The periglacial landforms of Svalbard

Content

- -) ice-wedges
- -) rock glaciers
- -) pingos
- -) solifluction
- -) avalanches
- -) debris flows
- -) rock falls
- -) nivation
- -) aeolian landforms

Glacial landforms
-) surge glaciers

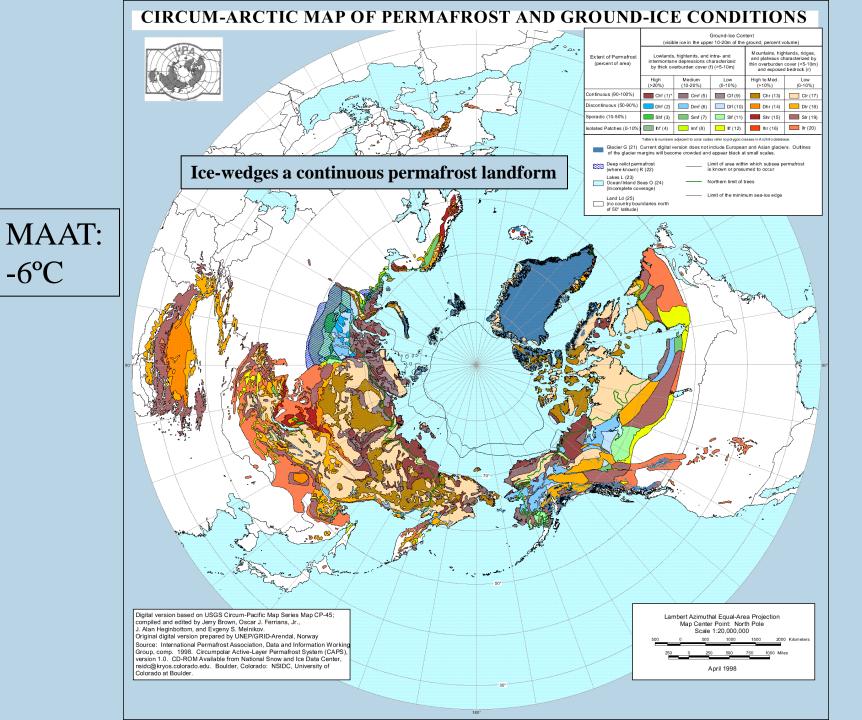
AND CCCH

Hanne H. Christiansen

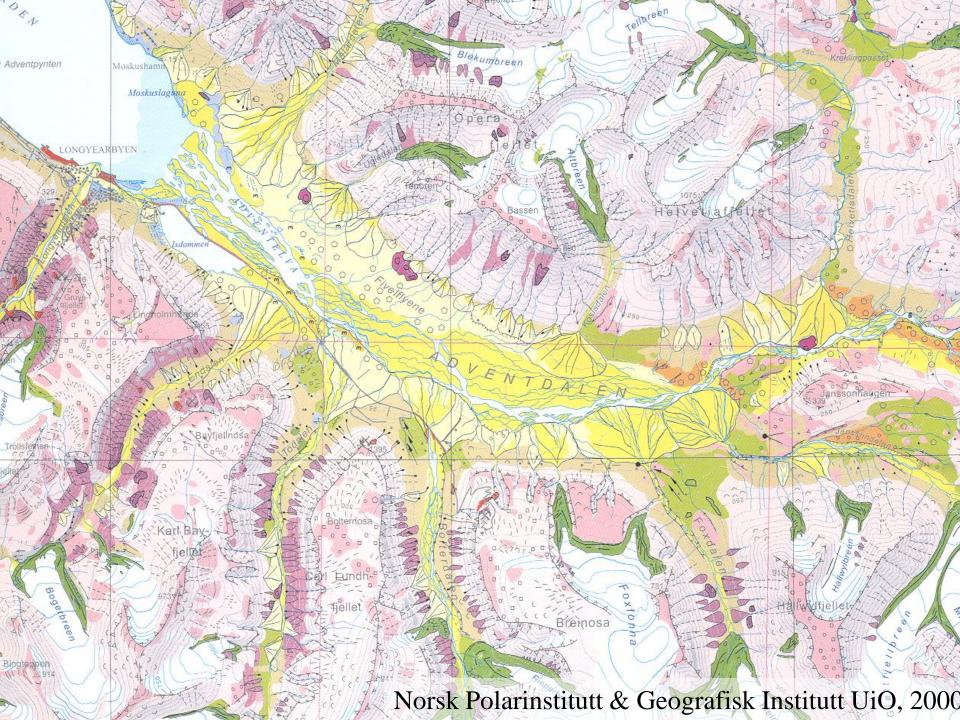
- -) Geology Department, The University Centre in Svalbard, UNIS
- -) Department of Geosciences, University of Oslo



Polygons – ice-wedges



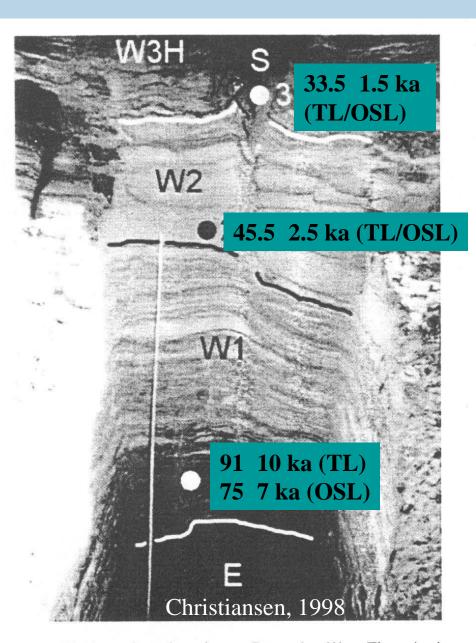
-6°C





Danish ice-wedge cast and sand-wedge





Pingos





Very active pingo in Adventdalen, Svalbard

Cyclic pingo formation

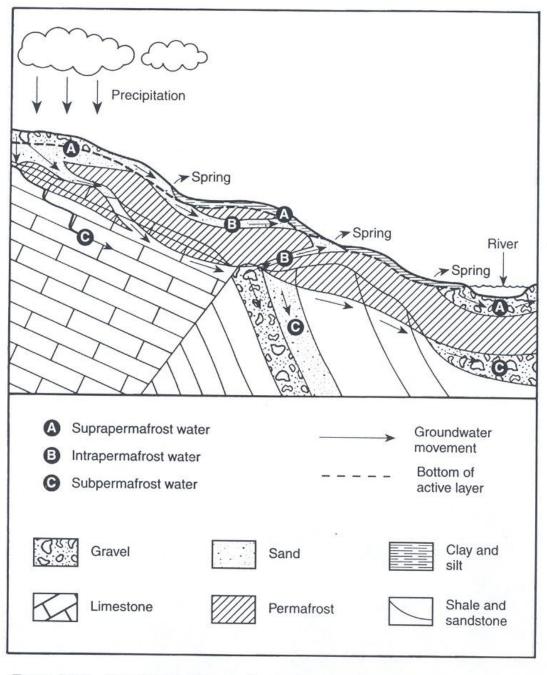
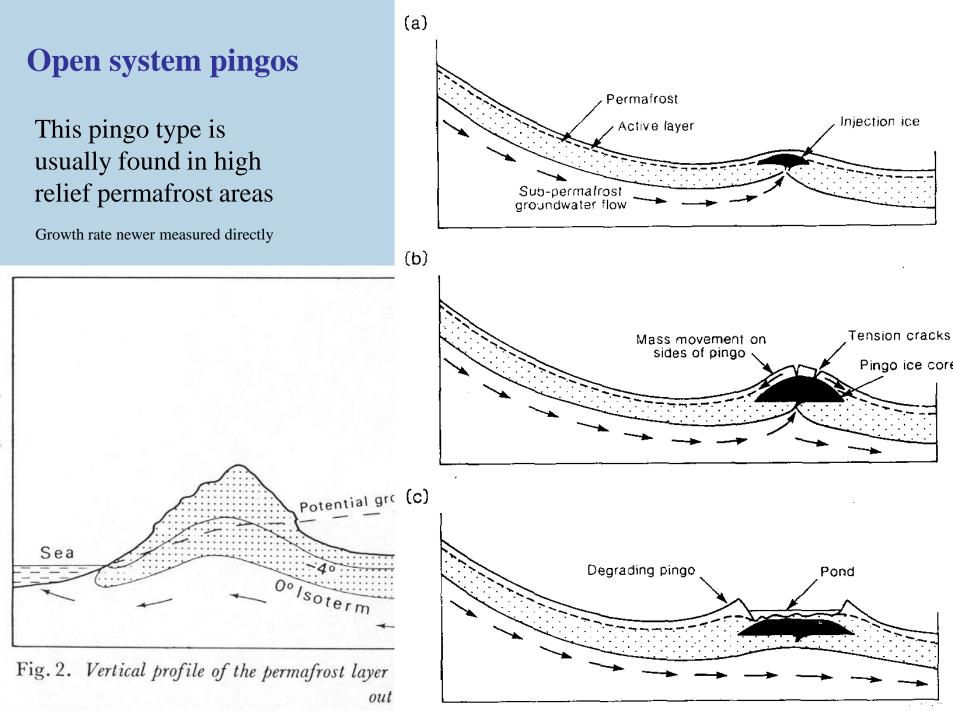


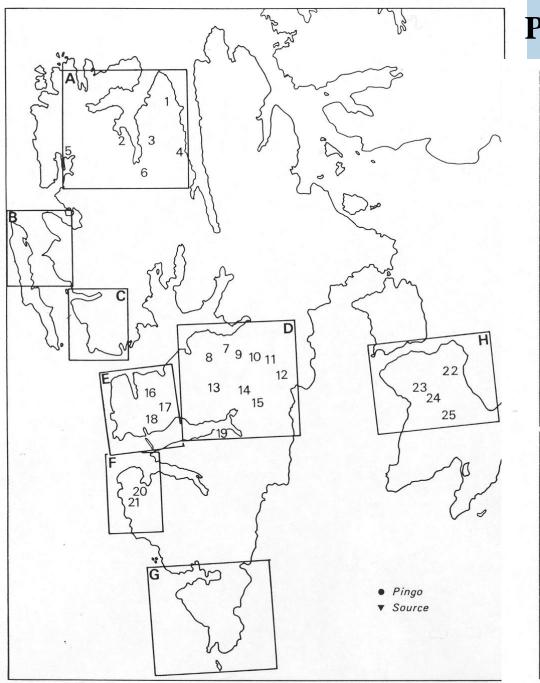
Figure 5.15 Occurrence of groundwater in permafrost areas

Permafrost hydrology

taliks



D



Pingo distribution on Svalbard

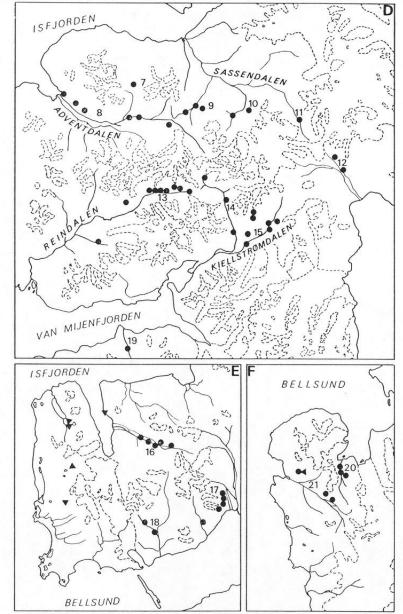


Fig. 4. Outline map of the more detailed maps showing the site of the different pingos and springs described

Fig. 8. Map of the area between Isfjorden, Agardhbukta and Van Mijenfjorden.

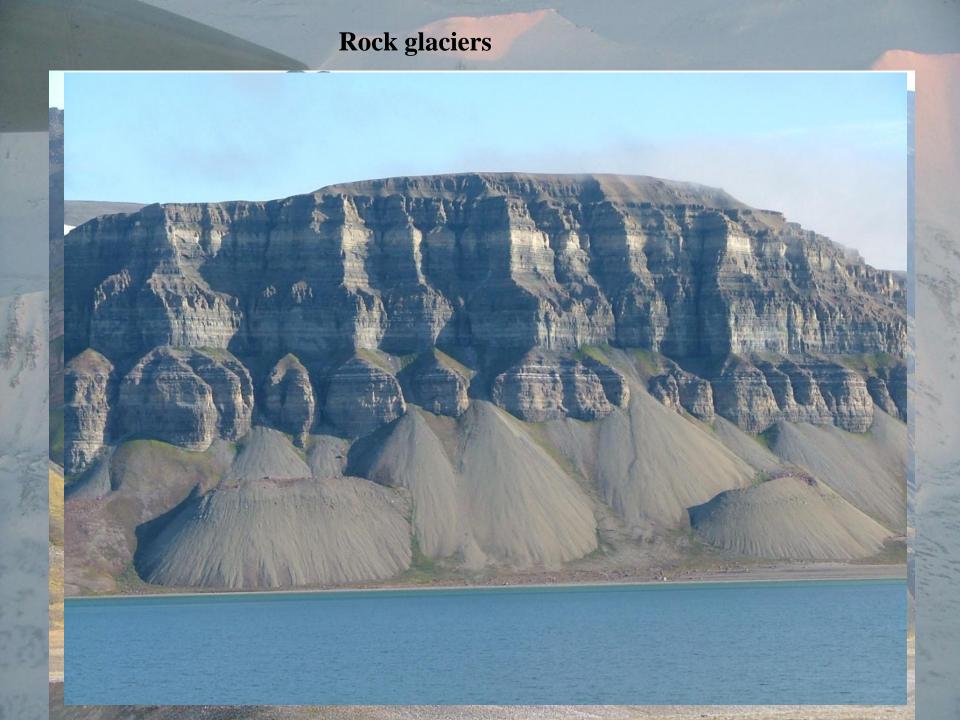
Rock glaciers

Rock glacier: A thick lobate or tongue-shaped body of frozen debris that moves slowly downslope through deformation of internal ice or ice-rich sediment. Moves due to permafrost creep. PF temp. and the amount of ice regulates the speed.

Two types:

Talus rock glaciers (talus-derived or protalus RGs) Glacigenic rock glaciers /glacier derived rock glaciers (debris or morainic RGs)

<u>Activity status:</u> Active (mobile) 0.005-2 m/yr. Inactive (immobile; contains internal ice) Relict (immobile: no perennial internal ice)



Rock glacier characteristics

Steep margins 10-70 m high

Surface microrelief:

closed depressions longitudinal ridges & furrows transverse ridges & depressions.

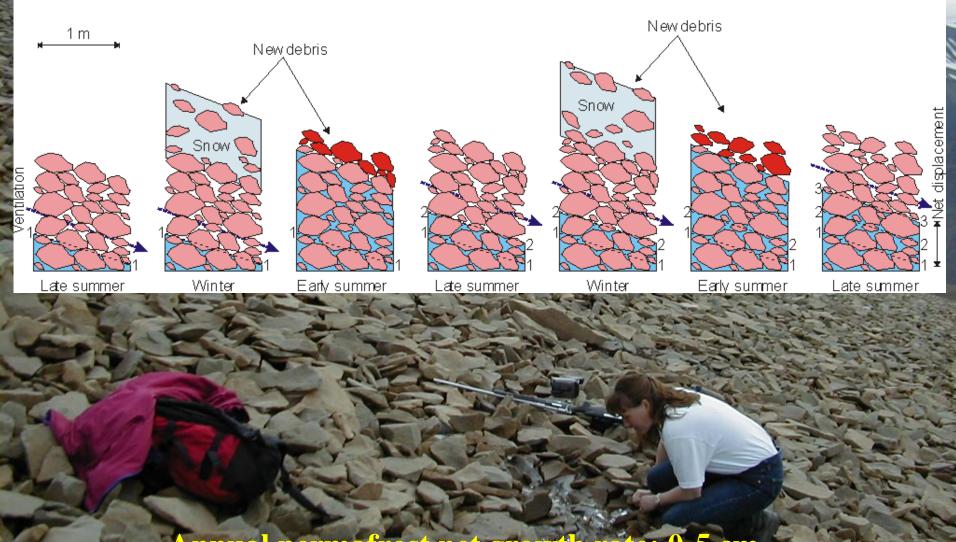
Surface layer of unfrozen coarse debris 1-5 m thick overlies ice-rich sediment and pure ice.

Internal ice: 50-90% of RG volume. Glacigenic RGs: glacier ice. Talus RGs: ice segregation, burial of surface snow and ice.





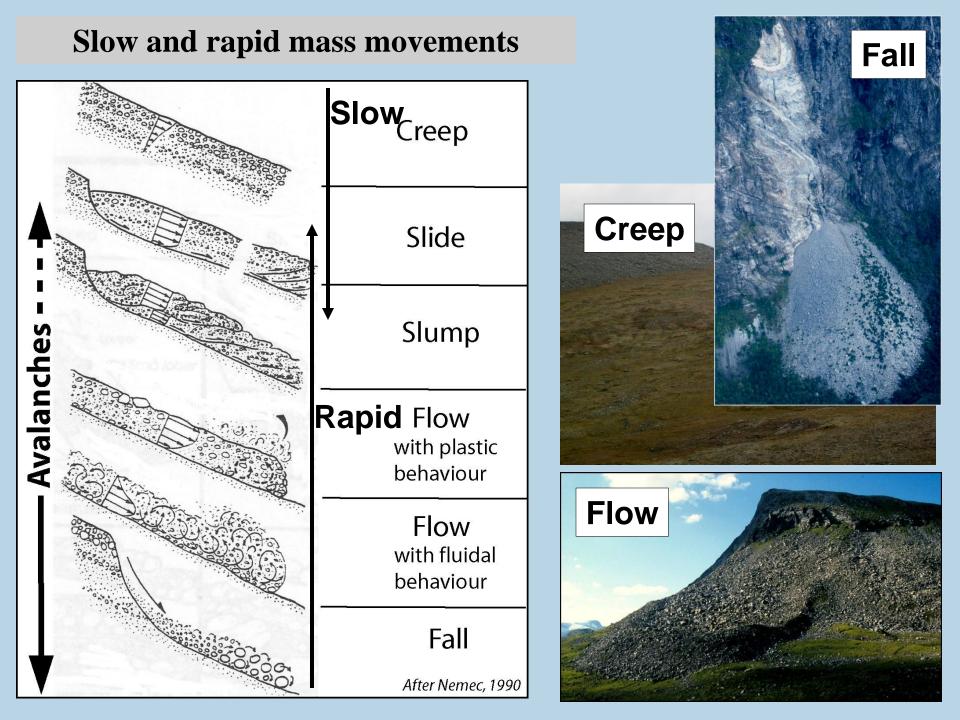
Displacement of active layer and permafrost caused by debris accumulation



Annual permafrost net growth rate: 0-5 cm







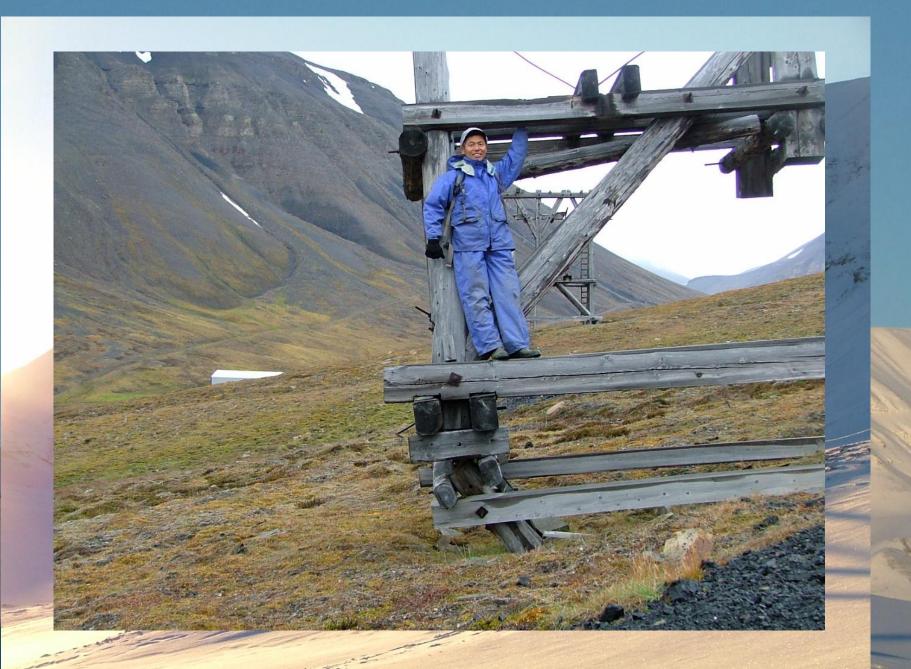
Solifluction lobes

Solifluction small lobes

Solifluction sheets

Solifluction/gelifluction in PF areas occur because of PF + segregated ice lenses providing excess water – reducing internal friction and cohesion in the soil

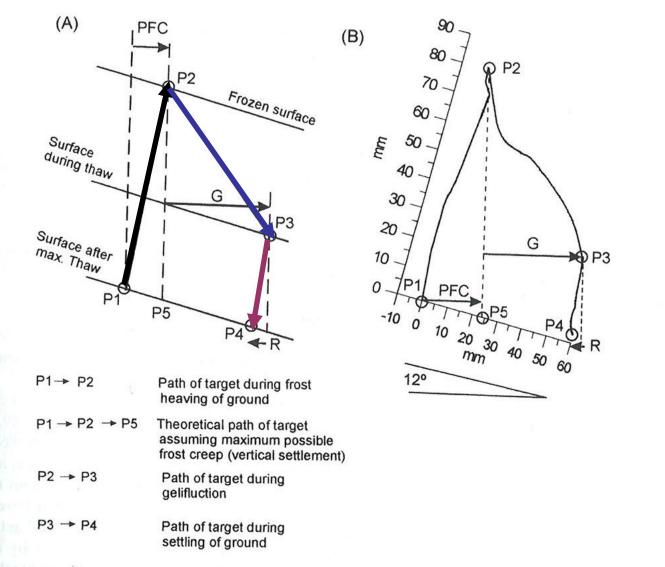
Taubane towers as solifluction monitoring instruments



Old Longyearbyen house piles (tree trunks) registering solifluction



Creep and gelifluction = solifluction



Creep = Heave*tan(slope°)

$\Delta l = h \tan(\sigma)$

Figure 9.7. The components of solifluction. (A) Theoretical displacements of a soil particle due to potential frost creep (PFC), gelifluction (G), and retrograde movement (R). (B) Interpretation of observed surface movement vector in a controlled large-scale laboratory simulation of gelifluction movement at CNRS, Caen, France. Data from Harris and Davies (2000).

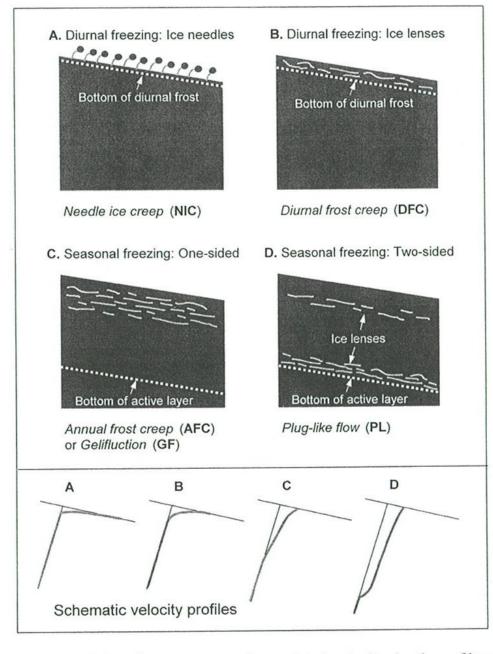


Figure 9.8. Summary of frost heave types and associated velocity-depth profiles of solifluction movement. From Matsuoka (2001c). Reprinted from Earth Science Reviews, with permission from Elsevier.

Solifluction styles

Small-scale slope failures in the active layer, >1 m deep.

Occur when shear strength is reduced, because of increased ice/water content.

Triggers: rapid spring thaw or summer rain

Active layer detachment slides

Debris flows

Erosion area

Statistic di

Transport area

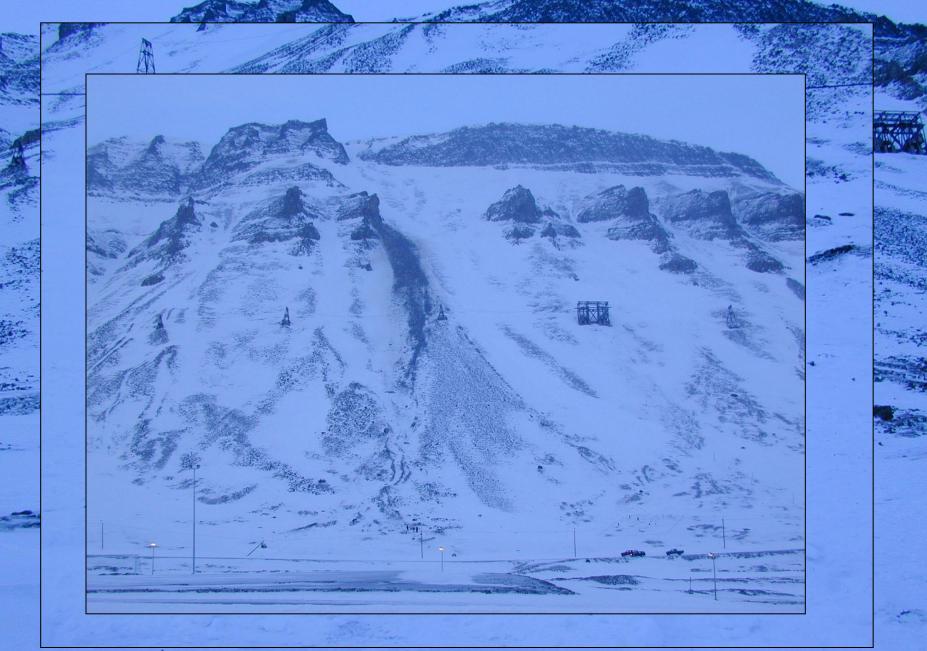
Depositional area

Debris flow June 2003 Longyearbyen





Rockfalls



Free faces

Transport & deposition Weathering:

Thawing of ice or formation of new permafrost

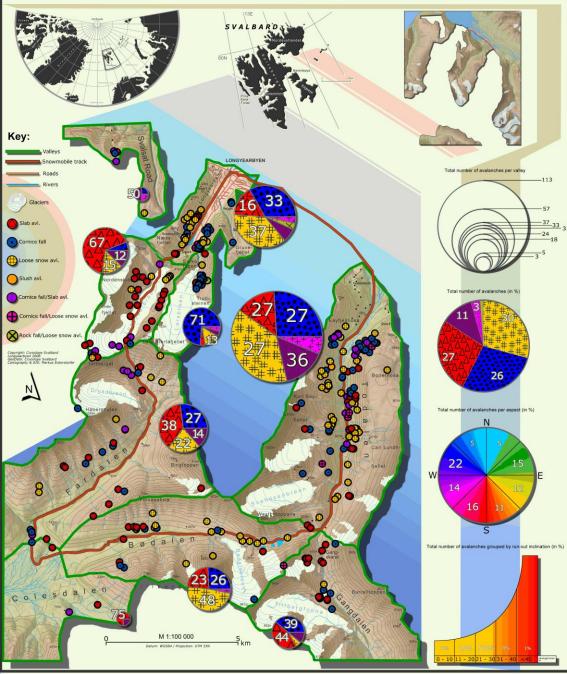
Loosening of rock particles in fractures

Shear strength reduced below stresses from gravity

Traffic and snow avalanches



Avalanches from 01.01.08 to 05.06.08 observed by the CRYOSLOPE SVALBARD PROJECT



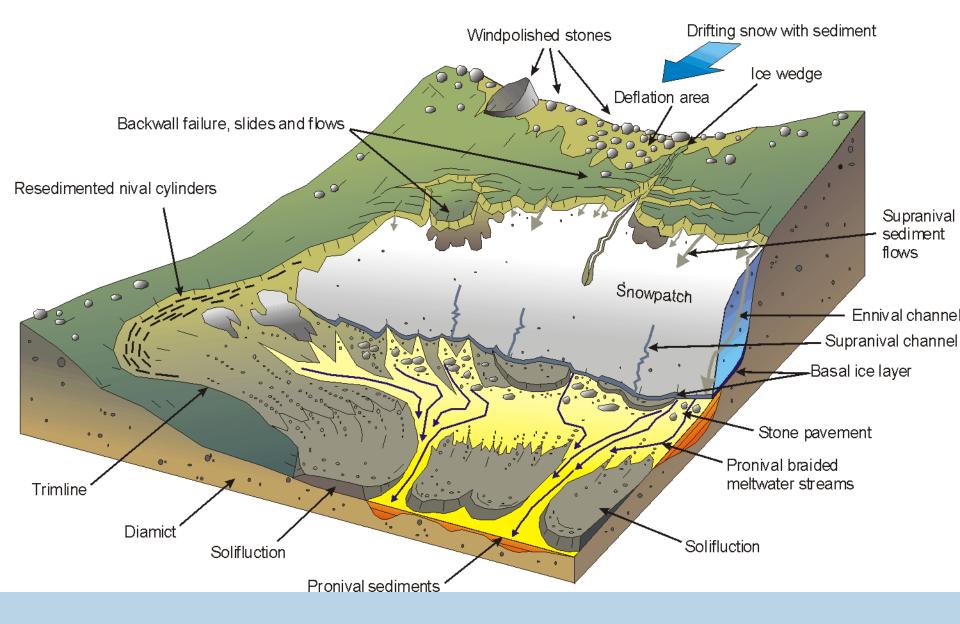
Building the first systematic database with all slope activity to access meteorological control on snow avalanches

Nov more than 600 avalanches since 1 January 2007





High Arctic Nivation Process-Form-Sediment Model



Free rock face / rockwall

Weathering landforms



Wind erosion and transport in periglacial landscapes - Svalbard



Wind transport and deposition in periglacial landscapes - Svalbard



Wind erosion in periglacial landscapes - Svalbard

NIKOI

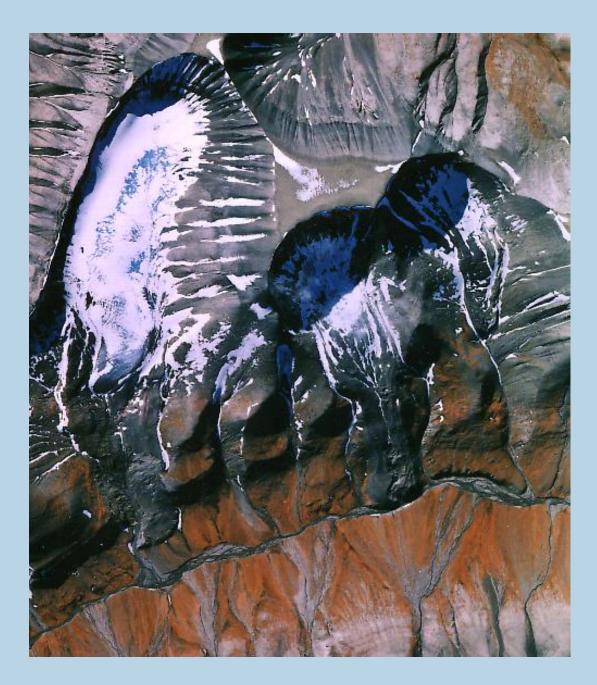
High geomorphic activity in permafrost regions at altitudes near the ELA

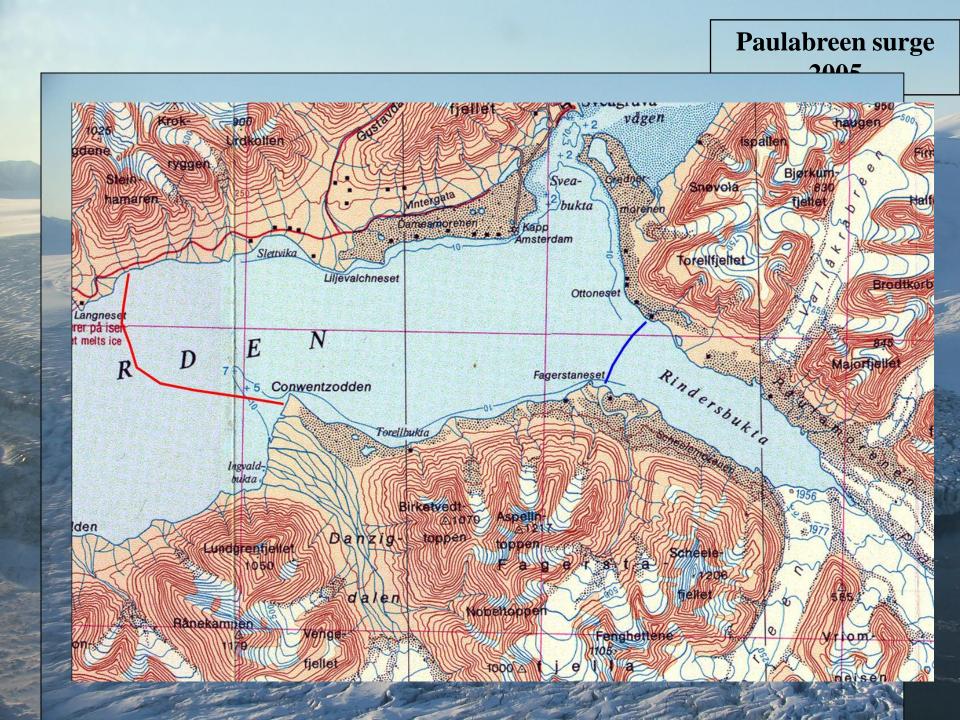
rock falls, avalanches, glaciers, rock glaciers, snowpatches, streams, solifluction, etc.

Site:

Northern side of Hiorthfjellet, near Longyearbyen, Svalbard. North is towards bottom of figure

Figure measures 3000 x 2400 m







Physical Geography at Department of Geology

Investigating modern meteorology at selected landforms Study periglacial processes pingo formation

Quantification of slope processes avalanches Study ice wedges and their palaeoclimatic history



Readings

-) H.M. French The Periglacial Environment, Third edition, 2007. Wiley, 458 p.

-) Humlum, O, Christiansen, H.H. & Juliussen, H. 2008. (2007) Avalanche Derived Rock Glaciers in Svalbard. *Permafrost and Periglacial Processes*, 18, 75-88.

-) Christiansen, H.H. (1998) Periglacial sediments in an Eemian - Weichselian succession at Emmerlev Klev, southwestern Jutland, Denmark. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 138, 245-258.

-) Liestøl, O., 1976: Pingos, springs, and permafrost in Spitsbergen. Norsk Polarinstitut Årbok 1975, pp. 7-29.