

Principles of snow melt

GEO4430 snow hydrology
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How does snow melt?

- We need energy to melt snow/ ice.


$$E = m_s \cdot L_f$$

$$\dots = \rho_w h_{we} \cdot L_f$$

$$h_{we} = \frac{E}{\rho_w L_f}$$

E – energy (J)
 m_s – mass (kg) of snow
 L_f – latent heat of fusion
 = 333400 J kg⁻¹

ρ_w – density (kg m⁻³) of water
 h_{we} – height of snow (m w.e.)



- Energy flux onto a unit surface:

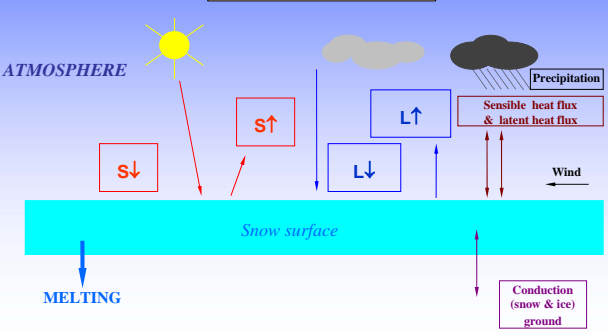
$$Q \text{ (Wm}^{-2}\text{)} = E \text{ (J)} / t \text{ (s)}$$

Amount of energy per unit time

Where does the energy come from?

Energy balance

$$0 = Q_R + Q_H + Q_L + Q_G + Q_P + Q_M$$

$$Q_R = S \downarrow - S \uparrow + L \downarrow - L \uparrow$$


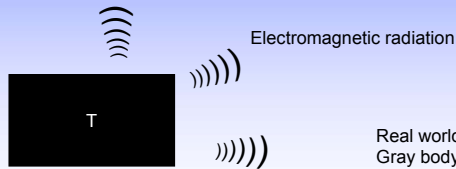
ATMOSPHERE
 Snow surface
 MELTING
 Conduction (snow & ice) ground
 Wind
 Precipitation
 Sensible heat flux & latent heat flux

Radiation

- Black body radiation (Stefan-Boltzmann law)

$$Q = \sigma T^4$$

$$\sigma = 5.6703 \cdot 10^{-8} \text{ W m}^{-2} \text{ K}^{-1}$$



Real world:
 Gray body radiation
 $Q = \epsilon \sigma T^4$
 ϵ – emissivity [0,1]

Radiation

– Basic Principle

- All bodies radiate; as temperature increases, the energy emitted increases, but the wavelength at which the peak radiation is emitted decreases.



Wien's law:

$$\lambda_{\text{max}} = 2.88 \cdot 10^{-3} T^{-1}$$

[m] [m K] [K]

Electromagnetic spectrum

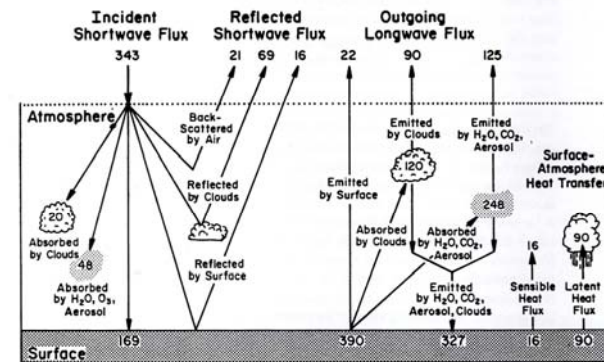
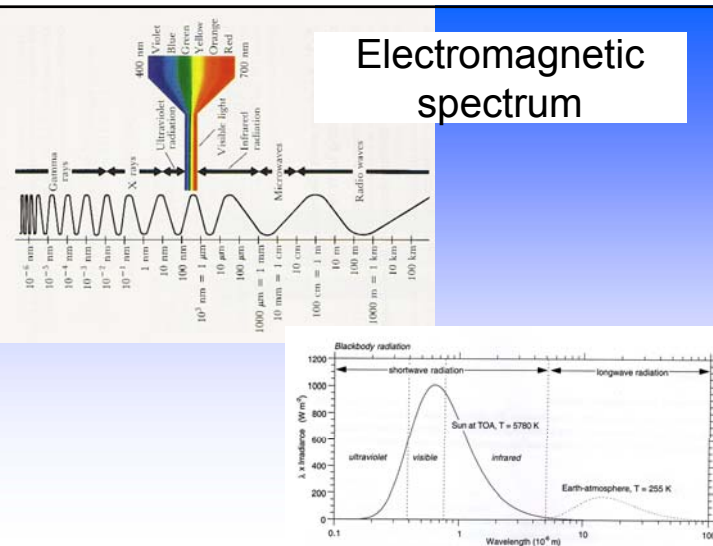


Figure 1.6 Estimated average energy budget of the Earth-atmosphere system (from Liou, 1980).

Radiation

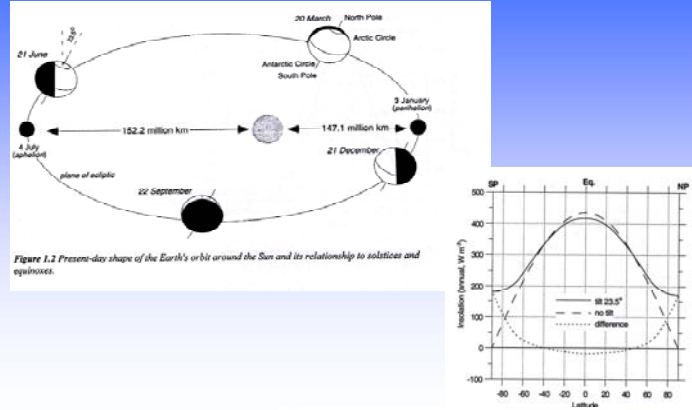


Figure 1.3 Annual mean insolation at the top of the atmosphere with/without tilt of the Earth's axis.

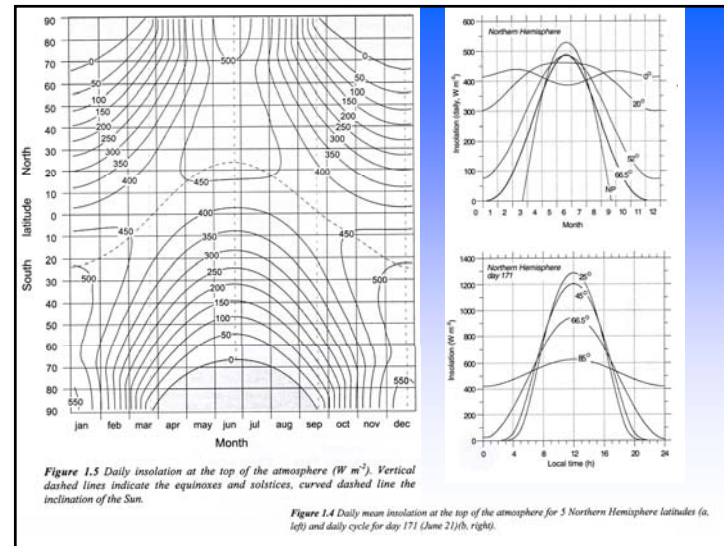
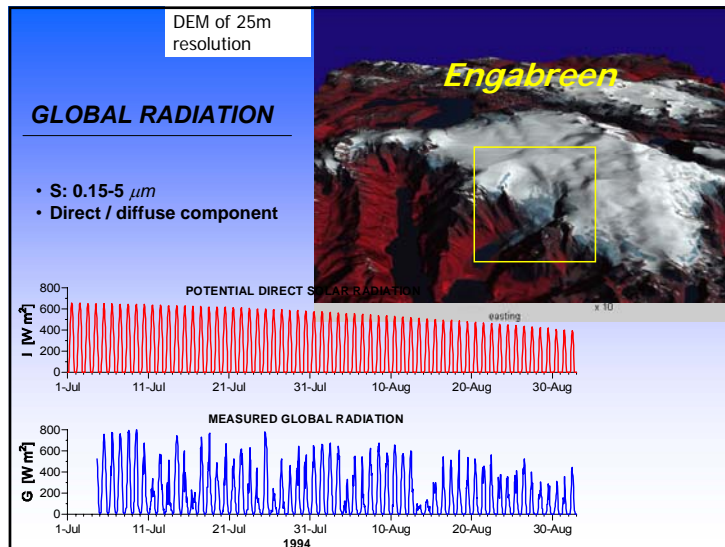


Figure 1.5 Daily insolation at the top of the atmosphere ($W m^{-2}$). Vertical dashed lines indicate the equinoxes and solstices, curved dashed line the inclination of the Sun.

Figure 1.4 Daily mean insolation at the top of the atmosphere for 5 Northern Hemisphere latitudes (a, left) and daily cycle for day 171 (June 21) (b, right).



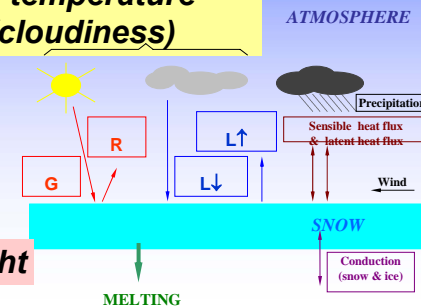
Longwave radiation

- 4-120 μm
- Emitted by atmosphere (water vapour, CO_2 , ozone)
- Function of air temperature and humidity (cloudiness)

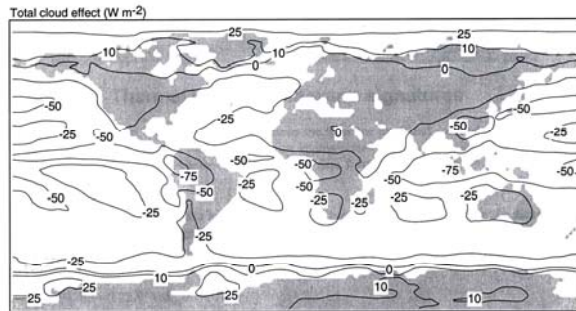
$$L \uparrow = \epsilon \sigma T^4$$

Max = 316 $W m^2$

L acts day & night



Cloud effect:



Effect of clouds on the annual mean net radiation at the surface. The total cloud effect is obtained by comparing the net radiation for clear-sky conditions with those for average conditions, i.e. with clouds.

Why are values positive in polar regions???

reflexion

shortwave reflectance = albedo

new snow	0.75 – 0.95
old snow	0.4 – 0.7
glacier ice	0.3 – 0.45
soil, dark	0.1
grass	0.2
rain forest	0.15

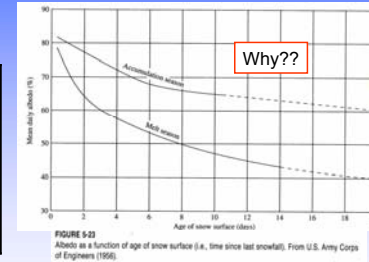


FIGURE 5-23 Albedo as a function of age of snow surface (i.e., time since last snowfall). From U.S. Army Corps of Engineers (1966).

Longwave reflectance of snow: < 0.1

snow is dark on IR image!

reflectance = 1 - emissivity

snow emits a lot $L \uparrow$

Turbulent heat fluxes

Sensible heat flux

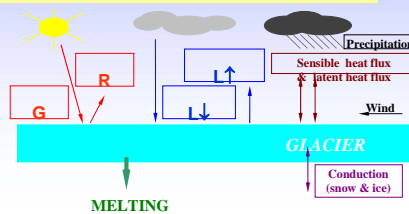
- Function of **temperature** gradient
- Function of **wind** speed

Latent heat flux

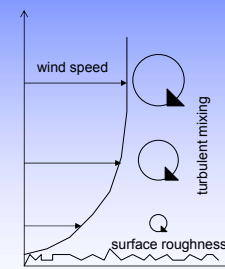
- Function of **vapour pressure** gradient
- Function of **wind** speed

Fluxes also affected by

- Surface roughness
- Atmospheric stability



Turbulent exchange



$$\begin{aligned} \text{Momentum } (\tau) &= \rho_a K_M \frac{\partial u}{\partial z} \\ \text{Heat } (H) &= \rho_a C_p K_H \frac{\partial \theta}{\partial z} \\ \text{Water vapor } (E) &= \rho_a L_v K_W \frac{\partial q}{\partial z} \end{aligned}$$

ρ_a is air density
 u is wind velocity (m s^{-1})
 θ is potential temperature (K)
 q is the specific humidity (dimensionless)
 C_p is the specific heat of air ($1.005 \times 10^3 \text{ J kg}^{-1} \text{ deg K}^{-1}$)
 L_v is the latent heat of vaporization
 K_M is the eddy diffusivity for momentum ($\text{m}^2 \text{ s}^{-1}$)
 K_H is the eddy diffusivity for heat in air ($\text{m}^2 \text{ s}^{-1}$)
 K_W is the eddy diffusivity for water vapor in air ($\text{m}^2 \text{ s}^{-1}$)

Melt physics

- To **melt** 1 kg snow/ice requires 334 000 J kg⁻¹
Latent heat of fusion
- To **sublimate** 1 kg of snow requires 2 600 000 J kg⁻¹
Latent heat of sublimation (8x L_f !!!)
- To **warm** 1 kg of snow 1 K requires 2009 J kg⁻¹ K⁻¹;
ice: 2097 J kg⁻¹ K⁻¹
Specific heat capacity

**Refreezing of 1 g water →
warms 160 g snow by 1 K**

Dry conditions:
Sublimation of snow occurs
 $L_s = 8 \cdot L_f \rightarrow 8x$ less ablation than under wet conditions

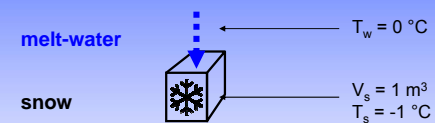


Melt physics

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Specific heat capacity

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Removing cold content



Condition for melt: snow must be at melting temperature, otherwise refreezing will occur

Cold content = energy needed to bring the snow / ice to 0 °C.

In the given example, refreezing of 2.5 l melt-water is needed to compensate for the cold content of the snow pack (snow density, $\rho_s=400$ kg m⁻³).

Cold content

$$Q_{cc} = C_i \rho_s h_s (T_s - T_m)$$

Q_{cc} is the cold content

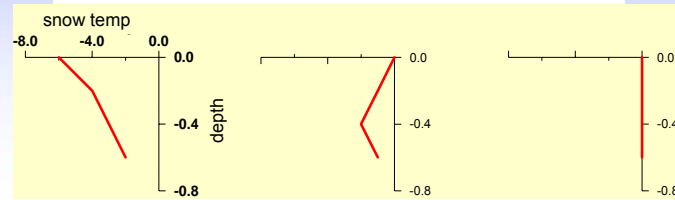
ρ_s is the density of snow

h_s is the snow height

C_i is the specific heat of ice

T_s is average snow temperature in K or C

T_m is the melting temperature of ice in K or C



Precipitation...

Heat Advected by Rain on Snow (Q_p)

– First Case

– **Rainfall on a melting snow pack, where the rain does not freeze**

- $Q_p = 4.2T_r P_r$ (kJ/m²-d)

– where T_r is the temperature of the rain (°C)

– and P_r is the depth of rain (mm/day)

- If $T_r = 2^\circ\text{C}$ and $P_r = 2$ mm, then $Q_p = 16.8$ kJ/m²-d or 0.19 Wm²

– Very small compared to 800 Wm² Incident Solar Radiation!

$$A = C_w \rho_w P (T_r - T_s)$$

A is the advected energy, usually rain on snow, in J m⁻² day⁻¹

C_w is the specific heat of liquid water

ρ_w is the density of liquid water

T_s is average snow temperature in K or C

T_r is the temperature of the rain

P is the precipitation amount (m day⁻¹)

Energy partitioning

TABLE 5-3
Energy-Balance Components (in cal cm⁻²) for Six Seasons at the U.S. National Weather Service—U.S. Agricultural Research Service
Snow Research Station, Danville, VT

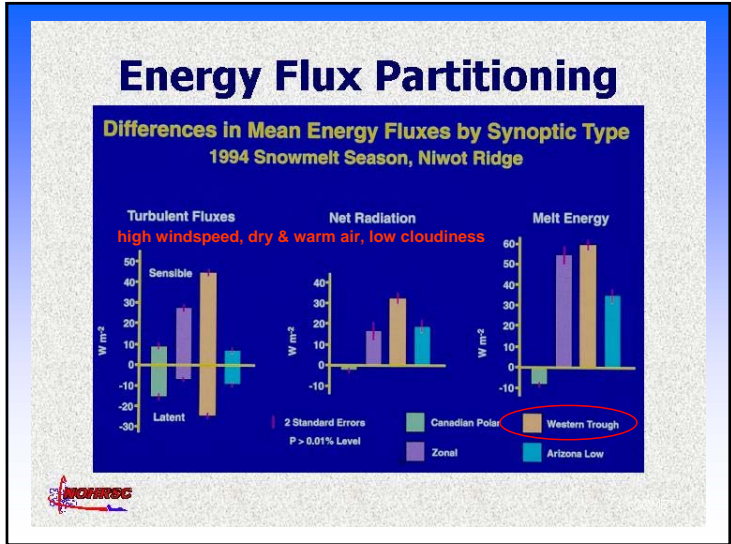
	68-69	69-70	70-71	71-72	72-73	73-74	Average	in [%]	out [%]
Accumulation Season									
Net shortwave radiation, K	4011	4096	4664	4053	2750	4580	4026	58	92
Net long-wave radiation, L	-6317	-6201	-6756	-5707	-4165	-5394	-5757		
Net radiation, $K + L$	-2306	-2105	-2092	-1654	-1415	-814	-1731		
Heat from rain, R	2	20	6	9	17	24	13	0.2	11
Heat from ground, G	735	1287	770	597	628	710	788		
$K + L + R + G$	-1569	-798	-1316	-1048	-770	-80	-930		
Turbulent exchange, sensible, H	2311	2015	2697	2209	1562	1955	2125	31	8
Turbulent exchange, latent, L_E	-286	-618	-648	-584	-384	-487	-501		
$H + L_E$	2025	1397	2049	1625	1178	1468	1624		
Net heat input	456	599	733	577	408	1388	694		
Melt Season									
Net shortwave radiation, K	4585	3081	3889	3600	3575	4031	3794	70	76
Net long-wave radiation, L	-2415	-1612	-1846	-1786	-2208	-2627	-2082		
Net radiation, $K + L$	2170	1469	2043	1814	1367	1404	1711		
Heat from rain, R	11	15	17	30	16	28	20	0.4	3
Heat from ground, G	125	85	121	102	228	222	147		
$K + L + R + G$	2306	1569	2181	1946	1611	1654	1878		
Turbulent exchange, sensible, H	1590	1030	1603	1573	1345	1452	1432	27	24
Turbulent exchange, latent, L_E	-677	-427	-565	-689	-836	-762	-659		
$H + L_E$	913	603	1038	884	509	690	773		
Net heat input	3219	2172	3219	2830	2120	2344	2651		

Energy Flux Partitioning

500 mb Synoptic Weather Patterns, Spring, 1994



Niwot Ridge, Colorado, USA



Energy partitioning (%)

<i>Glacier</i>	Q_R	Q_H	Q_L	Q_G	Q_M
Aletschgletscher, Switzerland	92	8	-6	0	-94
Hintereisferner, Austria	90	10	-2	0	-98
Peytoglacier, Canada	44	48	8	0	-100
Storglaciären, Sweden	66	30	5	-3	-97

- ### Summary
- **Ice and snow melt are determined by the energy balance**
→ Do not necessarily melt at air temperature $\geq 0^\circ\text{C}$
 - **Snow/ice surface temperature must be raised to 0°C before melting can occur (2 steps: warming, melting)**
 - **Fixed maximum surface temperature (0°C)**
→ under melting conditions: constant $L\uparrow = 316\text{ Wm}^2$, surface vapour pressure = 611 Pa
 - **Often net radiation dominant source of energy**
 - **Sublimation reduces energy available for melt**

