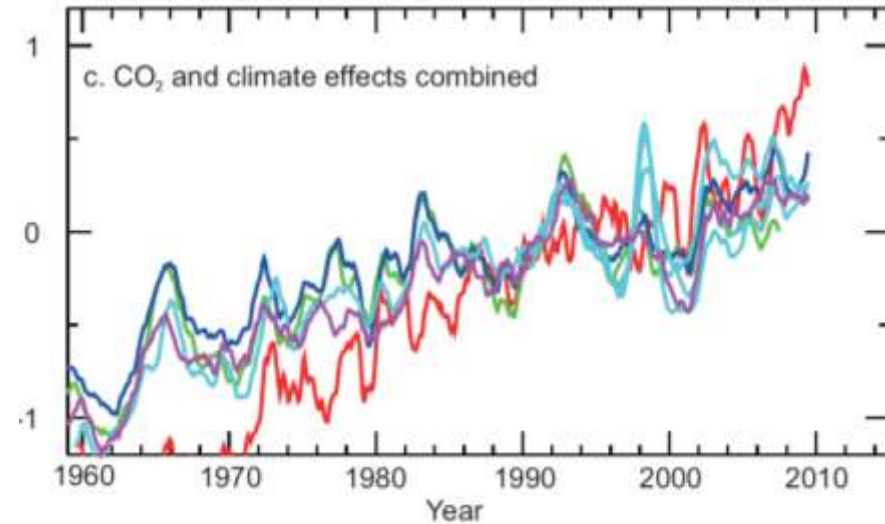
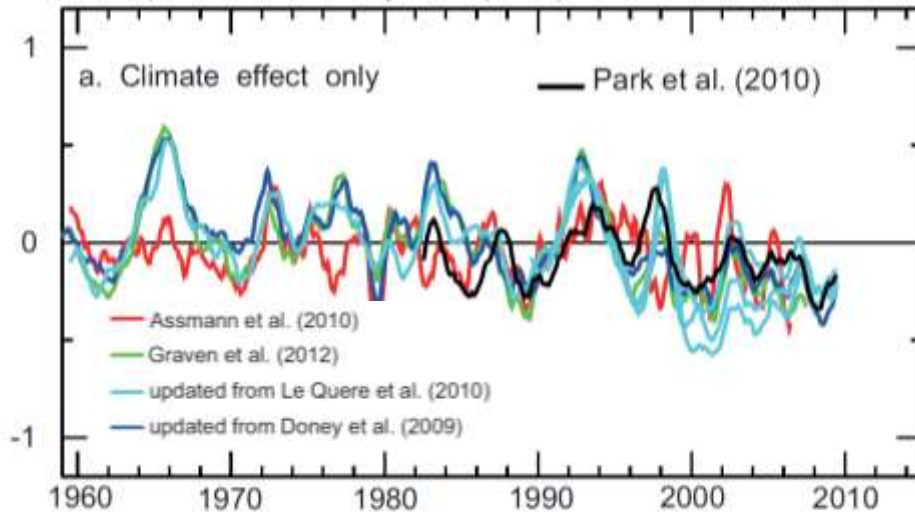
- SIO and NOAA GMD marine observations





# Ocean CO<sub>2</sub> flux variations - temporal

Anomalous CO<sub>2</sub> ocean-to-atmosphere fluxes wrt 1990-2000 avgs



— Ocean models

-- Indirect methods based on observations

-Low interannual variability.

-Ocean takes up more CO<sub>2</sub> during El Nino due to the suppression of CO<sub>2</sub> ota fluxes over the East Pacific upwelling.

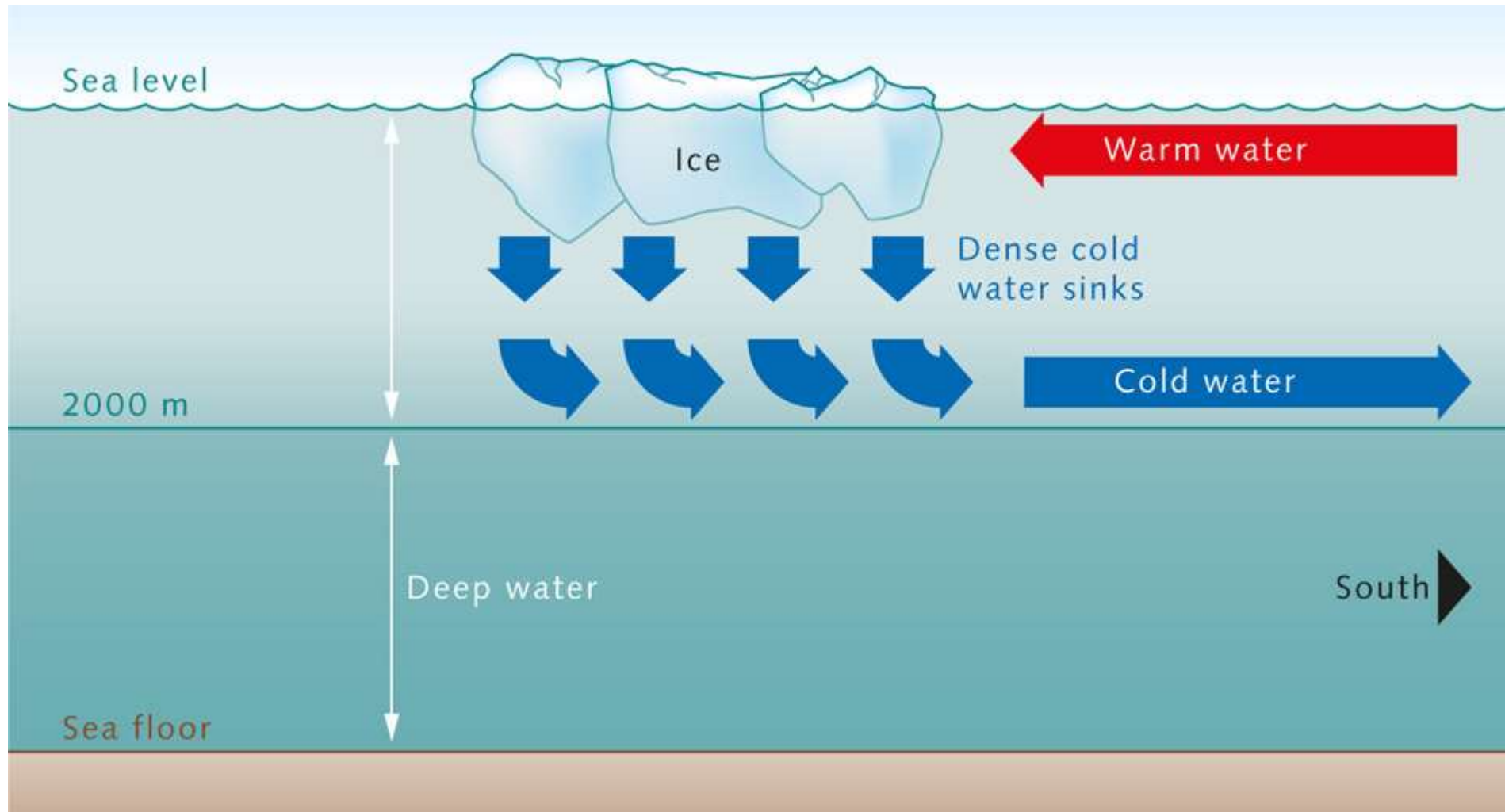
# Ocean inorganic carbon storage – regional changes

Section	Time	Storage Rate (mol C m <sup>-2</sup> yr <sup>-1</sup> )	Data Source
<b>Global average (used in Table 6.1)</b>	2007–2008	0.5 ± 0.2	Khatiwala et al. (2009)
<b>Pacific Ocean</b>			
Section along 30°S	1992–2003	1.0 ± 0.4	Murata et al. (2007)
N of 50°S, 120°W to 180°W	1974–1996	0.9 ± 0.3	Peng et al. (2003)
154°W, 20°N to 50°S	1991–2006	0.6 ± 0.1	Sabine et al. (2008)
140°E to 170°W, 45°S to 65°S	1968–1991/1996	0.4 ± 0.2	Matear and McNeil (2003)
149° W, 4°S to 10°N	1993–2005	0.3 ± 0.1	Murata et al. (2009)
149° W, 24°N to 30°N	1993–2005	0.6 ± 0.2	Murata et al. (2009)
Northeast Pacific	1973–1991	1.3 ± 0.5	Peng et al. (2003)
–160°E, –45°N	1997–2008	0.4 ± 0.1	Wakita et al. (2010)
North of 20°N	1994–2004/2005	0.4 ± 0.2	Sabine et al. (2008)
150°W, 20°S to 20°N	1991/1992–2006	0.3 ± 0.1	Sabine et al. (2008)
<b>Indian Ocean</b>			
<p>Anthropogenic Carbon uptake is observed to be larger over the high latitude oceans because of the more vigorous convection there.</p>			
<b>Atlantic Ocean</b>			
Section along 30°S	1992/1993–2003	0.6 ± 0.1	Murata et al. (2010)
–30°W, 56°S to 15°S	1989–2005	0.8	Wanninkhof et al. (2010)
20°W, 64°N to 15°N	1993–2003	0.6	Wanninkhof et al. (2010)
–25°W, 15°N to 15°S	1993–2003	0.2	Wanninkhof et al. (2010)
40°N to 65°N	1981–1997/1999	2.2 ± 0.7	Friis et al. (2005)
20°N to 40°N	1981–2004	1.2 ± 0.3	Tanhua et al. (2007)
Nordic Seas	1981–2002/2003	0.9 ± 0.2	Olsen et al. (2006)
<b>Sub-decadal variations</b>			
Irminger Sea	1981–1991	0.6 ± 0.4	Pérez et al. (2008)
Irminger Sea	1991–1996	2.3 ± 0.6	Pérez et al. (2008)
Irminger Sea	1997–2006	0.8 ± 0.2	Pérez et al. (2008)

**Table 6.5 |** Regional rates of change in inorganic carbon storage from shipboard repeated hydrographic cross sections.

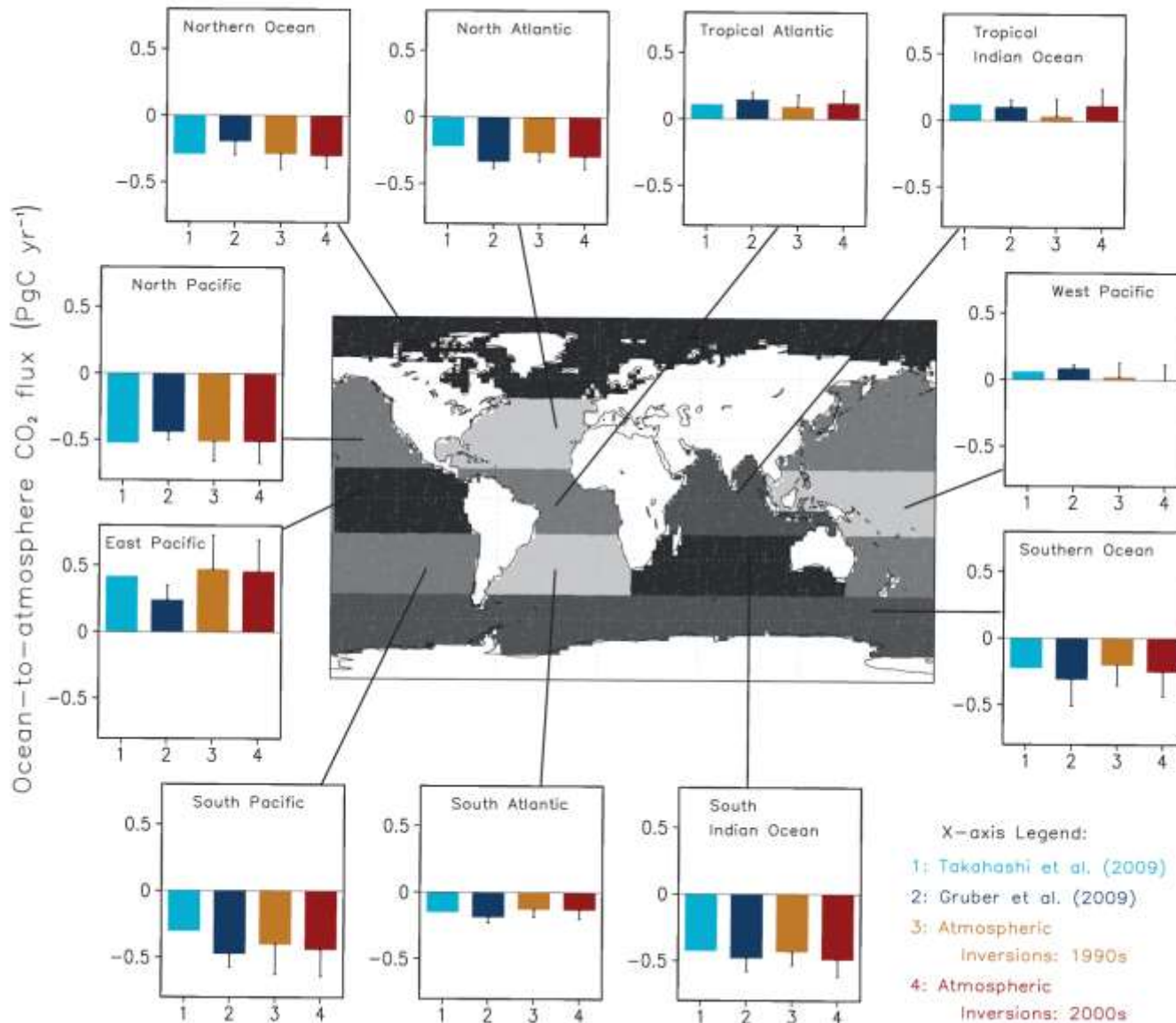


# Ocean convection



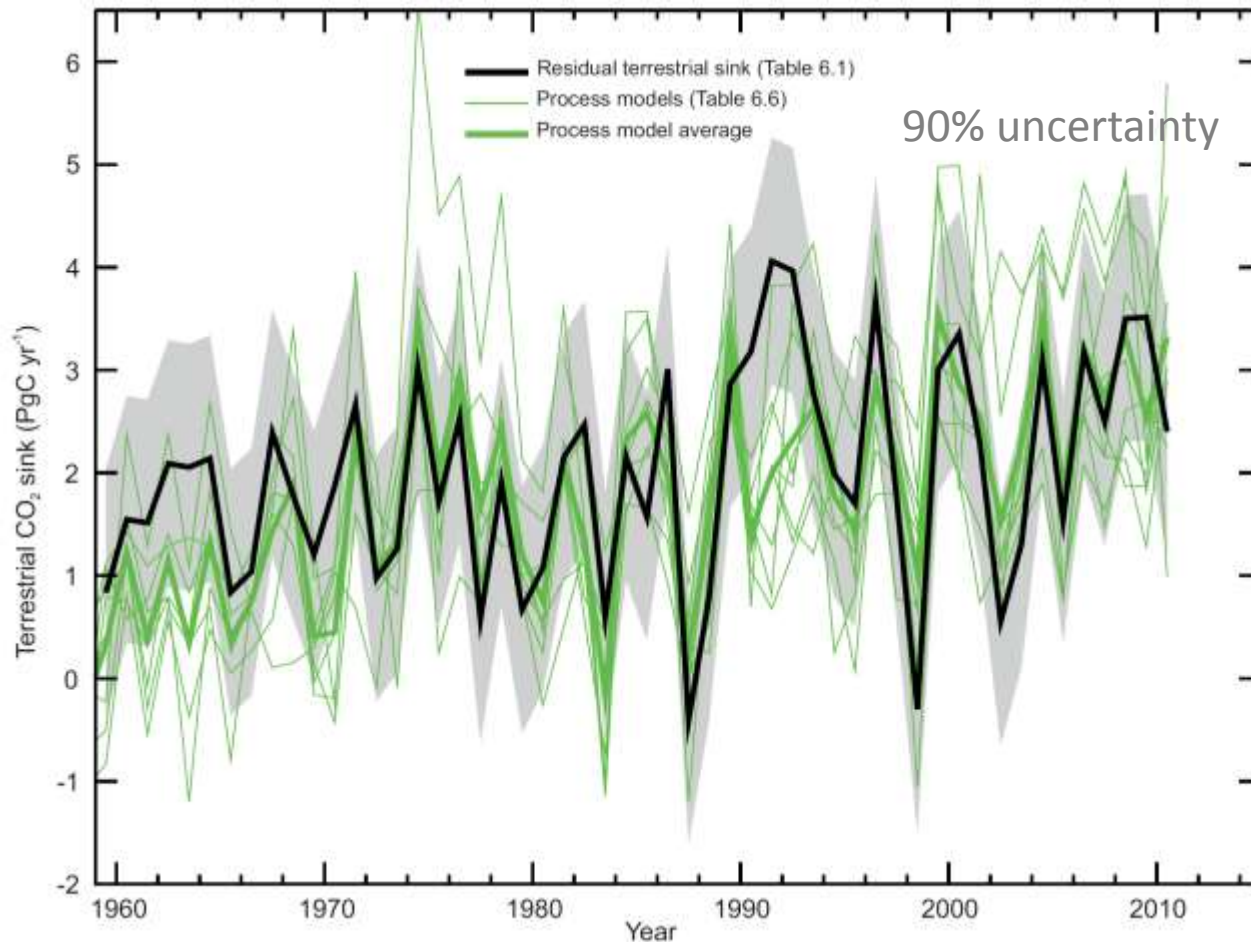
# Decadal oceanic CO<sub>2</sub> fluxes

2000s-1990s



# Terrestrial CO<sub>2</sub> sink – temporal

## models wo land use change



- Large interannual variability; large uncertainties.
- Tropical land ecosystems dominate global CO<sub>2</sub> variability.
- During El Nino/La Nina enhanced land CO<sub>2</sub> source/sink.
- ENSO-Volcanic Index time series explains 75% of the observed variability.



### Box 6.3 | The Carbon Dioxide Fertilisation Effect

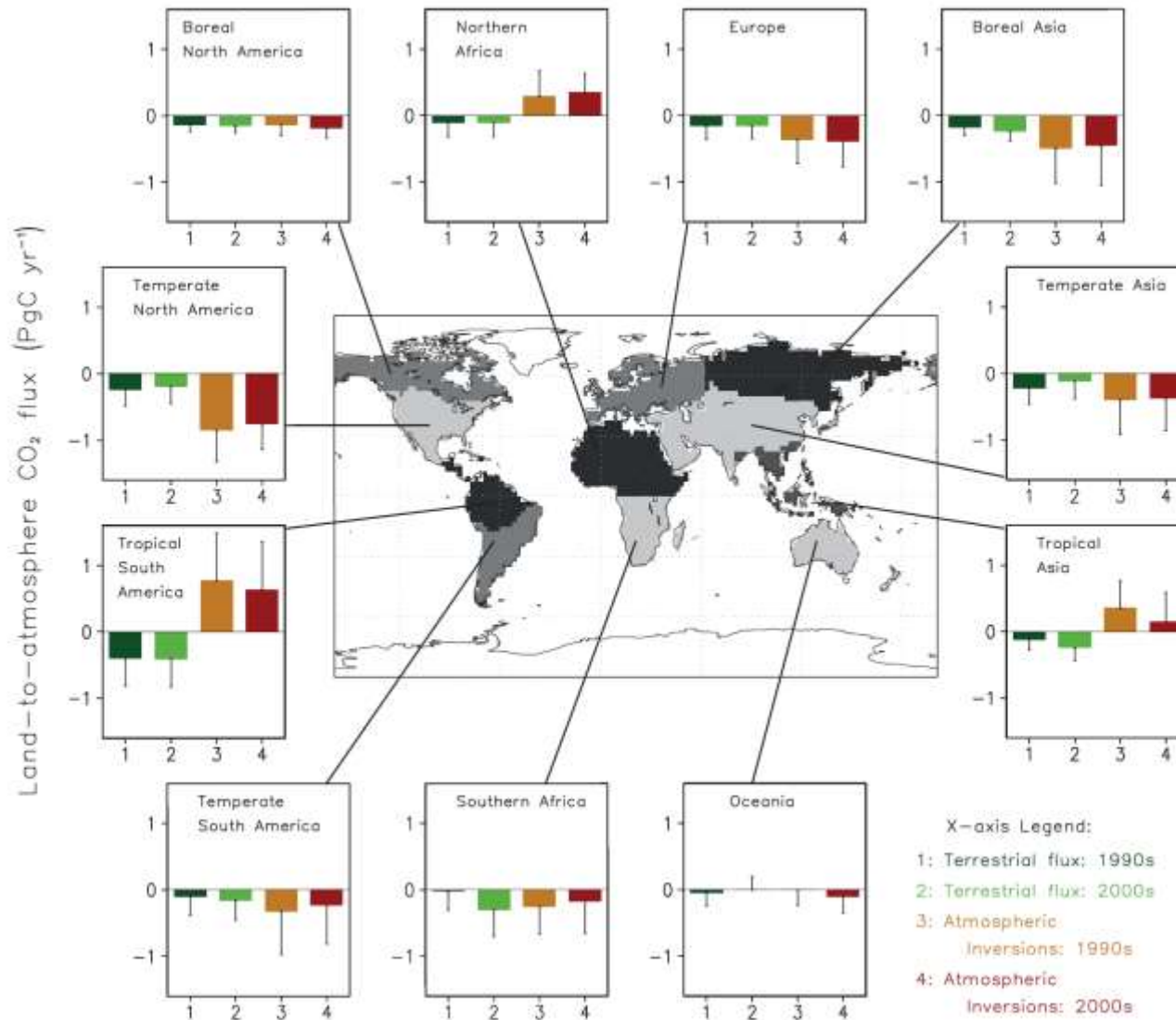
Elevated atmospheric CO<sub>2</sub> concentrations lead to higher leaf photosynthesis and reduced canopy transpiration, which in turn lead to increased plant water use efficiency and reduced fluxes of surface latent heat. The increase in leaf photosynthesis with rising CO<sub>2</sub>, the so-called CO<sub>2</sub> fertilisation effect, plays a dominant role in terrestrial biogeochemical models to explain the global land carbon sink (Sitch et al., 2008), yet it is one of most unconstrained process in those models.

Field experiments provide a direct evidence of increased photosynthesis rates and water use efficiency (plant carbon gains per unit of water loss from transpiration) in plants growing under elevated CO<sub>2</sub>. These physiological changes translate into a broad range of higher plant carbon accumulation in more than two-thirds of the experiments and with increased net primary productivity (NPP) of about 20 to 25% at double CO<sub>2</sub> from pre-industrial concentrations (Ainsworth and Long, 2004; Luo et al., 2004, 2006; Nowak et al., 2004; Norby et al., 2005; Canadell et al., 2007a; Denman et al., 2007; Ainsworth et al., 2012; Wang et al., 2012a). Since the AR4, new evidence is available from long-term Free-air CO<sub>2</sub> Enrichment (FACE) experiments in temperate ecosystems showing the capacity of ecosystems exposed to elevated CO<sub>2</sub> to sustain higher rates of carbon accumulation over multiple years (Liberloo et al., 2009; McCarthy et al., 2010; Aranjuelo et al., 2011; Dawes et al., 2011; Lee et al., 2011; Zak et al., 2011). However, FACE experiments also show the diminishing or lack of CO<sub>2</sub> fertilisation effect in some ecosystems and for some plant species (Dukes et al., 2005; Adair et al., 2009; Bader et al., 2009; Norby et al., 2010; Newingham et al., 2013). This lack of response occurs despite increased water use efficiency, also confirmed with tree ring evidence (Gedalof and Berg, 2010; Peñuelas et al., 2011).

Nutrient limitation is hypothesized as primary cause for reduced or lack of CO<sub>2</sub> fertilisation effect observed on NPP in some experiments (Luo et al., 2004; Dukes et al., 2005; Finzi et al., 2007; Norby et al., 2010). Nitrogen and phosphorus are *very likely* to play the most important role in this limitation of the CO<sub>2</sub> fertilisation effect on NPP, with nitrogen limitation prevalent in temperate and boreal ecosystems, and phosphorus limitation in the tropics (Luo et al., 2004; Vitousek et al., 2010; Wang et al., 2010a; Goll et al., 2012). Micronutrients interact in diverse ways with other nutrients in constraining NPP such as molybdenum and phosphorus in the tropics (Wurzburger et al., 2012). Thus, with *high confidence*, the CO<sub>2</sub> fertilisation effect will lead to enhanced NPP, but significant uncertainties remain on the magnitude of this effect, given the lack of experiments outside of temperate climates.

# Decadal terrestrial CO<sub>2</sub> fluxes - regional

2000s-1990s



# Regional CO<sub>2</sub> budgets

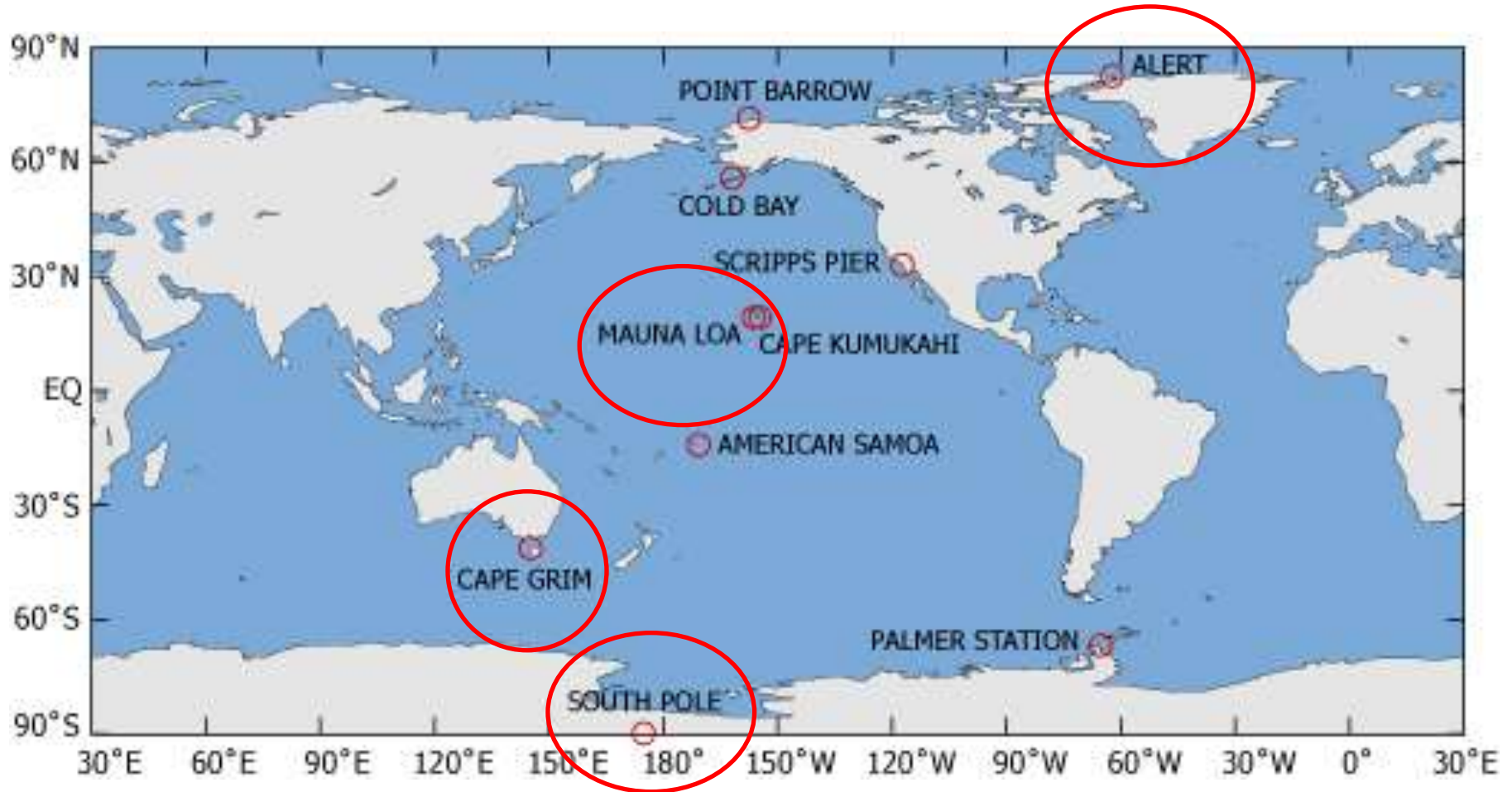
**Table 6.6 |** Regional CO<sub>2</sub> budgets using top-down estimates (atmospheric inversions) and bottom-up estimates (inventory data, biogeochemical modelling, eddy-covariance), excluding fossil fuel emissions. A positive sign indicates a flux from the atmosphere to the land (i.e., a land sink).

Region	CO <sub>2</sub> Sink (PgC yr <sup>-1</sup> )	Uncertainty <sup>a</sup>	Period	Reference
Arctic Tundra	0.1	±0.3 <sup>b</sup>	2000–2006	McGuire et al. (2012)
Australia	0.04	±0.03	1990–2009	Haverd et al. (2013)
East Asia	0.25	±0.1	1990–2009	Piao et al. (2012)
Europe	0.9	±0.2	2001–2005	Luyssaert et al. (2012)
North America	0.6	±0.02	2000–2005	King et al. (2012)
Russian Federation	0.6	–0.3 to –1.3	1990–2007	Dolman et al. (2012)
South Asia	0.15	±0.24	2000–2009	Patra et al. (2013)
South America	–0.3	±0.3	2000–2005	Gloor et al. (2012)

Notes:

- <sup>a</sup> One standard deviation from mean unless indicated otherwise.
- <sup>b</sup> Based on range provided.

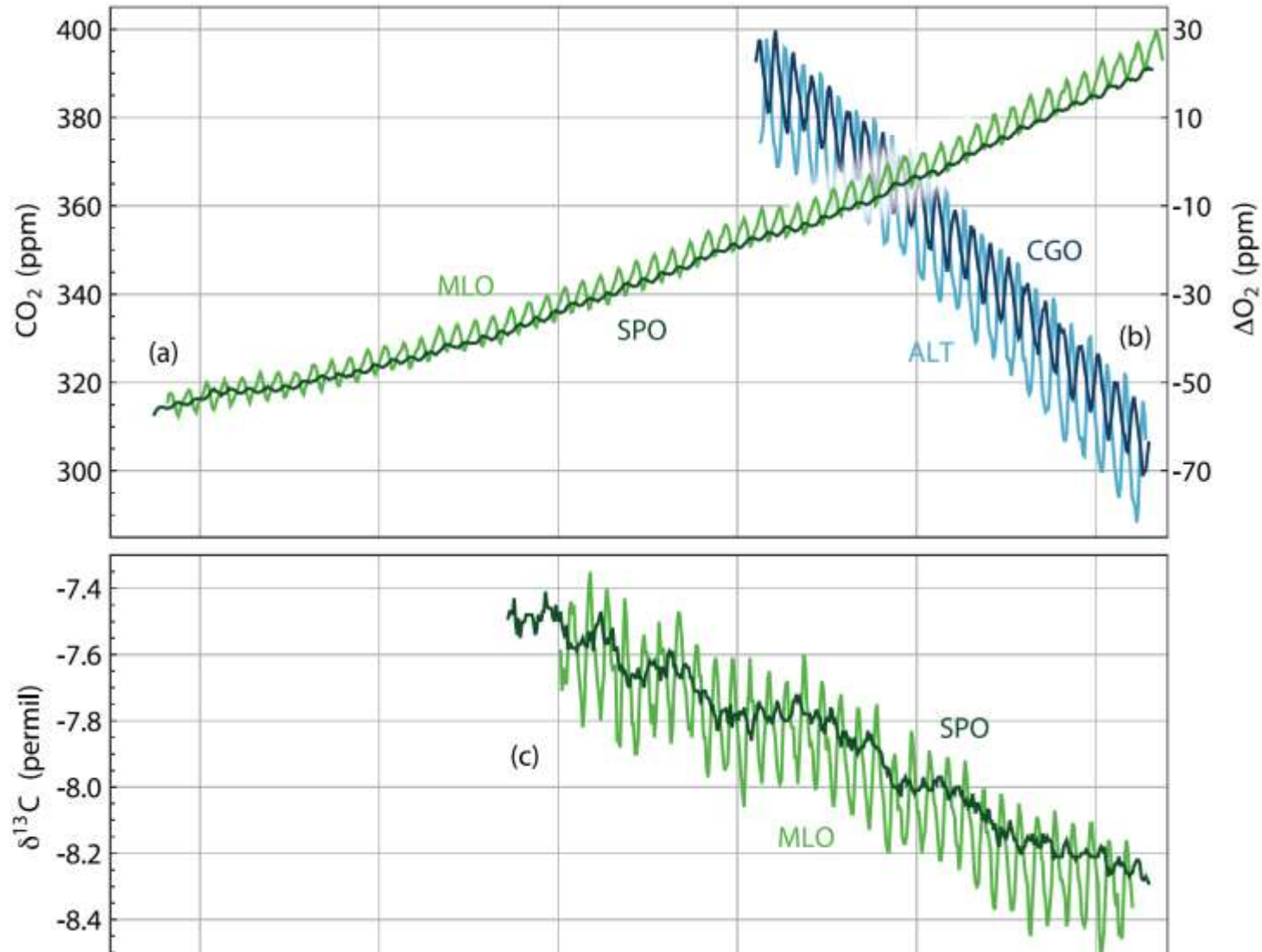




MLO: Mauna Loa Observatory SPO: South Pole Observatory  
ALT: Alert NH CGO: Cape Grim Observatory



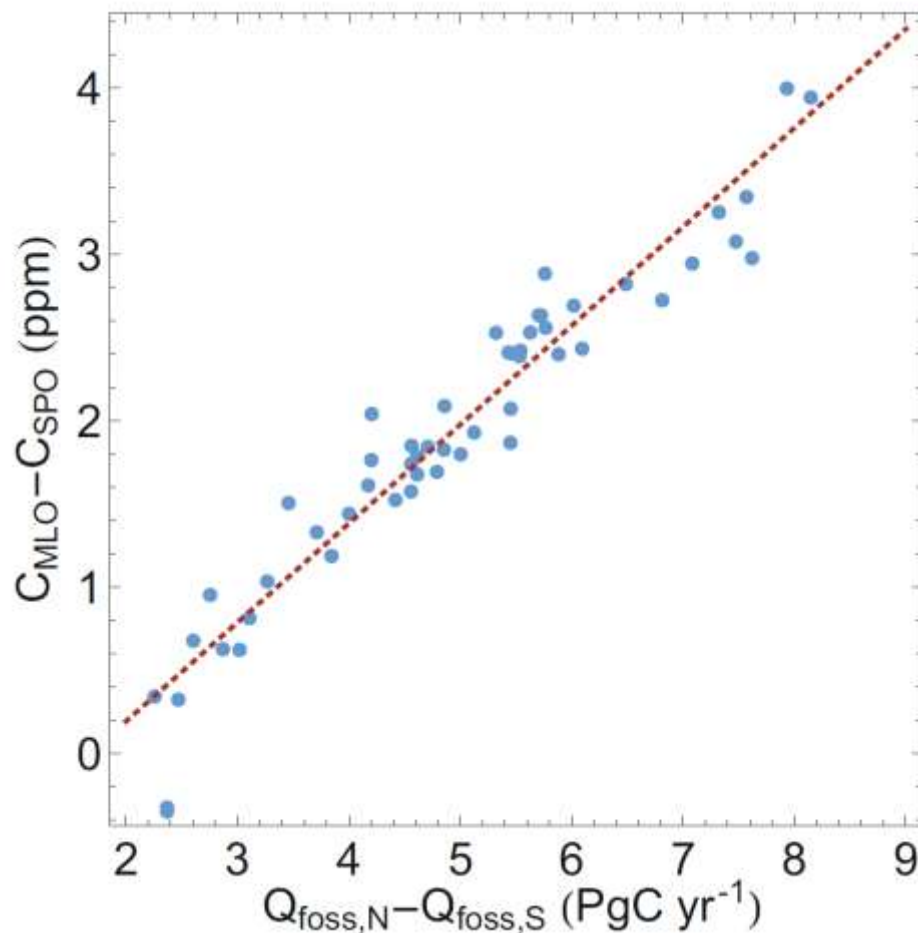
# Atmospheric concentrations of CO<sub>2</sub>, O<sub>2</sub>, <sup>13</sup>C/<sup>12</sup>C stable isotope ratio in CO<sub>2</sub>



CO<sub>2</sub> from fossil fuels and from land biosphere has a lower <sup>13</sup>C/<sup>12</sup>C stable isotope ratio than the CO<sub>2</sub> in the atmosphere.

>decreasing trend in atmospheric <sup>13</sup>C/<sup>12</sup>C ratio of atmospheric CO<sub>2</sub> concentration.

# Atmospheric CO<sub>2</sub> concentrations [ppm] versus fossil fuel combustion emissions [PgC/yr]: NH-SH differences



Fossil fuel CO<sub>2</sub> emissions take place mainly in industrialised countries (NH).

>Atmospheric CO<sub>2</sub> concentrations are higher in the NH than for the SH.

>Atmosphere CO<sub>2</sub> concentration diff follows linearly fossil fuel combustion emissions gradient.

# Summary

“With a **very high confidence**, the increase in **CO<sub>2</sub> emissions** from **fossil fuel burning** and those arising from **land use change** are the **dominant cause** of the observed **increase in atmospheric CO<sub>2</sub> concentration**. Several lines of evidence support this conclusion:

- The **observed decrease in atmospheric O<sub>2</sub> content** over past two decades and the **lower O<sub>2</sub> content in the northern compared to the SH** are **consistent with the burning of fossil fuels**.
- **CO<sub>2</sub> from fossil fuels and from the land biosphere has a lower <sup>13</sup>C/<sup>12</sup>C stable isotope ratio** than the CO<sub>2</sub> in the atmosphere. This induces a **decreasing temporal trend in the atmospheric <sup>13</sup>C/<sup>12</sup>C ratio of atmospheric CO<sub>2</sub> concentration** as well as, on annual average, slightly lower <sup>13</sup>C/<sup>12</sup>C values in the NH (Figure 6.3). These signals are measured.
- Most of the fossil fuel CO<sub>2</sub> emissions take place in the industrialised countries north of the equator. Consistent with this, on annual average, **atmospheric CO<sub>2</sub> measurement stations in the NH record increasingly higher CO<sub>2</sub> concentrations than stations in the SH**, as witnessed by the observations from Mauna Loa, Hawaii, and the South Pole.”

# Carbon Budget 2014 – Executive Summary

## All the data is shown in billion tonnes CO<sub>2</sub> (GtCO<sub>2</sub>)

1 Gigatonne (Gt) = 1 billion tonnes =  $1 \times 10^{15}$ g = 1 Petagram (Pg)

1 kg carbon (C) = 3.664 kg carbon dioxide (CO<sub>2</sub>)

1 GtC = 3.664 billion tonnes CO<sub>2</sub> = 3.664 GtCO<sub>2</sub>

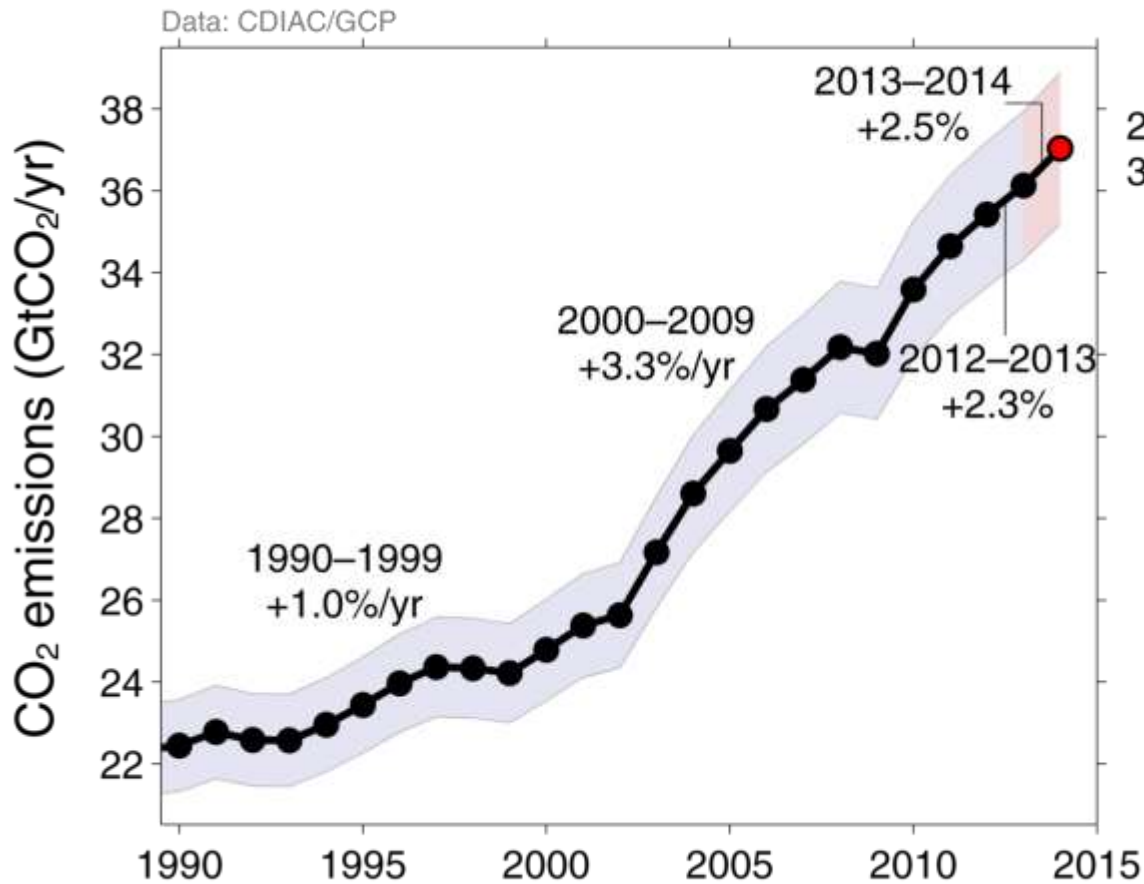
### Disclaimer

The Global Carbon Budget and the information presented here are intended for those interested in learning about the carbon cycle, and how human activities are changing it. The information contained herein is provided as a public service, with the understanding that the Global Carbon Project team make no warranties, either expressed or implied, concerning the accuracy, completeness, reliability, or suitability of the information.

# Fossil Fuel and Cement Emissions

Global fossil fuel and cement emissions:  $36.1 \pm 1.8$  GtCO<sub>2</sub> in 2013, 61% over 1990

● Projection for 2014 :  $37.0 \pm 1.9$  GtCO<sub>2</sub>, 65% over 1990

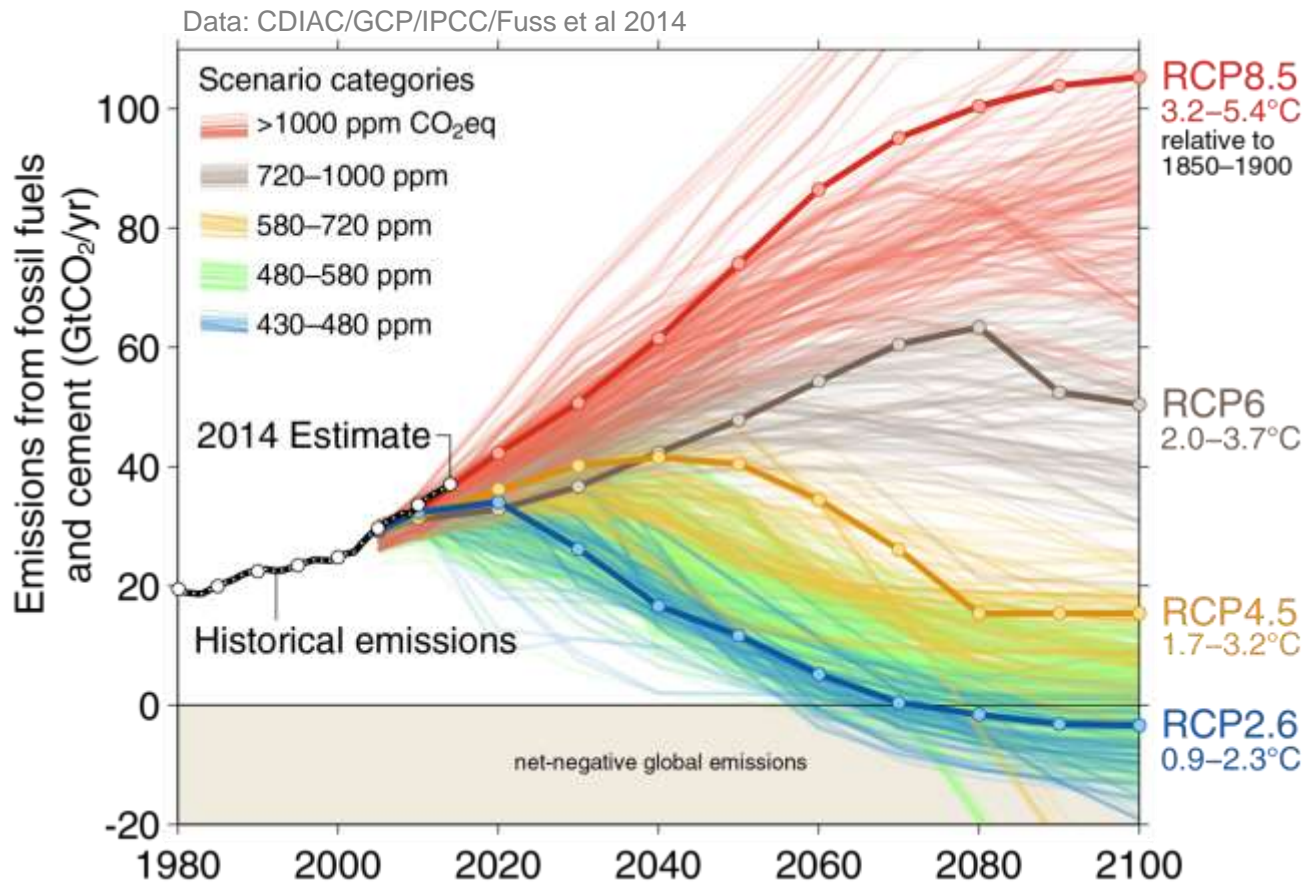


Uncertainty is  $\pm 5\%$  for one standard deviation (IPCC "likely" range)



# Observed Emissions and Emissions Scenarios

Emissions are on track for 3.2–5.4°C “likely” increase in temperature above pre-industrial  
 Large and sustained mitigation is required to keep below 2°C

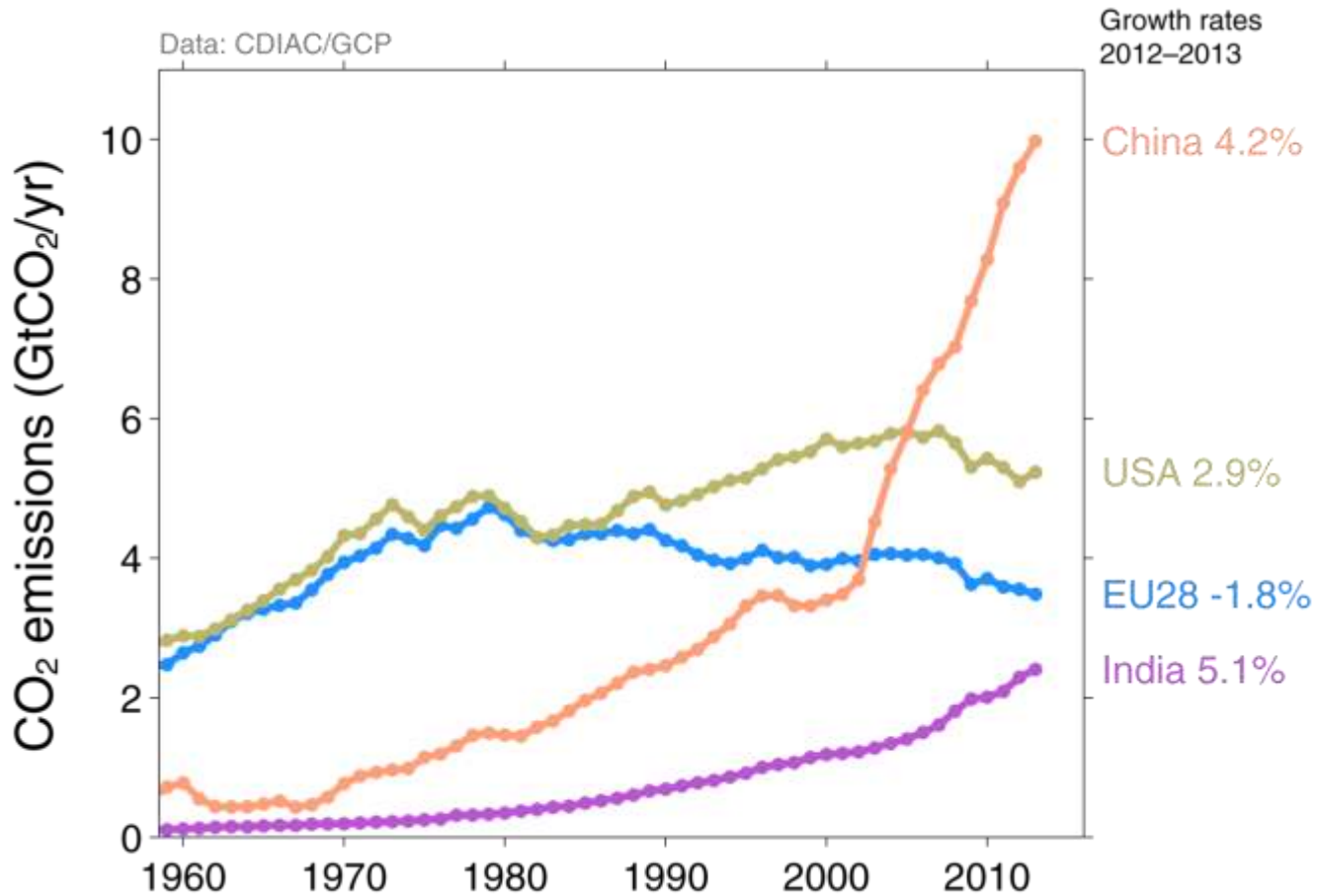


Over 1000 scenarios from the IPCC Fifth Assessment Report are shown

Source: [Fuss et al 2014](#); [CDIAC](#); [Global Carbon Budget 2014](#)

# Top Fossil Fuel Emitters (Absolute)

The top four emitters in 2013 covered 58% of global emissions  
 China (28%), United States (14%), EU28 (10%), India (7%)



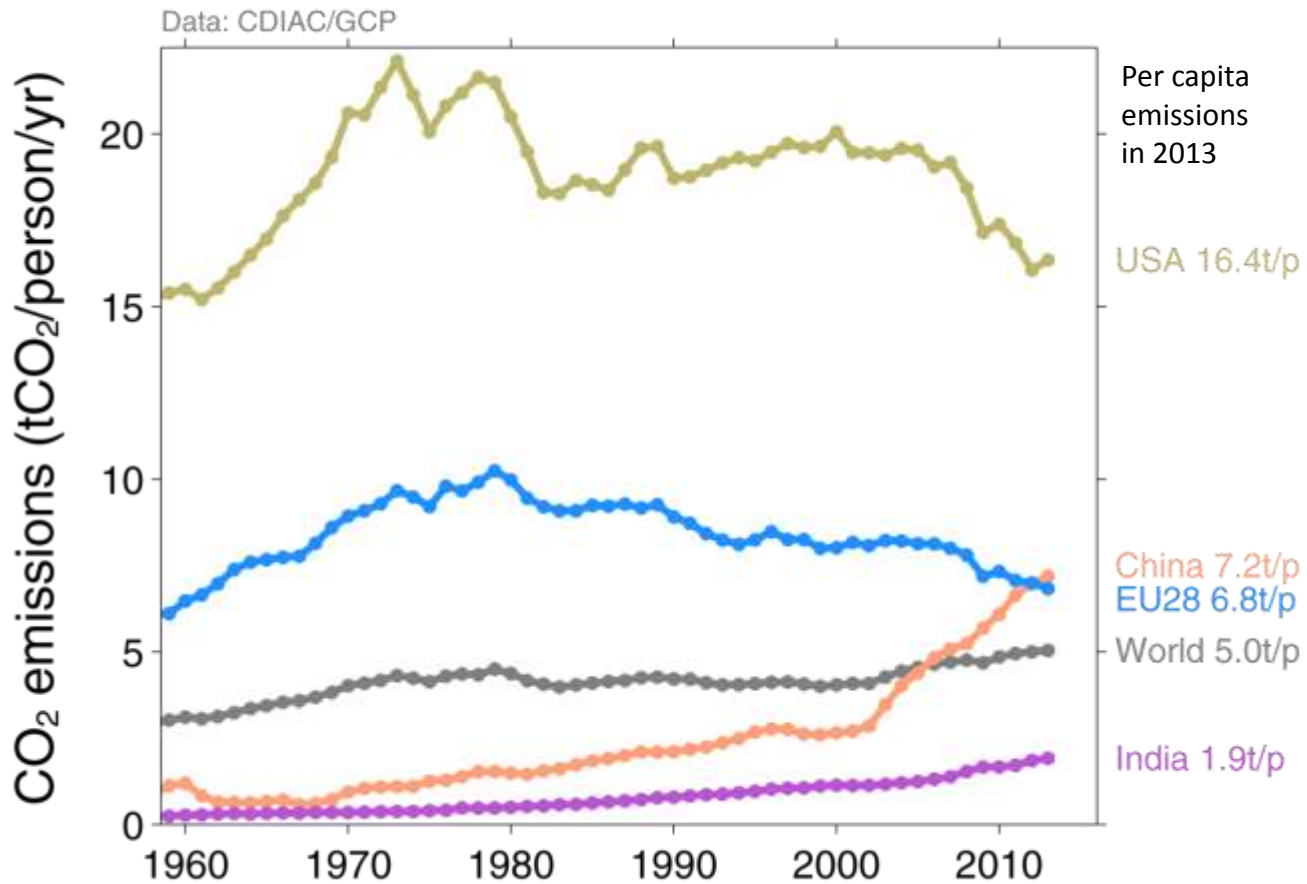
Bunkers fuel used for international transport is 3% of global emissions

Statistical differences between the global estimates and sum of national totals is 3% of global emissions

Source: [CDIAC](#); [Le Quéré et al 2014](#); [Global Carbon Budget 2014](#)

# Top Fossil Fuel Emitters (Per Capita)

China's per capita emissions have passed the EU28 and are 45% above the global average



# Closing the Carbon Budget

# Fate of Anthropogenic CO<sub>2</sub> Emissions (2004-2013 average)

32.4 ± 1.6 GtCO<sub>2</sub>/yr 91%



3.3 ± 1.8 GtCO<sub>2</sub>/yr 9%



15.8 ± 0.4 GtCO<sub>2</sub>/yr 44%



10.6 ± 2.9 GtCO<sub>2</sub>/yr 29%



Calculated as the residual of all other flux components

9.4 ± 1.8 GtCO<sub>2</sub>/yr 26%





# Changes in the Budget over Time

The sinks have continued to grow with increasing emissions, but climate change will affect carbon cycle processes in a way that will exacerbate the increase of CO<sub>2</sub> in the atmosphere

Data: GCP

